



Title	Photosynthesis in sediments determined at high spatial resolution by the use of microelectrodes
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3 **Photosynthesis in sediments determined at high spatial**
4 **resolution by the use of microelectrodes**
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6 Short running title: Photosynthesis in sediments
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8 By
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Abstract

The present study investigated photosynthetic rates and their regulation by light within the upper 5 mm of sediment in a tidal area of Niida River in Hachinohe, Japan. Steady-state concentration profiles of O_2 , NH_4^+ , NO_2^- , H_2S , and pH in the sediment were measured with microelectrodes. Microzonation of O_2 respiration, denitrification and SO_4^{2-} reduction was found in the sediment. When light intensities exceeded $1050 \mu\text{mol photons/m}^2/\text{s}$, net photosynthetic activity was detected in the upper 0.5 mm of the microbial mat colonizing on the sediment surface in the tidal area. In contrast, gross photosynthetic activity was detected in the upper 1.0 mm of the microbial mat at $1900 \mu\text{mol photons/m}^2/\text{s}$. As light intensity increased, the net photosynthetic rate and O_2 penetration depth increased. The maximal net photosynthetic rate and O_2 penetration depth were $6.1 \mu\text{mol } O_2/\text{cm}^3/\text{h}$ and 2.2 mm, respectively, at $1900 \mu\text{mol photons/m}^2/\text{s}$. Net photosynthetic rates in the microbial mat in the tidal area were lower than in the upstream sediment. The analysis of continuous O_2 concentration measurements in different layers of the microbial mat during artificial light-dark cycles demonstrated that the photosynthetic activity response to changes in light intensity was extremely fast (a few seconds) and the O_2 concentration in the microbial mat became stable within 200 s. The measurement of physical and chemical parameters in river water revealed that the study site was relatively polluted and sunlight intensity significantly fluctuated temporally. These results suggested that the in situ microbial processes occurring in the sediment fluctuated in accordance with periodic fluctuations in sunlight intensity.

Key words

Photosynthesis; Light intensity; River sediment; Tidal area; Microelectrodes

Introduction

Photosynthesis significantly affects carbon and nitrogen cycles in river systems (Falkowski and Raven, 1997b). While the primary photosynthetic microorganisms in the sea are planktonic microalgae, in rivers photosynthetic microorganisms in microbial mats,

1 which colonize on sediment surface, and in epilithic biofilms that grow on stones, are
2 significant contributors to photosynthetic activity. The microbial mats are compact
3 microbial communities with high specific rates of metabolic processes (Castenholz, 1994).
4 The high rates are due to dense populations of microorganisms and the high availability of
5 organic carbon and nutrients. Photosynthetic microorganisms in the microbial mats act as
6 sources of organic carbon and oxygen (O₂). In addition, the microbial mats receive organic
7 carbon and nutrients from overlying water.

8 Although the typical thickness of a microbial mat is only a few millimeters, a vertical
9 zonation of the microbial processes can be found in the microbial mat (Jensen *et al.*, 1994;
10 Wieland and Kühl, 2000). The use of microelectrodes has made it possible to study such
11 microzonation of the microbial processes and the tight spatial coupling between them at
12 high spatial resolution. The O₂ respiration process affected nitrification and denitrification
13 in freshwater sediment (Jensen *et al.*, 1994), sulfate (SO₄²⁻) reduction in a hot spring
14 cyanobacterial mat (Wieland and Kühl, 2000), and photosynthesis in hypersaline microbial
15 mats (Gröttschel and de Beer, 2002).

16 In the river systems, oxygenic photosynthesis, followed by increases in O₂ and organic
17 carbon concentrations, affects the microbial processes in the sediment. Although the
18 thickness of the photosynthetic zone is less than one millimeter due to limited light
19 penetration in the sediment, the specific photosynthetic rate is very high. Microelectrode
20 measurements have revealed the occurrence of oxygenic photosynthesis with high
21 photosynthetic rates in cyanobacterial mats (Wieland and Kühl, 2000; Ramsing *et al.*, 2000)
22 and natural biofilms (Glud *et al.*, 1992). Light (Ramsing *et al.*, 2000), O₂ concentration
23 (Gröttschel and de Beer, 2002) and temperature (Wieland and Kühl, 2000) affect the
24 photosynthetic activity in the cyanobacterial mats. However, very little is known about the
25 microbial processes and the effect of environmental factors in sediment at a river mouth.

26 Rivers flowing through urban areas receive anthropogenic inputs of organic carbon
27 and nutrients (e.g. nitrogen and phosphorus) from adjacent land (Chen *et al.*, 1999;
28 Sørensen and Revsbech, 1990). At the river mouth, diverse physical and chemical factors
29 (e.g. sunlight, concentrations of organic carbon and nutrients, changes in flow rates and
30 salinity), as well as nutrient retention occur because of tides (Feng *et al.*, 2002). These
31 environmental factors could affect the microbial photosynthesis, O₂ respiration,

1 nitrification, denitrification, and SO_4^{2-} reduction. Therefore, the microbial processes in the
2 sediment of a tidal area may be spatially and temporally dynamic when compared with other
3 environments.

4 In the present study, we investigated the effect of light intensity on the net
5 photosynthetic activity of sediment located in a tidal area of the Niida River, Hachinohe,
6 Japan. We further discussed the effect of light intensity on microzonation of the microbial
7 processes. Laboratory measurements of the O_2 , NH_4^+ , NO_2^- , H_2S , and pH concentration
8 profiles, and net and gross photosynthetic rates were performed with microelectrodes.
9 Furthermore, river water quality and environmental conditions of the sediments were
10 determined.

12 **2. Materials and methods**

14 *2.1. Sampling*

16 Sediment samples and river water were collected at two points (Point 1 and 2) in the Niida
17 River in Hachinohe, Japan (Fig. 1). The Niida River flows through agricultural areas and a
18 city area in Hachinohe. The population of Hachinohe City is about 245,000. There are
19 several factories of processed marine products around the downstream part of the Niida
20 River. River water samples were collected during April and December 2002. Flow rate at
21 the point 2 ranged from 212 to 2.3 m^3/s during the study period and the average was 15 m^3/s .
22 Maximal tidal range is 130 cm. The point 1 and 2 are located approximately 1.5 and 8.0 km
23 from the river mouth, respectively. Sediment samples were collected using 4 cm diameter
24 Plexiglas tubes with minimal disturbance. The sediment and river water samples were
25 transferred to the laboratory and were analyzed within 12 h using the methods described
26 below.

28 *2.2. Microelectrode measurements*

30 In the laboratory, steady state concentration profiles of O_2 , NH_4^+ , NO_2^- , H_2S , and pH
31 in the sediments were recorded according to the protocol reported elsewhere (Sato *et al.*,

2003). Clark-type microelectrodes for O₂ with a tip diameter of approximately 15 µm and a 90 % response time <0.5 s were prepared and calibrated as described by Revsbech (Revsbech, 1989). LIX-type microelectrodes for NH₄⁺, NO₂⁻ and pH (de Beer *et al.*, 1997) were constructed, calibrated, and used according to a protocol reported elsewhere (Okabe *et al.*, 1999b). H₂S microelectrodes were constructed as described by Revsbech *et al.* (1983) and were calibrated and used as described elsewhere (Okabe *et al.*, 1999a).

A flow cell reactor (60 (L) × 10 (W) × 5 (H) cm) for microelectrode measurements was constructed using an acrylic board (Fig. 2). The sediment samples were positioned in the reactor, where the surface of the microbial mat colonizing on the sediment surface was aligned with the bottom of the flow cell. 5.0 liter of the medium for microelectrode measurements was fed to the reactor at an average liquid velocity of 2 cm/s. To create a parallel, nonturbulent flow, a metallic net was mounted at the upstream of the reactor. The depth of the water above the microbial mat surface was about 1 cm. For the measurements of O₂ concentration profiles in the sediment, raw river water sampled at the sampling points was used as the medium for microelectrode measurements. Light intensity was changed from 10 to 1900 µmol photons/m²/s. When the microprofiles of NH₄⁺, NO₂⁻, H₂S, and pH in the sediment were measured, a synthetic medium was used to avoid interference with the LIX-type microelectrodes. The synthetic medium consisted of NH₄Cl (50 µM), NaNO₂ (40 µM), NaNO₃ (100 µM), Na₂SO₄ (2000 µM), Na₂HPO₄ (3000 µM), MgCl₂•6H₂O (84 µM), CaCl₂ (200 µM), EDTA (270 µM). pH was adjusted to 7.5. Light intensity was adjusted to 10 µmol photons/m²/s. The sediment was then acclimated in the medium at 20°C for at least 30 minutes before measurements to ensure that steady-state profiles were obtained. The microelectrodes were mounted on a motor-driven micromanipulator for accurate positioning and the concentration profiles in the sediment were recorded at intervals of 0.05 to 0.3 mm from the bulk liquid into the sediment (Okabe *et al.*, 1999b). The microbial mat surface was determined by using a dissection microscope (model Stemi 2000; Carl Zeiss). Halogen lamps supplied needed light during the experiments. At least three concentration profiles were measured at different positions in the sediment for each species and set of conditions. When O₂ concentrations were measured during artificial light-dark cycles, the sediment was illuminated at 1900 µmol photons/m²/s for 10 minutes and the electrode was introduced into the sediment. The sediment was then shaded and kept in the dark until the O₂

concentration became stable. The sediment was then illuminated again at 1900 $\mu\text{mol photons/m}^2/\text{s}$ and kept in the light until the O_2 concentration became stable. After two light-dark cycles, the electrode was introduced into deeper layers.

2.3. Calculation of metabolic rates

Net specific rates of O_2 production (P_{net}) or O_2 consumption (C_{net}) in the sediments were calculated from the mean steady state concentration profiles using Fick's second law of diffusion including a production or a consumption term:

$$D_s (d^2 C_s(z)/dz^2) = C_{\text{net}} \text{ or } -P_{\text{net}}$$

where D_s is the effective sediment diffusion coefficient for O_2 and $C_s(z)$ is the O_2 concentration at depth z . The details of this method were described previously by Lorenzen *et al.* (1998).

The D_s was calculated from the free solution molecular diffusion coefficient (D_0) for O_2 and the sediment porosity (ϕ) according to Ullman and Aller (1982):

$$D_s = \phi^2 D_0$$

The D_0 for O_2 ($2.09 \times 10^{-5} \text{ cm}^2/\text{s}$) at 20°C was found in the literature (Andrussow, 1969). Porosity and D_s at relevant depths of the investigated sediments are summarized in Table 2.

Gross photosynthetic rates (P_{gross}) were determined with same O_2 microelectrodes by the light-dark-shift method (Revsbech and Jørgensen, 1983; Revsbech *et al.*, 1981), in which the decrease in O_2 concentration was determined during a 2 s dark period and the P_{gross} at each depth could be calculated from the following equation:

$$P_{\text{gross}} = -dC_s(t)/dt$$

where $dC_s(t)/dt$ is the rate of O_2 disappearance during 2 s after darkening. Theory of this method was described elsewhere (Revsbech *et al.*, 1981). A short summary is given here. A steady state O_2 profile in the sediment will be approached after a long period of illumination. At steady state, the photosynthetic O_2 production in any layer will balance the combined losses due to respiration and to diffusion of O_2 away from that layer (calculated negative if there is a net diffusion to the layer). When illumination is suddenly stopped, diffusion and respiration are initially unchanged, so that the O_2 concentration in the layer will decrease at a rate equal to the former photosynthetic rate. Furthermore, specific O_2 respiration rates

(R_{O2}) were calculated from these rates by subtracting the P_{net} from the P_{gross} for each layer of the sediment.

2.4. Analytical methods

Chemical Oxygen Demand (COD), using KMnO₄ as an oxidant, Biochemical Oxygen Demand (BOD) and Dissolved Organic Carbon (DOC) were analyzed according to the Standard Methods (APHA, 1998). The NH₄⁺ and NO₂⁻ concentrations were determined colorimetrically. The NO₃⁻ and SO₄²⁻ concentration was determined using an ion chromatograph (HIC-6A; Shimadzu) equipped with a Shim-pack IC-A1 column. The samples for NH₄⁺, NO₂⁻, NO₃⁻ and SO₄²⁻ were filtered with 0.2 µm membrane filters before analysis. Porosity of the sediment was analyzed according to the methods of Murdroch and Azcue (1995). Ignition loss was measured from the loss of volatile matter upon ignition at 600°C for 2 h. Total iron (T-Fe) concentration in the sediment was analyzed using an atomic adsorption spectrophotometer (AA-670; Shimadzu) after a proper sample pretreatment (APHA, 1998). Light intensity was measured by a quantum meter (Fujiwara Scientific Company), which senses only in the 400 to 700 nm region. All measurements were made above the water surface. Salinity was measured using a salinometer. The O₂ concentration and pH were determined using an O₂ electrode and a pH electrode, respectively.

3. Results and discussion

3.1. Physical and chemical parameters in river water and sediments at the study sites

Physical and chemical parameters (average ± standard deviation) of the river water (n=57) at the point 1 are shown in Table 1. The river water quality at the point 1 was relatively polluted while BOD was low and the NH₄⁺ concentration was below the detection limit at the point 2 (data not shown). The average NO₃⁻ concentration was 127 µM at the point 1. SO₄²⁻ concentrations and salinity fluctuated, indicating that the point 1 was located in a tidal area. The sediment located at the point 1 was exposed to atmosphere for several hours during the spring tide. The sediment at the point 1 consisted of sand and the sediment

at the point 2 consisted of silt (Murdroch and Azcue, 1995). A ca. 1 mm thick brownish microbial mat colonized on the sediment surface at the point 1. A black layer underlay the microbial mat, indicating the precipitation of iron sulfide (FeS) in the sediment. Average values of porosity and D_s for O_2 , which was calculated from the molecular diffusion coefficient for O_2 in water and the average value of porosity, at relevant depths of the investigated sediments are summarized in Table 2. The average value of porosity ($n=6$) slightly decreased with depth. Ignition loss in the upper 5 mm of the sediments at the point 1 and point 2 ($n=6$) was $4.8 \pm 0.5 \%$ and $3.4 \pm 0.5 \%$, respectively. Benthic animals (*Tylorrhynchus heterochaetus*) were found at a density of about 2,000 individuals/m² in the sediment at the point 1. The average concentration profile of T-Fe in the sediment at the point 1 ($n=6$) is shown in Fig. 3. The average T-Fe concentration was 21 mg/g dry weight in the microbial mat. The T-Fe concentration slightly increased with depth and the maximum concentration was 25 ± 12 mg/g at a depth of 4-5 mm.

The diurnal cycle of sunlight intensity on the surface of the sediment at the point 1 in June 2003 is shown in Fig. 4. The light intensity increased as the sun rose and reached its maximum ($1920 \mu\text{mol photons/m}^2/\text{s}$) just before noon on sunny days. The light intensity then gradually decreased during the afternoon. The light intensities on cloudy days showed the same trend, but were less than half of the light intensities on sunny days. Even on sunny days, when clouds shaded the sediment surface, the light intensity decreased to the level of cloudy days within a few seconds.

3.2. Concentration profiles in sediments

Representative 2-D contour plots of O_2 in the sediment at the point 1 at $10 \mu\text{mol photons/m}^2/\text{s}$, constructed from 11 vertical O_2 profiles, are shown in Fig. 5. The sediment sample was collected on 6 June 2001. The O_2 concentration in the bulk liquid was ca. $130 \mu\text{M}$. O_2 decreased in the diffusive boundary layer above the microbial mat. The O_2 concentrations at the microbial mat surface were almost the same ($90 \mu\text{M}$) at all measuring points. The contour lines of 10, 30 and $60 \mu\text{M}$ of O_2 followed the microbial mat surface.

Mean steady-state concentration profiles of O_2 , NH_4^+ , NO_2^- , H_2S , and pH in the sediment at the point 1 at $10 \mu\text{mol photons/m}^2/\text{s}$ are shown in Fig. 6. The sediment sample

was collected on 23 January 2002. O_2 penetrated 1.3 mm into the microbial mat. The NH_4^+ produced in deeper parts of the sediment diffused toward the microbial mat and was consumed in the microbial mat. NO_2^- concentration decreased in the upper parts of the oxic zone due to nitrification. NO_2^- concentration profile showed a peak of ca. 150 μM at a depth of 1.0 mm. Decreased NO_2^- concentrations in the lower parts of the oxic zone and anoxic sediment were probably due to denitrification. H_2S was detected in the sediment below a depth of 3.0 mm, indicating SO_4^{2-} reduction occurred in this zone. These results clearly indicate the microzonation of O_2 respiration, denitrification and SO_4^{2-} reduction in the upper 5 mm of the microbial mat and sediment. Decrease in H_2S concentration at a depth of 2.7 mm could be explained by H_2S oxidation by NO_2^- and NO_3^- and H_2S fixation with Fe. The abundance of T-Fe in the deeper parts of the sediment (Fig. 3) reflected this result.

3.3. Effect of light intensity on O_2 concentration profiles and photosynthesis in the sediments

The effects of oxygenic photosynthesis on O_2 concentration in the sediments at the point 1 were investigated because O_2 concentration affected the microzonation of O_2 respiration, denitrification and SO_4^{2-} reduction (Jensen *et al.*, 1994; Wieland and Kühl, 2000). Mean steady-state O_2 concentration profiles measured at five different light intensities and calculated net rates of O_2 production (P_{net}) and O_2 consumption (C_{net}) in the sediment are shown in Fig. 7A to 7E. The sediment sample was collected on 1 July 2003. The photosynthetic activities were detected in the upper 0.5 mm of the microbial mat when light intensities were more than 1050 μmol photons/ m^2/s . As light intensity increased, the P_{net} in the microbial mat increased, resulting in increases in O_2 penetration depth and maximal O_2 concentration in the sediment. The maximal O_2 concentration in the microbial mat increased to 240 μM at a depth of 0.2 mm at 1900 μmol photons/ m^2/s , which was 1.4 times higher than that of the bulk liquid. In this layer the P_{net} was 6.1 μmol $O_2/cm^3/h$. The average O_2 penetration depth at 1900 μmol photons/ m^2/s was 2.2 mm, which was about three times deeper than in the sediment at 10 μmol photons/ m^2/s (Fig. 7A). These results indicate that in nature, the O_2 penetration depth in the sediment could change from ca. 0.7 mm to ca. 2.2 mm in accordance with fluctuations in sunlight intensity (Fig. 4).

Increase in O₂ penetration depth results in higher redox potential in the upper parts of the sediment (Jørgensen *et al.*, 1979) and a shift in the zones of anoxic microbial processes (i.e. denitrification and SO₄²⁻ reduction) to the deeper layers of the sediment (Jensen *et al.*, 1994; Wieland and Köhl, 2000). In this study, microzonation of O₂ respiration, denitrification and SO₄²⁻ reduction was found in the sediment (Fig. 6). Consequently, the locations of the zones of anoxic microbial processes should fluctuate in accordance with a diurnal cycle of sunlight intensity. In the deeper parts of the sediment, the metabolic rates were low because biodegradable organic carbon and nutrients were usually limited due to the long diffusional distance from overlying water, although limited nutrients were supplied from below. Therefore, denitrification and SO₄²⁻ reduction rates might be higher during the night as compared with in the daytime. These results indicate that fluctuations in sunlight intensity might indirectly affect the carbon and nitrogen cycles in the sediment. However, many physical and chemical factors (e.g. vertical hydraulic flow within pores of the sediment induced by tide, water turbulence on the sediment surface, and fluctuations in organic carbon and nutrient concentrations in river water) affect O₂ distribution in the sediment in nature. Consequently, our results present the mechanism of change in O₂ distribution in the sediment by light rather than O₂ concentration profiles in the sediment in nature.

C_{net} was high just below the net O₂ production zone. With increasing light intensity, the maximal C_{net} increased; the C_{net}s were 0.7, 1.3 and 1.8 µmol O₂/cm³/h at 1050, 1550 and 1900 µmol photons/m²/s, respectively. This was probably due to increasing concentrations of O₂ and organic carbon provided from photosynthetic zone in this zone.

The profiles of the average gross photosynthetic rate (P_{gross}) and O₂ respiration rate (R_{O2}) in the sediment at the point 1 at 1900 µmol photons/m²/s are shown in Fig. 8. The sediment sample was collected on 16 July 2003. The R_{O2} was calculated by subtracting the P_{net} shown in Fig. 7E from the P_{gross}. The P_{gross} peaked at a depth of 0.2 mm with the rate of 8.6 µmol O₂/cm³/h and then decreased gradually with depth. Gross photosynthesis occurred from the surface to a depth of 1.0 mm, while the net photosynthesis was detected in the upper 0.5 mm of the microbial mat (Fig. 7E). This indicates that the photic zone ranged from the surface to a depth of 1.0 mm. The R_{O2}s in the photic zone were below 2.9 µmol O₂/cm³/h. These values were comparable with the C_{net}s in the sediment at 10 µmol photons/m²/s (Fig.

7A).

Representative steady-state O₂ concentration profiles measured at five different light intensities and P_{net} and C_{net} in the sediment located at the point 2 are shown in Fig. 9A to 9E. The sediment sample was collected on 9 September 2001. The P_{net} in the sediment at the point 2 was higher than in the sediment at the point 1 (Fig. 7). In the sediment photosynthesis occurred even at 100 μmol photons/m²/s (Fig. 9B). The P_{net} (6.6 μmol O₂/cm³/h) at 100 μmol photons/m²/s was highest at a depth of 0.1 mm in the sediment. This value was higher than in the sediment at the point 1 at 1900 μmol photons/m²/s (Fig. 7E). With increasing light intensity, the maximal P_{net} in the sediment at the point 2 increased. The P_{net} increased to 13.2 μmol O₂/cm³/h at a depth of 0.1 mm at 1250 μmol photons/m²/s. The difference in the photosynthetic rates between these two types of sediment may be attributed to difference in photosynthetic microorganisms between the sampling sites, although the photosynthetic microorganisms in the sediments were not investigated.

Continuous readings of O₂ concentrations in different layers of the sediment at the point 1 during artificial light-dark cycles are shown in Fig. 10A to 10D. The sediment sample was collected on 1 August 2003. At depths of 0.2, 0.4 and 0.6 mm, the O₂ concentration immediately decreased after darkening (0 s) and became stable within 200 s. Thereafter, when the sediment was illuminated again, the O₂ concentration increased immediately at depths of 0.2 and 0.4 mm, however, a small lag (a few seconds) occurred before O₂ started to increase at a depth of 0.6 mm (Fig. 10D). The steady-state O₂ concentration in the light gradually decreased as the light-dark cycle progressed. This was significant in the deeper layers, that is, in the later measurements. The possible explanations for a decrease in the maximal O₂ concentration might be an inherent circadian rhythm of the photosynthetic microorganisms (Falkowski and Raven, 1997a), migration of photosynthetic microorganisms into deeper layers of the sediment (Ramsing *et al.*, 2000), and/or increased photorespiration (Glud *et al.*, 1992).

Conclusions

Photosynthetic activities in the sediment in the tidal area of Niida River in Hachinohe, Japan, were determined at high spatial resolution by microelectrode technique.

1 Photosynthesis occurred in the upper 0.5 mm of the sediment. Microelectrode
2 measurements revealed that a vertical microzonation of O₂ respiration, denitrification and
3 SO₄²⁻ reduction was found in the upper 5 mm of the sediment. As light intensity increased,
4 the net photosynthetic rate and O₂ penetration depth in the sediment increased.
5 Consequently, the location and activities of anoxic microbial processes occurring in the
6 sediment could fluctuate in accordance with periodic fluctuations in sunlight intensity.

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COD (mg/L)	BOD (mg/L)	DOC (mg/L)	NH ₄ ⁺ (μM)	NO ₂ ⁻ (μM)	NO ₃ ⁻ (μM)
16.8 ± 8.2	4.4 ± 2.0	11.2 ± 7.3	44 ± 39	1 ± 5	127 ± 168
O ₂ (μM)	SO ₄ ²⁻ (μM)	Salinity (%)	Temperature (°C)	pH (-)	
190 ± 40	3400 ± 3500	0.23 ± 0.24	17.4 ± 3.9	7.3 ± 0.8	

Table 1
Photosynthesis in sediment determined at high spatial resolution by the use of microelectrodes
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Sample	Distance from the surface (mm)	Porosity (-)	Ds ($10^{-5} \text{ cm}^2/\text{s}$)
Point 1	0-5	0.82 ± 0.09	1.40
	5-10	0.82 ± 0.07	1.41
	10-15	0.79 ± 0.06	1.31
	15-20	0.79 ± 0.08	1.30
	20-30	0.78 ± 0.07	1.27
	30-40	0.76 ± 0.09	1.20
	40-50	0.77 ± 0.11	1.23
	50-60	0.75 ± 0.11	1.18
Point 2	0-50	0.61 ± 0.04	0.78

Table 2

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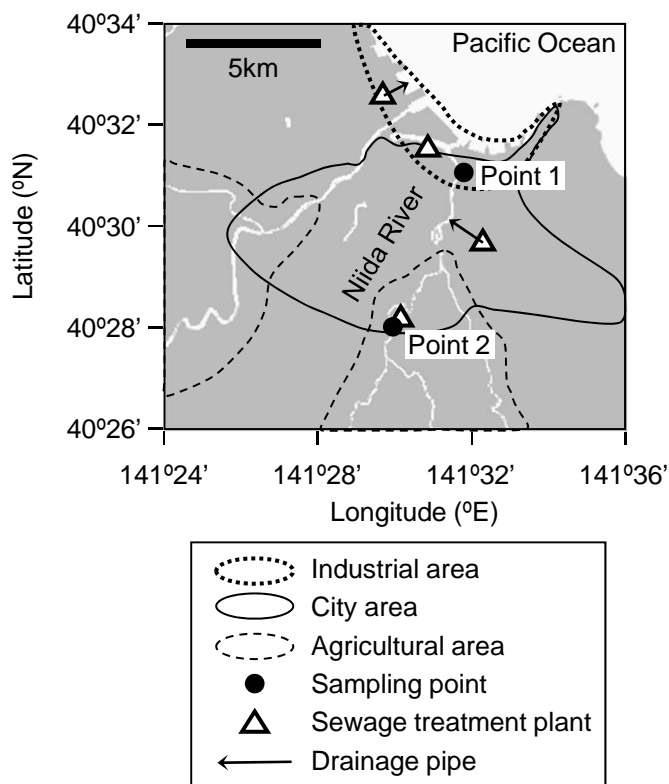


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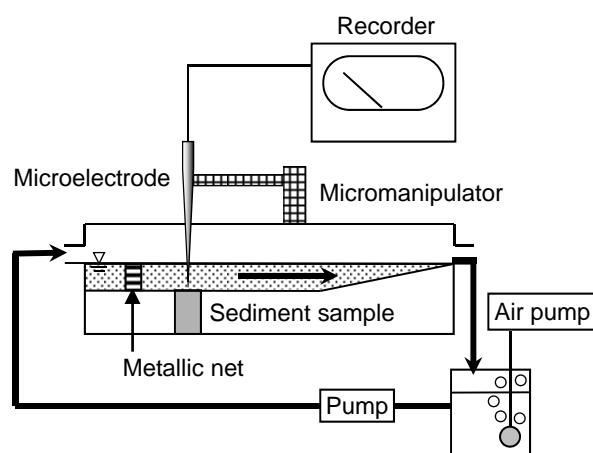


Fig. 2

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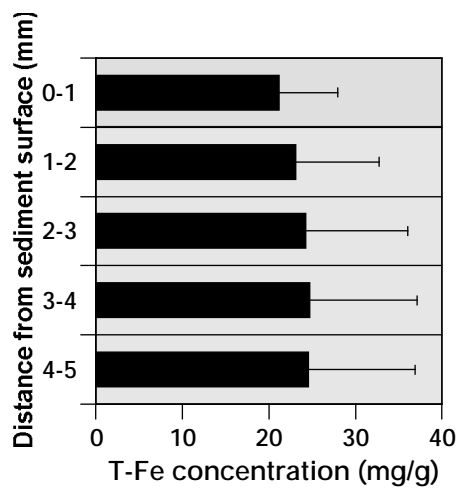


Fig. 3
Photosynthesis in sediment determined at high spatial resolution by the use of microelectrodes
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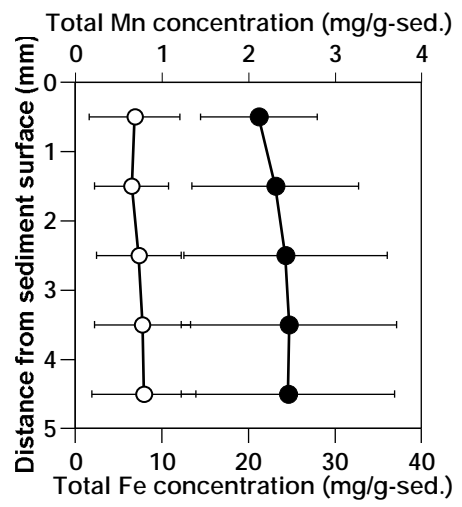
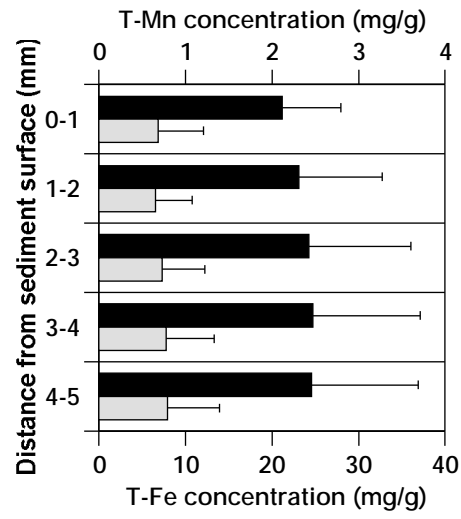


Fig.*
 Macroscale.
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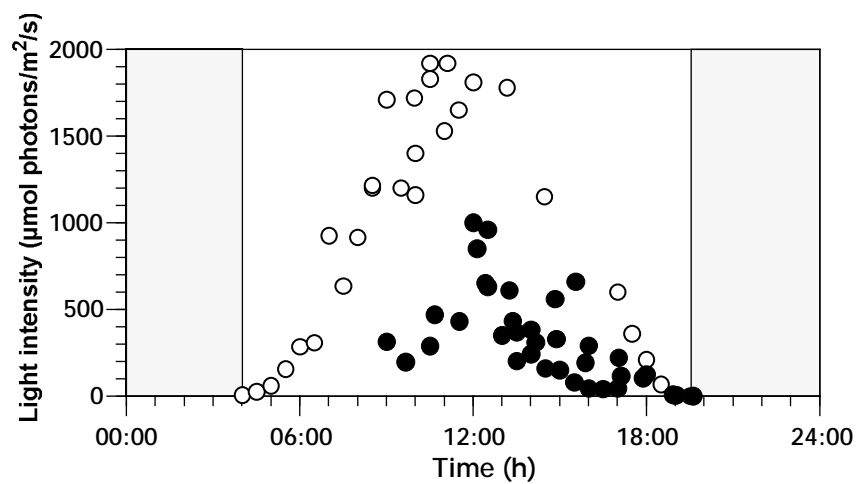


Fig. 4
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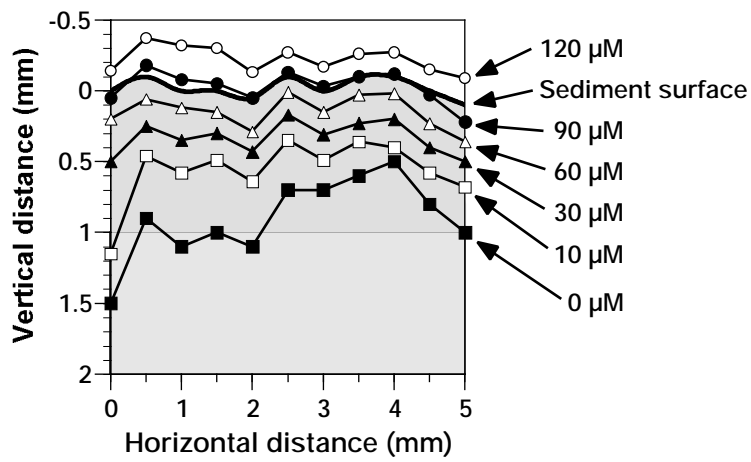


Fig. 5
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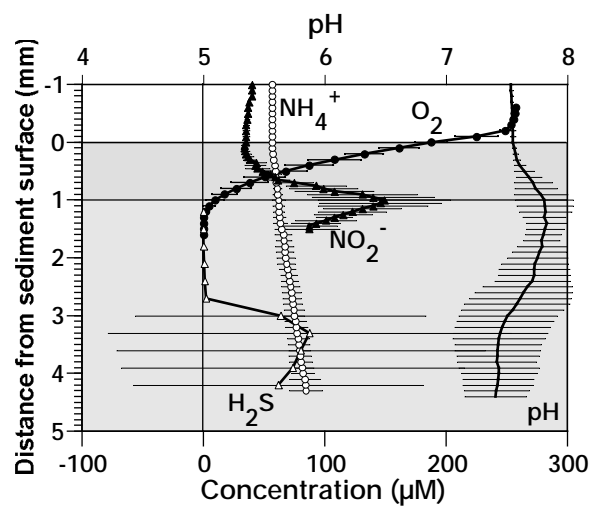


Fig. 6
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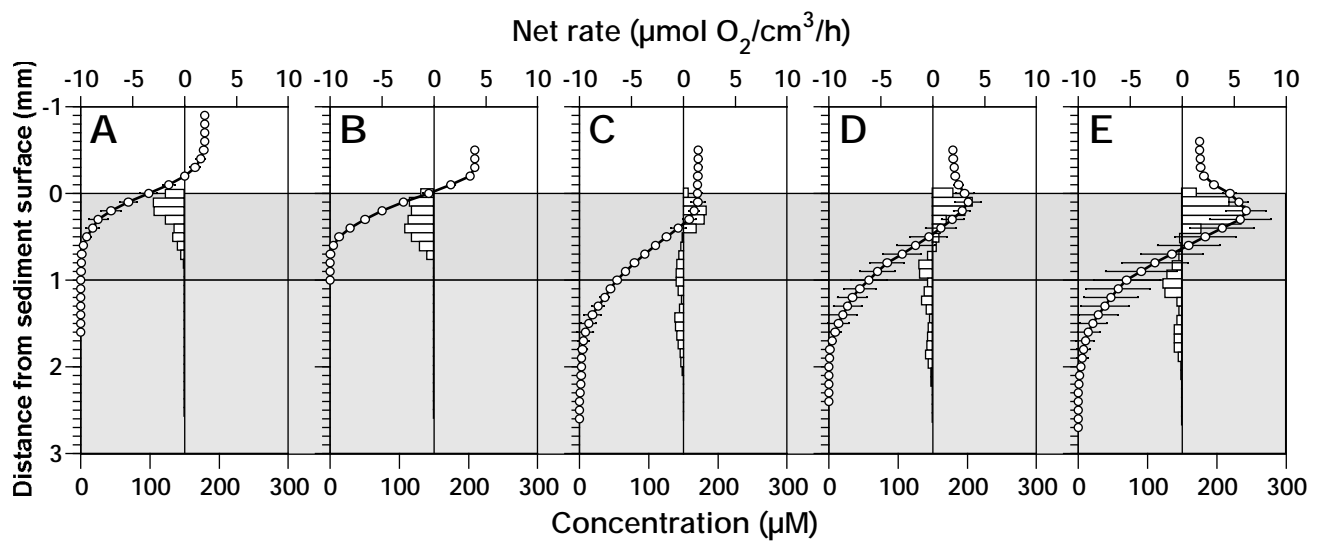
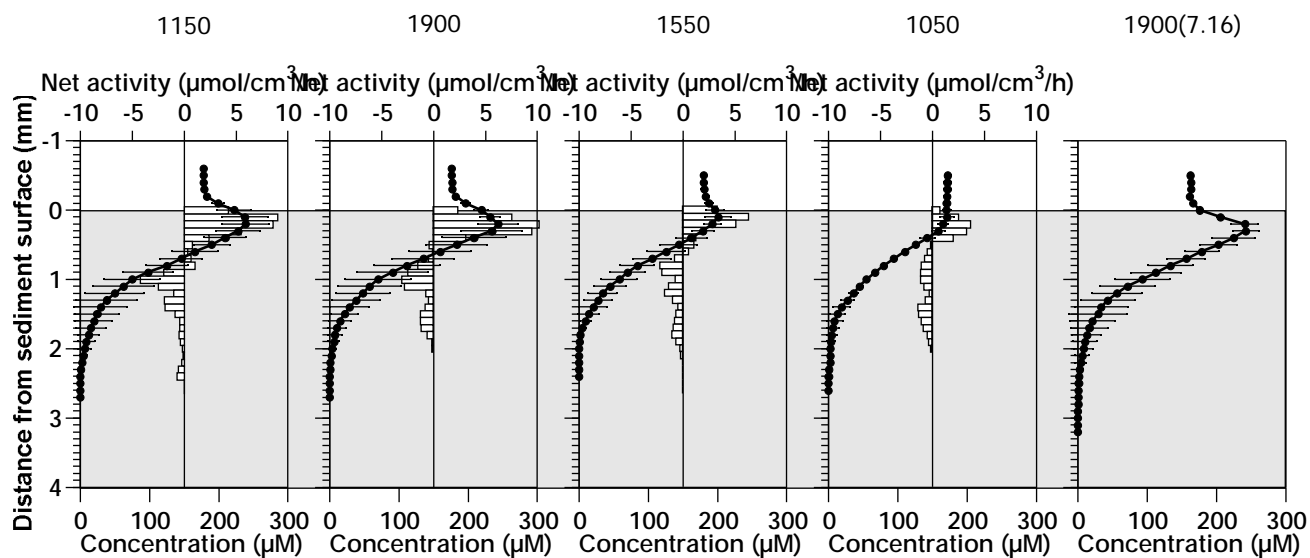
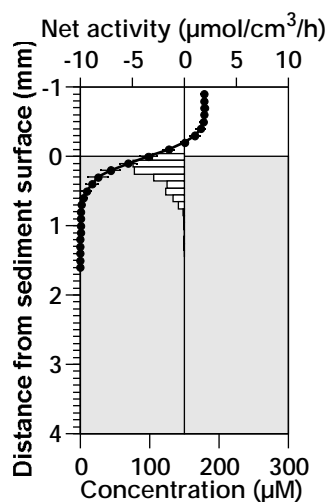


Fig. 7
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10(01.9.20)



900(8.1-1)

1900(8.1-2)

400(8.1-4)

10(8.1-5)

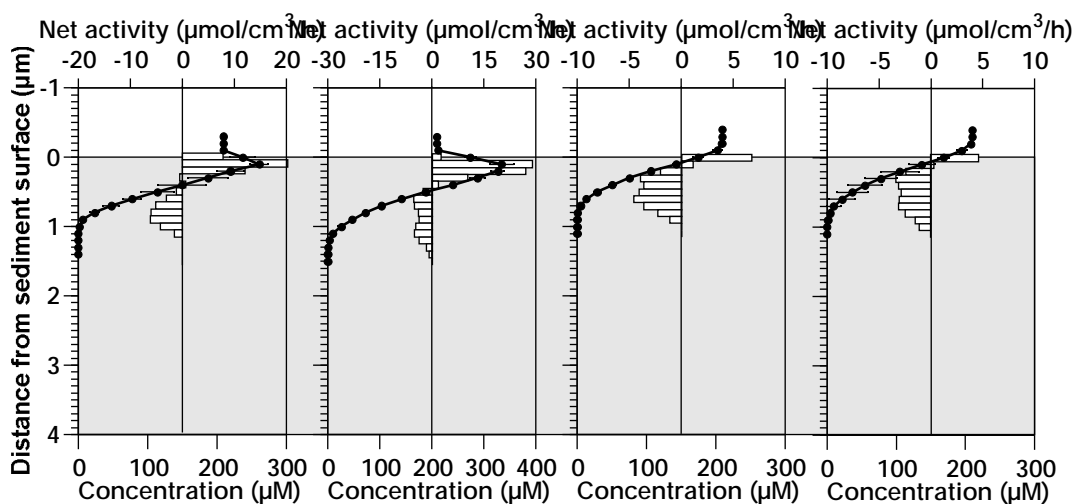


Fig.*

Macroscale.

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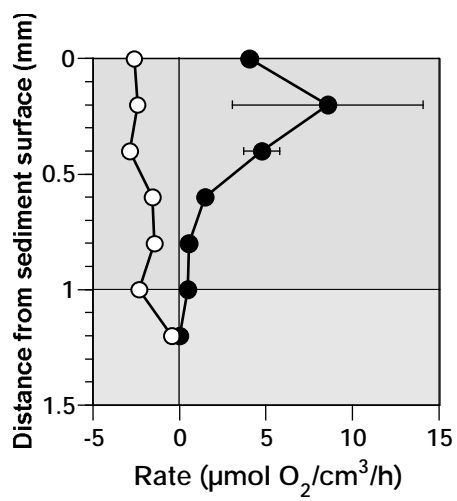


Fig. 8
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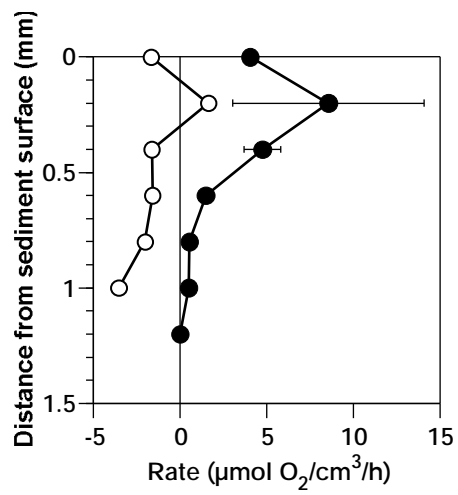
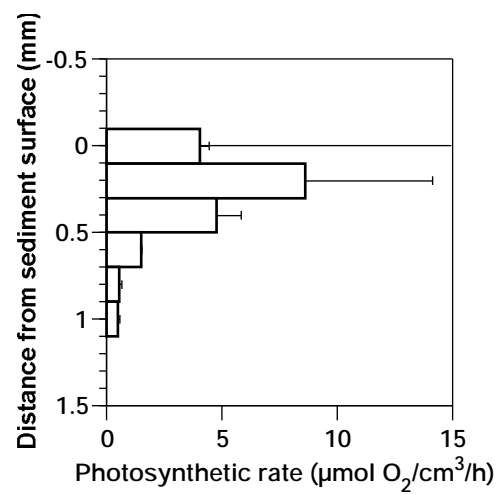


Fig.*
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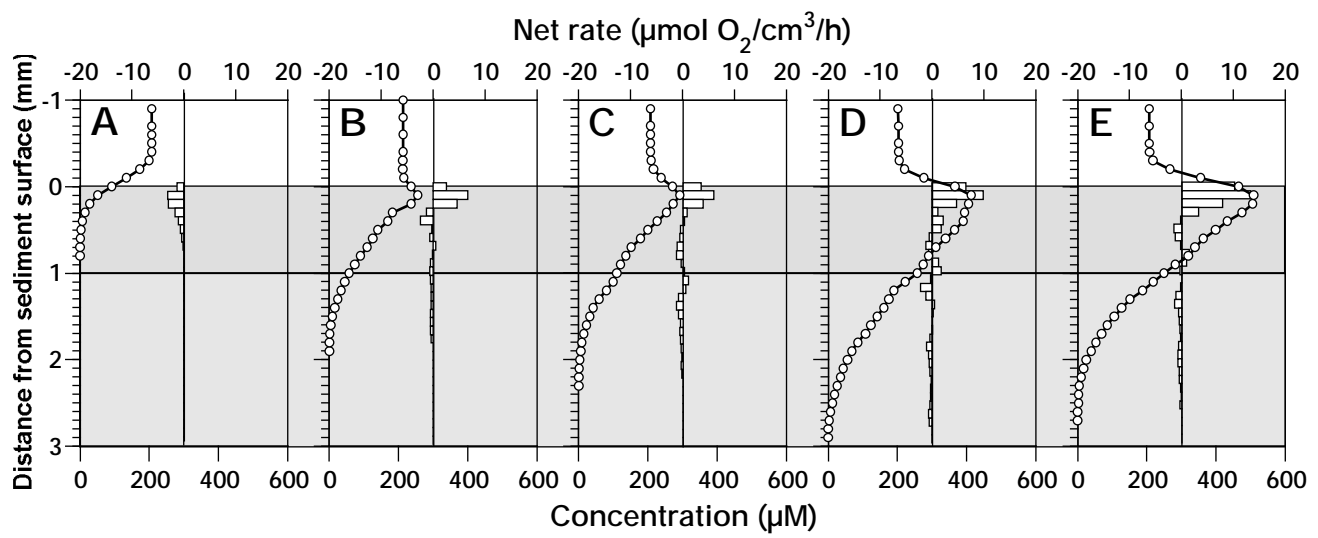
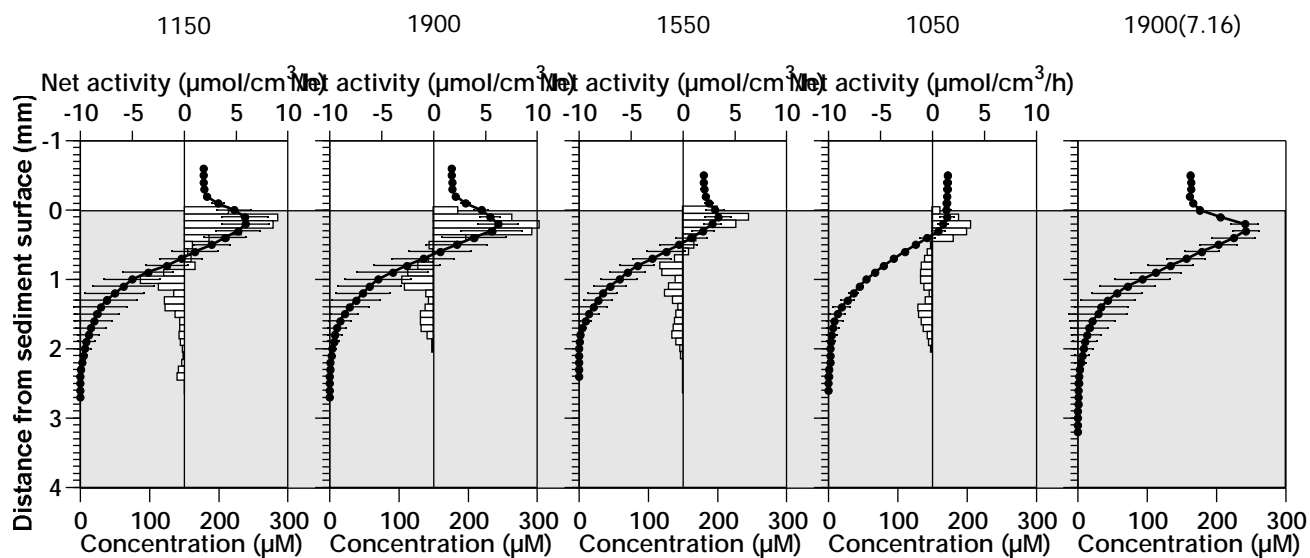
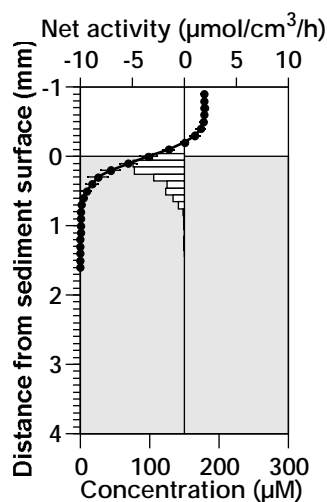


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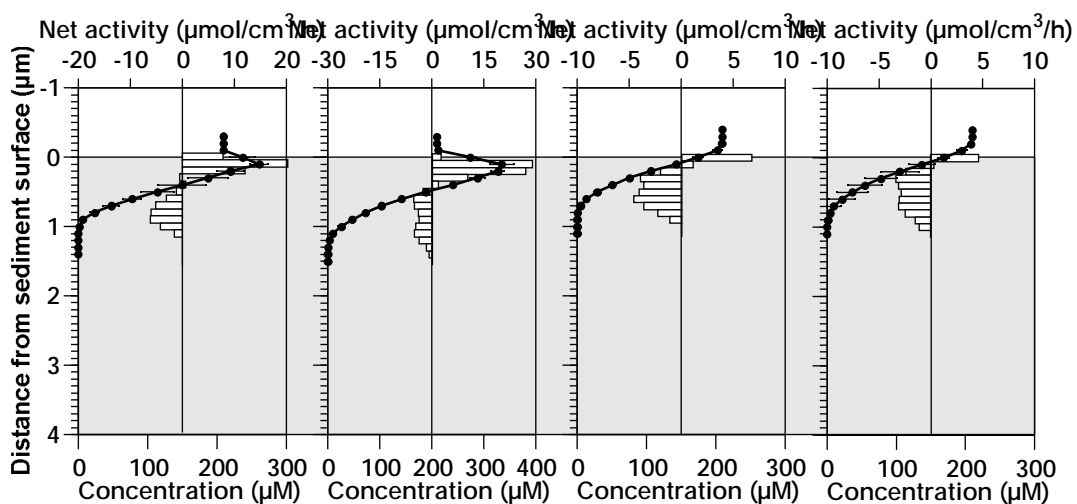


Fig.*

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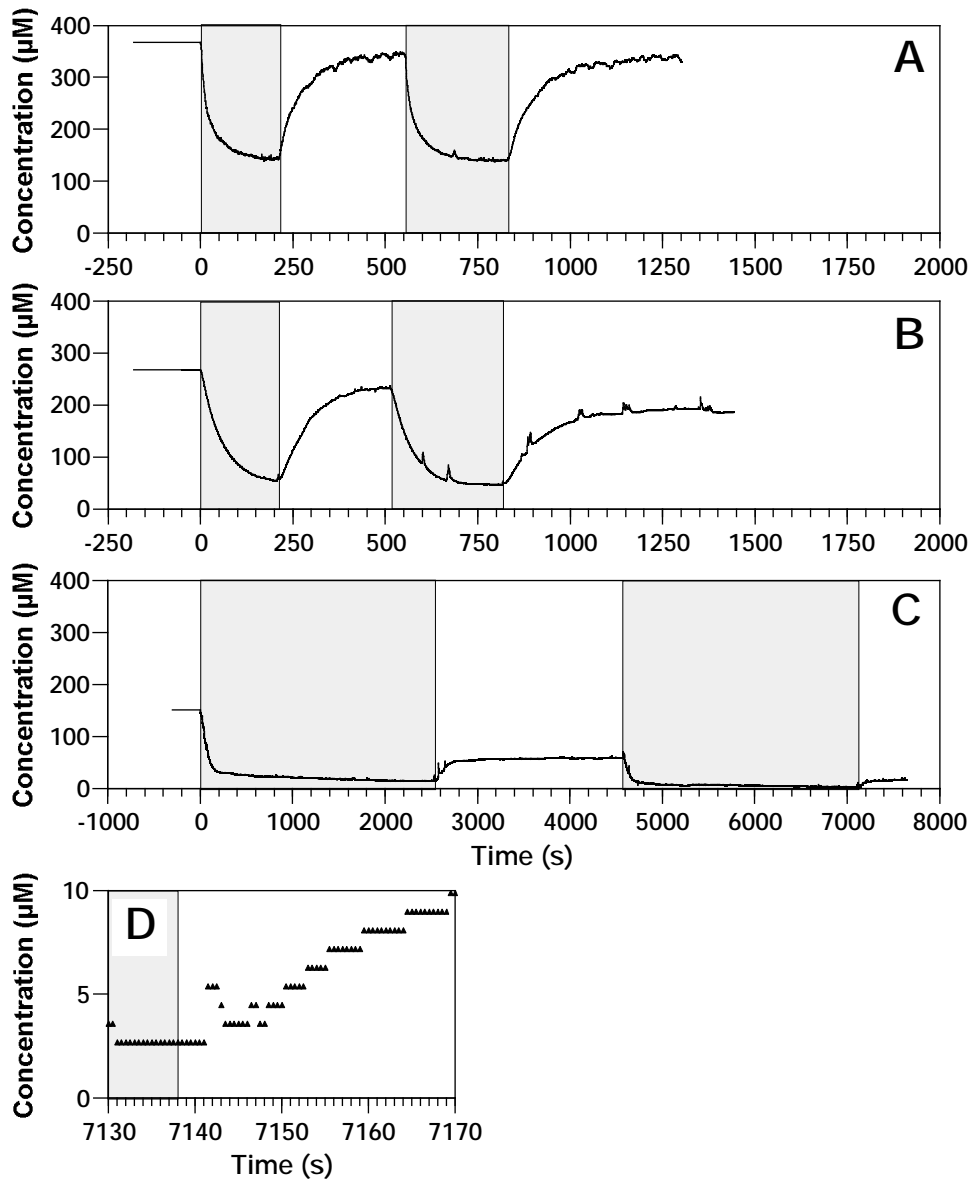


Fig. 10
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