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Two-dimensional similarity solutions for finite-mass granular avalanches with Coulomb- and viscous-type frictional resistance

KOLUMBAN HUTTER AND RALF GREVE

Institut für Mechanik, Technische Hochschule Darmstadt, D-W-6100 Darmstadt, Germany

ABSTRACT. This paper is concerned with the motion of an unconfined finite mass of granular material down an inclined plane when released from a rest position in the shape of a circular or elliptical paraboloid. The granular mass is treated as a frictional Coulomb-like continuum with a constant angle of internal friction. The basal friction force is assumed to be composed of a Coulomb-type component with a bed-friction angle that is position-dependent and a viscous Voellmy-type resistive stress that is proportional to the velocity squared. The model equations are those of Hutter and others (in press b) and form a spatially two-dimensional set for the evolution of the avalanche height and the depth averaged in-plane velocity components; they hold for a motion of a granular mass along a *plane* surface.

Similarity solutions, i.e. solutions which preserve the shape and the structure of the velocity field, are constructed by decomposing the motion into that of the centre of mass and the deformation relative to it. This decomposition is possible provided the effect of the Voellmy drag on the deformation is ignored. With it, the depth and velocities relative to those of the centre of mass of the moving pile can be determined analytically. It is shown that the pile has a parabolic cap shape and contour lines are elliptical. The semi-axes and the position and velocity of the centre of mass are calculated numerically. We explicitly show that

- (i) For two-dimensional spreading, a rigid-body motion does not exist, no matter what be the values of the bed-friction angle and the coefficient of viscous drag.
- (ii) A steady final velocity of the centre of the mass cannot be assumed, but the motion of the centre of mass depends strongly on the value of the Voellmy coefficient.
- (iii) The geometry of the moving pile depends on the variation of the bed-friction angle with position, as well as on the value of the coefficient of viscous drag.

1. INTRODUCTION

This paper — the second in a sequence of the study of the behaviour of the motion of a finite mass of granular mass subject to Coulomb and Voellmy-type resistive drag — is thought to be a further contribution to the understanding of the dynamics of flow avalanches, which have a negligible air-borne powder-snow contribution. It is intended to extend the work of Hutter and Savage and co-workers and, in particular, that of Hutter and Nohguchi (1990) and aims at an improved description of the classical Voellmy (1955), Salm (1968) and Perla and others (1980) models.

Recent years have witnessed an increased impetus in the study of landslides, rockfalls and snow and ice avalanches, viz. Alean (1984, 1985), Beghin and others (1981), Hermann and Hutter (1991), Greve and Hutter (1993), Gubler (1987, 1989, unpublished), Gubler and Hiller (1984), Hopfinger (1983), Hopfinger and Beghin (1980), Hsü (1975, 1978), Hutter (1992), Hutter and Koch (1991), Hutter and Savage (1988a, b), Hutter and others (in press), Lang and others (1989), Lang and

Martinelli (1979), Norem and others (1987, 1988), Perla and Martinelli (1978), Perla and others (1980), Salm (1966, 1968), Savage and Hutter (1989, 1991), Savage and Nohguchi (1988), Scheiwiller (1986), Scheiwiller and Hutter (1982), Scheiwiller and others (1987), Tochon-Danguy (1977), Tochon-Danguy and Hopfinger (1975), Vila (1987), Voellmy (1955), and others. All these works deal in one way or another with the mathematical formulation of the model equations of such catastrophic motions, their integration and, if possible, their verification against laboratory and field observations.

It is obvious that the direct observation of the dynamics of rockfalls or avalanches is extremely difficult and is probably only possible by remote-sensing techniques. Gubler (1987, unpublished) measured velocities and depths of flow avalanches that were artificially released in central Switzerland using radar-Doppler techniques, while Norem and others (1986, 1988) and Norem and Kristensen (1988) measured avalanche speeds and forces in the Ryggfonn Project of the Norwegian Geotechnical Institute when real avalanches were traced in three consecutive winters. The existing field data,

against which theoretical models can be tested, are still very limited but it is known that the classical theoretical formulations (Voellmy, 1955; Salm, 1966, 1968; Perla and Martinelli, 1978) were known to be oversimplified. Run-out distances and deposition areas cannot be sufficiently accurately predicted by these models. Reasons for these inadequacies must probably be sought in (i) the difficulties of parameter identification and in (ii) an insufficient resolution of the physics (rheological properties, sliding conditions), and of the geometry of the moving avalanche.

In this paper, our aim is not so much in a demonstration that our model will do better than the classical ones when compared with observations. For chute flows this has been demonstrated by Hutter and Koch (1991), Greve and Hutter (1993) and Hutter and others (in press). Here we take our new model, which is the first to our knowledge that describes the two-dimensional motion of the moving surface, and construct particular solutions in semi-analytical form. These solutions have diagnostic value insofar as they permit parameter studies and thereby inform us about the performance of the model; this provides physical insight.

We use the governing equations of Hutter and others (in press). These were derived from the three-dimensional balance laws of mass and momentum of an incompressible continuum obeying a Coulomb-friction law. The equations emerged from depth-averaging of three-dimensional equations and incorporate a bed-friction law that is composed of a Coulomb-type and a Voellmy-type viscous contribution. We assume the angle of internal friction is constant but allows the bed-friction angle to vary with position. It is at this point where our model provides the necessary flexibility sought by the avalanche practitioners.

Conceptually, it may be of advantage if the equations governing the motion of the moving mass are decomposed into those governing the motion of the centre of mass and those describing the deformation. For the equations of Hutter and others (in press b), this decomposition is not possible in general, but when the effect of the Voellmy drag on the deformation of the moving pile is ignored, then this decomposition describes the complete motion very accurately. The decomposition is the basis for the construction of *similarity solutions*. Such solutions have previously been constructed for the one-dimensional chute-flow situation (Savage and Nohguchi, 1988; Nohguchi and others, 1989; Savage and Hutter, 1989; Hutter and Nohguchi, 1990); they have the property of *permanent shape*. In the spatially two-dimensional situation of unconfined flow, a similarity solution means that the (mathematical) structure of the shape is preserved but its aspect ratio may still change.

We show that a finite mass of gravel that starts from rest in the form of a circular or elliptical parabola will maintain its elliptical-paraboloidal shape. We derive the governing equations for the position and velocity of its centre of mass and the evolution of the semi-axis of the ellipse. These are non-linear ordinary differential equations which must be solved numerically.

In general, the motion of the centre of mass and the evolution of the deformation of the pile are coupled and separate only in special cases. We construct numerical

solutions with the Runge-Kutta technique and analyze the performance of the model when the inclination angle of the plane, the initial depth to length ratio, the bed-friction angle and the Voellmy drag coefficient are varied. Solutions are discussed in detail.

2. GOVERNING EQUATIONS

Consider free surface flow of a granular material down a slowly varying topography. Identify the mean plane surface of this topography with a plane that is parallel to the xy -plane of a three-dimensional Cartesian coordinate system. Let the x -coordinate follow the direction of steepest descent, the y -coordinate the horizontal lines and let the z -coordinate be perpendicular to these (see Fig. 1). Thus, the z -axis will be inclined with respect to the vertical by the angle ζ . The bottom and free surface of the moving mass will be defined by $z = b(x, y)$ and $z = f(x, y, t)$, respectively; the margin curve is therefore given by the condition $f - b = 0$. The space within $0 \leq z - b \leq f - b = h$ is filled with a granular material, which is assumed to be treatable as a fluid-like continuum. This supposes that the thickness, $h = f - b$, of the sliding and deforming body extends over several particle diameters. We accept the fact that this condition cannot be satisfied close to the margin.

The granular continuum is regarded as incompressible, an assumption one may debate about, but this condition is sufficiently satisfied in the entire space filled by the granular material except in a very thin fluidized layer close to the bottom. This layer will be absorbed into a sliding condition.

Balance laws of mass and momentum are

$$\nabla \cdot \mathbf{u} = 0, \quad \frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} = -\frac{1}{\rho} \nabla \cdot \mathbf{p} + \mathbf{g}, \quad (2.1)$$

in which $\mathbf{u}, \rho, \mathbf{p}, \mathbf{g}$ are the velocity vector, constant density, pressure tensor and gravity vector. In the coordinates of Figure 1 we have

$$\mathbf{g} = g(\sin \zeta, 0, -\cos \zeta). \quad (2.2)$$

Boundary conditions have to be formulated at the free surface and the base, and comprise kinematic and

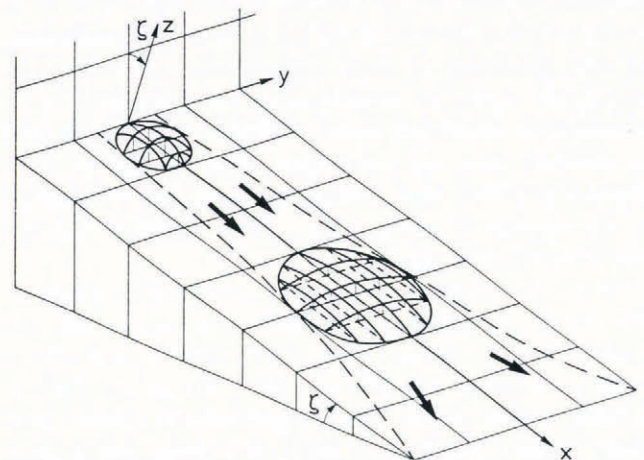


Fig. 1. Definition of configuration and coordinate system (physical coordinates).

dynamic statements. At the free surface, $F_S = f(x, y, t) - z \equiv 0$, we have

$$\left. \begin{aligned} \frac{\partial F_S}{\partial t} + \nabla F_S \cdot \mathbf{u} &= 0, \\ \mathbf{p}\mathbf{n} &= 0, \end{aligned} \right\} \text{ at } F_S(\mathbf{x}, t) = 0, \quad (2.3)$$

where \mathbf{n} is the unit-normal vector. At the base the tangency condition for the velocity must be fulfilled. With $F_B = b - z \equiv 0$ this reads

$$\nabla F_B \cdot \mathbf{u} = 0, \quad (\mathbf{u} \cdot \mathbf{n} = 0). \quad (2.4)$$

We regard the material as a cohesionless granular body obeying a Mohr–Coulomb yield criterion with a constant internal friction angle. This states that yielding will occur on a plane element when

$$\|\mathbf{S}\| = \|\mathbf{N}\| \tan \phi \quad (2.5)$$

where \mathbf{S} and \mathbf{N} are, respectively, the shear and normal stress acting on the element. By contrast, at the basal surface, the friction law is assumed to consist of two components

$$\mathbf{S} = \mathbf{S}_C + \mathbf{S}_V. \quad (2.6)$$

The first is a Coulomb-type dry-friction law, where \mathbf{S} and \mathbf{N} are now related by

$$\mathbf{S}_C = -\|\mathbf{N}\| \tan \delta \frac{\mathbf{u}}{\|\mathbf{u}\|}, \quad (2.7)$$

in which δ is the bed-friction angle (which is generally smaller than ϕ). The second is a *viscous* drag, very much like the classical Voellmy drag in the early avalanche models and has the form

$$\mathbf{S}_V = \rho q (\|\mathbf{u}\|, \|\mathbf{N}\|) \|\mathbf{u}\| \mathbf{u}, \quad (2.8)$$

in which q is the dimensionless drag coefficient that may depend on the modulus of the velocity vector and the stress-normal to the basal surface. For a q independent of $\|\mathbf{u}\|$, Equation (2.8) corresponds to a quadratic dependence of \mathbf{S}_V on the velocity, but, clearly, any other dependence is also possible. Furthermore, avalanche dynamicists who are used to working with the Voellmy model usually define \mathbf{S}_V in the form

$$\mathbf{S}_V = \frac{\rho g}{\xi} \|\mathbf{u}\| \mathbf{u} \quad (2.9)$$

where ξ is a “viscosity” having the dimension of acceleration and

$$q = \frac{g}{\xi}. \quad (2.10)$$

Hutter and others (in press b) have motivated a reduction of the initial boundary-value problem outlined above. It is not our goal here to repeat this derivation; in short, the intention is not to determine the three-dimensional distribution of the velocity field, but only depth averages of this vector field. As a result, the spatially three-dimensional problem is reduced by one spatial dimension and therefore becomes much simpler.

In the derivation of the simplified equations Hutter and others (in press b) introduced the Cartesian coordinates of Figure 1 and wrote the previous equations in dimensionless form by introducing the following scales:

$$(x, y, z) = ([L_x]x^*, [L_y]y^*, [H]z^*),$$

$$(f(x, y, t), b(x, y)) = [H](f^*(x^*, y^*, t^*), b^*(x^*, y^*)),$$

$$(u, v, w) = ([U]u^*, [V]v^*, [W]w^*),$$

$$(2.11)$$

$$t = [T]t^*,$$

$$(p_{xx}, p_{yy}, p_{zz}, p_{xz}) = \rho g [H](p_{xx}^*, p_{yy}^*, p_{zz}^*, p_{xz}^*),$$

$$(p_{yz}, p_{xy}) = \rho g [H](\lambda_1 p_{yz}^*, \lambda_2 p_{xy}^*).$$

Quantities in brackets are typical scales for the physical quantities under consideration, the only exceptions being the shear stresses, for which λ_1 and λ_2 account for their smallness. Variables carrying an asterisk are dimensionless. Hutter and others (in press b) further supposed that the following order relations hold:

$$\frac{[L_x]}{[U]} = \frac{[L_y]}{[V]} = \frac{[H]}{[W]} = [T], \quad [U]^2 = g[L_x],$$

$$\epsilon_x = \frac{[H]}{[L_x]}, \quad \epsilon_y = \frac{[H]}{[L_y]}, \quad \epsilon_{xy} = \frac{[L_y]}{[L_x]}, \quad (2.12)$$

$$\lambda_1 \ll 1, \quad \lambda_2 \ll 1$$

$$\epsilon_x \ll 1, \quad \epsilon_y \ll 1, \quad \epsilon_{xy} \leq 1.$$

In other words, the ratio of the velocity scales equals the corresponding length scale ratio and, the time-scale equals any of the length scales divided by the corresponding velocity scale. Moreover, we have set the characteristic longitudinal velocity square proportional to the product of the gravity constant with the longitudinal characteristic length. The ratios $\epsilon_x, \epsilon_y, \epsilon_{xy}$ are simply definitions but the last three of Equations (2.12) suppose that the moving masses are *long*, *wide* and also *shallow*. Different ordering symbols are used to indicate that the moving piles may possess a large range of aspect ratios, but always such that $\epsilon_{xy} < 1$. (A convenient choice would be $\epsilon_y = \epsilon_x^\alpha$, $0.5 \leq \alpha \leq 1$, so that $\epsilon_{xy} = \epsilon_x^{1-\alpha}$, $0 \leq (1-\alpha) \leq 0.5$; if $\alpha = 1$ then $\epsilon_x = \epsilon_y$ and $\epsilon_{xy} = 1$, a very common case we shall adopt here.) Furthermore, in presenting the reduced equations below we shall suppose that ϵ_x, ϵ_y are sufficiently small so that terms of order ϵ_x^2 and ϵ_y^2 can be ignored.

The dimensionless equations that emerge from this procedure involve the longitudinal and transverse (side-wise) pressure p_{xx}^*, p_{yy}^* which through the Mohr–Coulomb yield criterion are expressed in terms of the overburden pressure p_{zz}^* . We shall distinguish between *active* and *passive* stress states by writing

$$p_{xx}^* = k_{\text{actpass}}^x p_{zz}^*, \quad (2.13)$$

$$p_{yy}^* = k_{\text{actpass}}^y p_{zz}^*.$$

In view of the dominant downward motion, we expect one principal axis of strain rate to be very nearly the xz -plane. The stress p_{xx}^* can then easily be related to p_{zz}^* with the aid of the Mohr–Coulomb yield circle (see Fig. 2), so that

$$k_{act}^x = \frac{2}{\cos^2 \phi} \left(1 - \sqrt{1 - (1 + \tan^2 \delta) \cos^2 \phi} \right) - 1, \quad \text{for } \frac{\partial u}{\partial x} > 0, \quad (2.14)$$

$$k_{pass}^x = \frac{2}{\cos^2 \phi} \left(1 + \sqrt{1 - (1 + \tan^2 \delta) \cos^2 \phi} \right) - 1, \quad \text{for } \frac{\partial u}{\partial x} < 0.$$

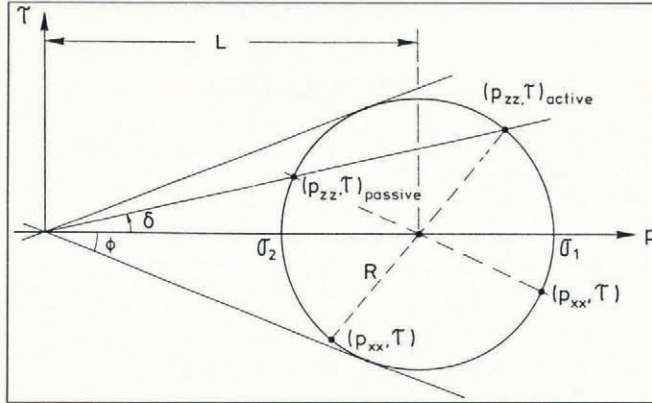


Fig. 2. Mohr's circle representation of the active and passive stress states.

It is common to assume that, if the material is at the point of yielding, the stress in the other direction (p_{yy}^*) corresponds either to the major or minor principal stress σ_1 or σ_2 . Again, with reference to Figure 2, we may thus deduce

$$k_{act}^y = \frac{1}{2} \left(k_{act}^x + 1 - \sqrt{(k_{act}^x - 1)^2 + 4 \tan^2 \delta} \right), \quad \text{for } \frac{\partial v}{\partial y} > 0, \quad (2.15)$$

$$k_{pass}^y = \frac{1}{2} \left(k_{pass}^x + 1 + \sqrt{(k_{pass}^x - 1)^2 + 4 \tan^2 \delta} \right), \quad \text{for } \frac{\partial v}{\partial y} < 0.$$

Relations (2.13)–(2.15) are reasonable at the bed and again at the free surface, where $p_{zz} = 0$ implies $p_{xx} = p_{yy} = 0$. So, by way of continuity, the postulates (2.13)–(2.15) are logical assumptions. They are common in the soil literature. For motions along the plane-inclined surface, we have $b = 0$ and the depth-averaged field equations take the form

$$\frac{\partial h}{\partial t} + \frac{\partial(h\bar{u})}{\partial x} + \frac{\partial(h\bar{v})}{\partial y} = 0,$$

$$\frac{\partial(h\bar{u})}{\partial t} + \frac{\partial(h\bar{u}^2)}{\partial x} + \frac{\partial(h\bar{u}\bar{v})}{\partial y} = Ah - B \frac{\partial(h^2/2)}{\partial x} - C \text{sgn}(\bar{u})h - \frac{\bar{u}}{\Xi} \|\bar{\mathbf{u}}\|,$$

$$\frac{\partial(h\bar{v})}{\partial t} + \frac{\partial(h\bar{u}\bar{v})}{\partial x} + \frac{\partial(h\bar{v}^2)}{\partial y} = -D \frac{\partial(h^2/2)}{\partial y} - Ch \frac{\bar{v}}{\|\bar{\mathbf{u}}\|} - \frac{\bar{v}}{\Xi} \|\bar{\mathbf{u}}\|, \quad (2.16)$$

with

$$A = \sin \zeta$$

$$B = \epsilon_x \cos \zeta k_{actpass}^x,$$

$$C = \cos \zeta \tan \delta,$$

$$D = \frac{\epsilon_y}{\epsilon_{xy}} \cos \zeta k_{actpass}^y, \quad (2.17)$$

$$\|\bar{\mathbf{u}}\| = \sqrt{\bar{u}^2 + \epsilon_{xy}^2 \bar{v}^2} \cong \bar{u}$$

$$\Xi = \frac{\epsilon}{q} = \frac{\epsilon \zeta}{g}$$

(see Hutter and others, in press b, equations (3.1) and (3.2)). Here we have kept $\epsilon_{xy}^2 \bar{v}^2$ in the expression for $\|\bar{\mathbf{u}}\|$, because this term is not negligible at the initial stage of the motion from a state at rest. At later times — in reality immediately after the motion has set in — this higher-order term may be dropped. We shall omit this term at all times and accept the inconsistency in the procedure at early times. In Equations (2.16) and (2.17) h is the depth, \bar{u} and \bar{v} are the depth-averaged dimensionless velocity components

$$h\bar{u} = \int_0^h u(x, y, z, t) dz,$$

$$h\bar{v} = \int_0^h v(x, y, z, t) dz,$$

and all quantities are dimensionless (asterisks have been dropped for simplicity). Equation (2.16)₁ is the mass balance in depth-integrated form, and Equations (2.16)_{2,3} represent the in-plane longitudinal and transverse momentum balances. Equations (2.16) are solved subject to the boundary conditions

$$h_M = h(x_M(t), y_M(t), t) = 0 \quad (2.18)$$

and prescribed initial conditions, usually

$$h(x, y, 0) = h_I(x, y), \quad (2.19)$$

$$\bar{u}(x, y, 0) = \bar{v}(x, y, 0) = 0.$$

Equation (2.18) describes the margin and Equation (2.19)₁ the initial mass distribution. Integration of Equations (2.16)–(2.19) is likely only numerically possible.

3. EQUATIONS GOVERNING THE CENTRE OF MASS MOTION AND THE DEFORMATION

3.1. Symbolic form of the field equations

Equations (2.16) are now separated into two sets of equations, one governing the motion of the centre of mass, the other describing the deformation or the deviation from the rigid-body motion. To this end, we write them as

$$\begin{aligned} \frac{\partial h}{\partial t} + \operatorname{div}(h\bar{\mathbf{v}}) &= 0, \\ \frac{\partial(h\bar{\mathbf{v}})}{\partial t} + \operatorname{div} \mathbf{t} &= A h \mathbf{e} - \mathbf{B} \operatorname{grad} \left(\frac{h^2}{2} \right) \\ &\quad - C \frac{h\bar{\mathbf{v}}}{\|\bar{\mathbf{v}}\|} - \frac{1}{\Xi} \|\bar{\mathbf{v}}\| \bar{\mathbf{v}}. \end{aligned} \quad (3.1)$$

Here, the vector and tensor quantities and the corresponding operators are two-dimensional. In particular,

$$\mathbf{t} = \begin{pmatrix} h\bar{u}^2 & h\bar{u}\bar{v} \\ h\bar{u}\bar{v} & h\bar{v}^2 \end{pmatrix}, \quad \mathbf{B} = \begin{pmatrix} B & 0 \\ 0 & D \end{pmatrix}, \quad (3.2)$$

and $\bar{\mathbf{v}} = (\bar{u}, \bar{v})$, while \mathbf{e} is the unit vector

$$\mathbf{e} = (1, 0)^T, \quad (3.3)$$

in which T denotes transposition.

3.2. Equations describing the centre of mass motion

Let $\mathcal{A}(t)$ be the domain in the (x, y) space that is, at time t , covered by the avalanching mass; let, further, $\partial\mathcal{A}(t)$ be its boundary (margin). Then integrate Equation (3.1)₁ over \mathcal{A} , so that

$$\iint_{\mathcal{A}} \frac{\partial h}{\partial t} da + \iint_{\mathcal{A}} \operatorname{div}(h\bar{\mathbf{v}}) da = 0. \quad (3.4)$$

By using Gauss's theorem in the second term and Reynolds' transport theorem in the first,

$$\iint_{\mathcal{A}} \operatorname{div}(h\bar{\mathbf{v}}) da = \oint_{\partial\mathcal{A}} h\bar{\mathbf{v}} \cdot \mathbf{n} ds, \quad (3.5)$$

$$\iint_{\mathcal{A}} \frac{\partial h}{\partial t} da = \frac{d}{dt} \iint_{\mathcal{A}} h da - \oint_{\partial\mathcal{A}} h\bar{\mathbf{v}} \cdot \mathbf{n} ds,$$

where \mathbf{n} is the exterior unit vector normal to $\partial\mathcal{A}$, Equation (3.4) becomes

$$\frac{d}{dt} \iint_{\mathcal{A}} h da = 0. \quad (3.6)$$

In other words, the total volume of the avalanche is conserved. Because of incompressibility, this is tantamount to the conservation of mass. In much the same way, the momentum Equation (3.1)₂ may be transformed. The computations are straight-forward and the results are

$$\begin{aligned} \frac{d}{dt} \iint_{\mathcal{A}} h\bar{\mathbf{v}} da &= \mathbf{e} \iint_{\mathcal{A}} A h da + \iint_{\mathcal{A}} (\operatorname{div} \mathbf{B}) \frac{h^2}{2} da \\ &\quad - \iint_{\mathcal{A}} C \frac{h\bar{\mathbf{v}}}{\|\bar{\mathbf{v}}\|} da - \iint_{\mathcal{A}} C \frac{\bar{\mathbf{v}}\|\bar{\mathbf{v}}\|}{\Xi} da. \end{aligned} \quad (3.7)$$

In the derivation of Equation (3.7), the divergence theorem has been employed at various places, and the boundary condition that h vanishes along $\partial\mathcal{A}$ has been invoked. Furthermore, Reynolds' transport theorem has been employed. Equation (3.7) is the expression of global momentum balance.

We now define *averaged field quantities* as follows

$$\begin{aligned} \bar{h} &= \frac{1}{\mathcal{A}} \iint_{\mathcal{A}} h da, \\ \bar{\mathbf{v}}_c &= \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} h\bar{\mathbf{v}} da, \\ \bar{A} &= \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} A h da, \end{aligned} \quad (3.8)$$

$$\begin{aligned} \frac{1}{2} \bar{\mathbf{B}} \bar{h} &= \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} \operatorname{div} \mathbf{B} \frac{h^2}{2} da, \\ \overline{C \frac{\bar{\mathbf{v}}}{\|\bar{\mathbf{v}}\|}} &= \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} C \frac{h\bar{\mathbf{v}}}{\|\bar{\mathbf{v}}\|} da, \\ \frac{\overline{\bar{\mathbf{v}}\|\bar{\mathbf{v}}\|}}{\Xi h} &= \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} h \frac{\bar{\mathbf{v}}\|\bar{\mathbf{v}}\|}{\Xi h} da. \end{aligned}$$

\bar{h} is the depth averaged over the pile volume, $\bar{\mathbf{v}}_c$ the velocity of the centre of mass. The averages of the other quantities are formal means which will further be reduced below. Note, however, that owing to definition (3.2)₂, $\bar{\mathbf{B}}$ is a vector with the components

$$\bar{\mathbf{B}} = \frac{2}{h^2 \mathcal{A}} \left\{ \iint_{\mathcal{A}} \frac{\partial B}{\partial \bar{x}} \frac{h^2}{2} da, \iint_{\mathcal{A}} \frac{\partial D}{\partial \bar{y}} \frac{h^2}{2} da \right\}^T \quad (3.9)$$

where B and D are defined in Equation (2.17). With the definitions (3.8), Equation (3.7) takes the form

$$\frac{d\bar{\mathbf{v}}_c}{dt} = \bar{A} \mathbf{e} + \frac{1}{2} \bar{\mathbf{B}} \bar{h} - \overline{C \frac{\bar{\mathbf{v}}}{\|\bar{\mathbf{v}}\|}} - \frac{\overline{\bar{\mathbf{v}}\|\bar{\mathbf{v}}\|}}{\Xi h}. \quad (3.10)$$

We next assume geometric and dynamic symmetry of the avalanche with respect to the x -axis of Figure 1 that is in the direction of steepest descent,

$$\begin{aligned} h(x, y, t) &= h(x, -y, t), \\ \bar{u}(x, y, t) &= \bar{u}(x, -y, t), \\ \bar{v}(x, y, t) &= -\bar{v}(x, -y, t). \end{aligned} \quad (3.11)$$

Then Equation (3.8)₂ implies $\bar{v}_c \equiv 0$ for all times, and the second component of Equation (3.10) is satisfied if, for instance

$$\begin{aligned} C(x, y) &= C(x, -y), \\ \Xi(x, y) &= \Xi(x, -y), \\ \bar{B}_y &= 0. \end{aligned} \quad (3.12a)$$

In view of Equation (3.9), the last of these requires D to be symmetric,

$$D(x, y) = D(x, -y). \quad (3.12b)$$

Going back to the original definitions (2.17) of the coefficients C , D and Ξ , it is seen that the requirements (3.12) presuppose the friction angles ϕ , δ and the Voellmy coefficient Ξ to be symmetrically distributed. This is what we shall assume. In the above Equation (3.10), $\bar{\mathbf{v}}_c$ is the centre of mass velocity whose position, $\bar{\mathbf{x}}_c$, is given by

$$\bar{\mathbf{x}}_c = \frac{1}{h\mathcal{A}} \iint_{\mathcal{A}} h \bar{\mathbf{x}} da. \quad (3.13)$$

It can be determined by integrating the equation of motion

$$\frac{d\bar{\mathbf{x}}_c}{dt} = \bar{\mathbf{v}}_c. \quad (3.14)$$

It is not difficult to show that Equation (3.14) is in conformity with Equation (3.8)₂. Indeed,

$$\begin{aligned} \frac{d\bar{x}_c}{dt} &= \frac{1}{\bar{h}\mathcal{A}} \frac{d}{dt} \iint_{\mathcal{A}} h\bar{x} da \\ &= \frac{1}{\bar{h}\mathcal{A}} \iint_{\mathcal{A}} \left(h\bar{v} + \frac{\partial h}{\partial t} \bar{x} + (\text{grad } h \cdot \bar{v})\bar{x} \right. \\ &\quad \left. + h\bar{x} \text{div } \bar{v} \right) da \\ &= \frac{1}{\bar{h}\mathcal{A}} \iint_{\mathcal{A}} \left(h\bar{v} + \left[\frac{\partial h}{\partial t} + \text{div}(h\bar{v}) \right] \bar{x} \right) da \\ &= \frac{1}{\bar{h}\mathcal{A}} \iint_{\mathcal{A}} h\bar{v} da = \bar{v}_c \quad \text{q.e.d. ;} \end{aligned}$$

in the second to last step use has been made of Equation (3.4). Clearly, because of the symmetry assumptions Equations (3.11) and (3.12), only one component of Equation (3.14) is non-trivial, namely

$$\frac{d\bar{x}_c}{dt} = \bar{u}_c \tag{3.15}$$

while $d\bar{y}_c/dt = 0$, since $\bar{v}_c \equiv 0$.

3.3. Deformation equations

Having defined the centre of mass motion through Equations (3.10) and (3.14), we now proceed to the derivation of the deformation equation. To this end, it is convenient to define the new independent variables

$$\xi = \bar{x} - \bar{x}_c, \eta = \bar{y}, \tau = t \tag{3.16}$$

from which one obtains

$$\frac{\partial}{\partial \bar{x}} = \frac{\partial}{\partial \xi}, \frac{\partial}{\partial \bar{y}} = \frac{\partial}{\partial \eta}, \frac{\partial}{\partial t} = \frac{\partial}{\partial \tau} - \bar{u}_c \frac{\partial}{\partial \xi} \tag{3.17}$$

Introducing the difference velocity

$$\tilde{u} = \bar{u} - \bar{u}_c, \tilde{v} = \bar{v}_c, \tag{3.18}$$

or

$$\tilde{v} = \bar{v} - \bar{v}_c,$$

and the transformation rules (3.16) and (3.17) into the governing Equations (3.1), we may, on account of Equations (3.10), derive the following *deformation equations*

$$\begin{aligned} \frac{\partial h}{\partial \tau} + \text{div}_{\xi,\eta}(h\tilde{v}) &= 0, \tag{3.19} \\ \frac{\partial \tilde{v}}{\partial \tau} + (\text{grad}_{\xi,\eta} \tilde{v})\tilde{v} &= \mathbf{e}(A - \bar{A}) - \mathbf{B}(\text{grad}_{\xi,\eta} h) - \frac{1}{2} \bar{\mathbf{B}}\bar{h} \\ &\quad - \left\{ C \frac{\tilde{v}}{\|\tilde{v}\|} - \overline{C \frac{\tilde{v}}{\|\tilde{v}\|}} \right\} - \left\{ \frac{\tilde{v}\|\tilde{v}\|}{\Xi h} - \overline{\frac{\tilde{v}\|\tilde{v}\|}{\Xi h}} \right\}. \end{aligned}$$

The derivation of these equations is somewhat long but not difficult. We have also indicated that the differentiations in the operators are with respect to the coordinates ξ and η . On an inclined plane with constant internal and bed-friction angles A, B, C and D are constant, implying $A - \bar{A} = 0, \bar{\mathbf{B}} = 0$. Furthermore, if it is assumed that the difference velocities \tilde{u}, \tilde{v} are small in comparison to \bar{u}_c , so that quadratic and higher-order terms can be ignored, then

$$C \frac{\tilde{v}}{\|\tilde{v}\|} \cong C \left\{ \frac{\bar{v}_c}{|\bar{u}_c|} + \frac{\tilde{v}_c}{|\bar{u}_c|} - \frac{\bar{v}_c \tilde{u}}{|\bar{u}_c|^2} \right\},$$

$$\overline{C \frac{\tilde{v}}{\|\tilde{v}\|}} \cong C \frac{\bar{v}_c}{|\bar{u}_c|}, \tag{3.20}$$

$$\frac{\tilde{v}\|\tilde{v}\|}{\Xi h} \cong \frac{|\bar{u}_c|(\bar{v}_c + \tilde{v} + (\tilde{u}/\bar{u}_c)\bar{v}_c)}{\Xi \bar{h}},$$

$$\overline{\frac{\tilde{v}\|\tilde{v}\|}{\Xi h}} \cong \frac{|\bar{u}_c| \bar{v}_c}{\Xi \bar{h}}.$$

In expressions (3.20)_{3,4} we have also replaced h by \bar{h} , the mean depth, to avoid formation of singularities at the margin. In compensation, Ξ has been replaced by $\bar{\Xi}$, a suitable average over the avalanche. While the above relations correspond to additional ad hoc assumptions, they are needed if the separation into field variables describing the centre of mass motion and those describing the deformation is meaningful at all. On the other hand, the expressions (3.20) are also physically reasonable because, with the exception of the onset of the motion, the dispersive velocity is small in comparison to the centre of mass velocity. Moreover, singularities at the margins of the Voellmy resistive force are physically very unlikely and variations of the drag coefficient Ξ within the avalanche are difficult to determine.

With expressions (3.20), Equations (3.10), (3.15) and (3.19) reduce to the form

$$\begin{aligned} \frac{d\bar{x}_c}{dt} &= \bar{u}_c, \\ \frac{d\bar{u}_c}{dt} &= \sin \zeta - C \text{sgn}(\bar{u}_c) - \frac{1}{\bar{\Xi} \bar{h}} \bar{u}_c |\bar{u}_c|, \\ \frac{\partial h}{\partial \tau} + \frac{\partial(h\tilde{u})}{\partial \xi} + \frac{\partial(h\tilde{v})}{\partial \eta} &= 0, \tag{3.21} \\ \frac{\partial \tilde{u}}{\partial \tau} + \frac{\partial \tilde{u}}{\partial \xi} \tilde{u} + \frac{\partial \tilde{u}}{\partial \eta} \tilde{v} &= -B \frac{\partial h}{\partial \xi} - 2 \frac{|\bar{u}_c| \tilde{u}}{\bar{\Xi} \bar{h}}, \\ \frac{\partial \tilde{v}}{\partial \tau} + \frac{\partial \tilde{v}}{\partial \xi} \tilde{u} + \frac{\partial \tilde{v}}{\partial \eta} \tilde{v} &= -C \frac{\tilde{v}}{|\bar{u}_c|} - D \frac{\partial h}{\partial \eta} - \frac{|\bar{u}_c|}{\bar{\Xi} \bar{h}} \tilde{v}. \end{aligned}$$

The first two equations correspond to the classical Voellmy model in which the resistive force on a moving mass of snow is composed of a dry Coulomb drag and a viscous drag that is proportional to the squared velocity. These two equations are equivalent to Equations (3.10) and (3.14) and thus give the Voellmy model (and its equivalents due to Salm, McClung and others) a clear interpretation. However, Equations (3.21) go beyond this simple model. The remaining Equations (3.21) describe for a restricted class of pile geometry also the *deformation* of the avalanche. The two sets of the equations are coupled; consequently they cannot, in general, be integrated independently.

Several effects influence the spreading of the avalanche pile. In the downhill direction $B(\partial h/\partial x)$ contributes to a dilatation as $\partial h/\partial x > 0 (< 0)$ in front of (behind) the centre of gravity. The drag forces, on the other hand, contribute to a contraction; interestingly, only the viscous drag, but not the Coulomb friction force contributes to this contraction. Sidewise spreading is

enhanced by the term $D(\partial h/\partial \eta)$ and is reduced by the friction forces; here both viscous and Coulomb-type friction contribute.

4. SIMILARITY SOLUTIONS

4.1. Theoretical considerations

We follow a procedure much the same as in Hutter and others (in press b) and will now establish the governing equations for the longitudinal and sidewise spreading rates. More explicitly, we will derive similarity solutions for a granular pile with elliptical shape in plan view and parabolic distribution of the height. Strictly speaking, such similarity solutions do not exist for the model considered here; however, the reduced Equations (3.21) with the particular assumptions that were invoked to obtain them do permit existence of such similarity solutions.

Similarity solutions were previously constructed, for chute flows along a plane bed by Savage and Hutter (1989) and for chute flows along curved beds by Savage and Nohguchi (1988). The effect of variable friction and of a Voellmy resistive drag were then analyzed by Nohguchi and others (1989) and Hutter and Nohguchi (1990). The first analysis of an unconfined flow along an inclined bed by Hutter and others (in press b) contained a study of similarity solutions in which the Coulomb friction angles ϕ and δ were both kept constant and the Voellmy term was ignored. Here we generalize their solution.

We anticipate a solution in which the shape and difference velocity distributions are preserved, and the profiles are merely stretched or compressed in the streamwise and lateral directions. In view of this, we choose new similarity variables that are normalized by the half length of the pile (see Fig. 3), viz.

$$\nu = \frac{\xi}{g(\tau)}, \mu = \frac{\eta}{f(\tau)}, t = \tau \tag{4.1}$$

from which we may deduce

$$\frac{\partial}{\partial \xi} = \frac{1}{g} \frac{\partial}{\partial \nu}, \frac{\partial}{\partial \eta} = \frac{1}{f} \frac{\partial}{\partial \mu}, \tag{4.2}$$

$$\frac{\partial}{\partial \tau} = \frac{\partial}{\partial t} - \frac{g'}{g} \nu \frac{\partial}{\partial \nu} - \frac{f'}{f} \mu \frac{\partial}{\partial \mu}.$$

Primes denote univariate differentiation. Substituting the expressions (4.2) into the last three of Equations (3.21) and seeking a solution for the difference velocity in the form

$$\tilde{u} = g'\nu, \tilde{v} = f'\mu, \tag{4.3}$$

shows that expressions (4.2) will reduce to the following equation set

$$\begin{aligned} \frac{\partial h}{\partial t} + \frac{h}{g} \frac{\partial \tilde{u}}{\partial \nu} + \frac{h}{f} \frac{\partial \tilde{v}}{\partial \mu} &= 0, \\ \frac{\partial \tilde{u}}{\partial t} &= -\frac{B}{g} \frac{\partial h}{\partial \nu} - 2 \frac{|\tilde{u}_c|}{\Xi h} g' \nu, \\ \frac{\partial \tilde{v}}{\partial t} &= -C \frac{f'}{|\tilde{u}_c|} \mu - \frac{D}{f} \frac{\partial h}{\partial \mu} - \frac{|\tilde{u}_c|}{\Xi h} f' \mu, \end{aligned} \tag{4.4}$$

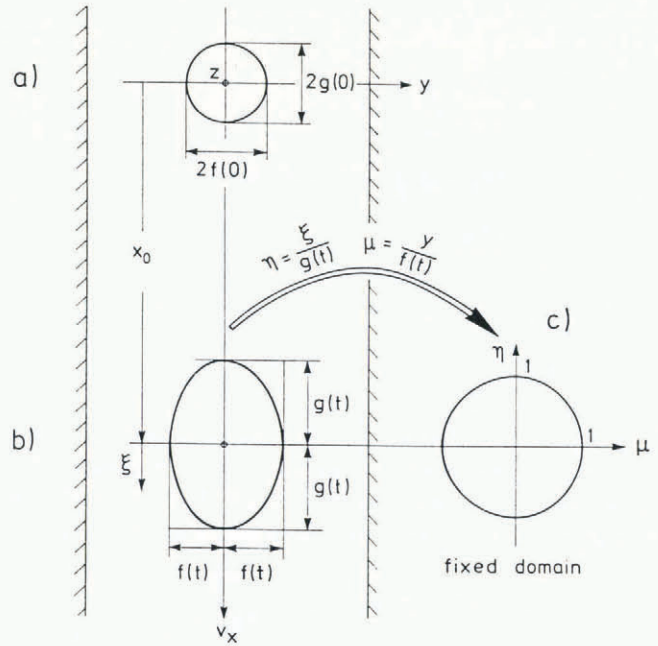


Fig. 3. Top view of the granular pile in the transformed dimensionless coordinates shown in its initial position (a), in a general position (b), and after a fixed-domain transformation (c), explaining the similarity solution and the related coordinate systems.

or after replacement of \tilde{u} and \tilde{v} by Equations (4.3)

$$\frac{\partial h}{\partial \nu} = -\frac{\nu}{B} g \left\{ g'' + 2 \frac{|\tilde{u}_c|}{\Xi h} g' \right\}, \tag{4.5}$$

$$\frac{\partial h}{\partial \mu} = -\frac{\mu}{D} f \left\{ f'' + \left(\frac{C}{|\tilde{u}_c|} + \frac{|\tilde{u}_c|}{\Xi h} \right) f' \right\}. \tag{4.6}$$

Notice that both equations appear in separable form: their righthand sides are each a product of a function of ν (and μ , respectively) and a function of time, t . To solve these equations, assume that the granular mass starts initially from a circular paraboloidal shape that in the course of motion deforms into an ellipsoid with parabolic distribution of its height. More specifically, g and f are interpreted as the principal semi-axes of the ellipse so that the transformation of the ellipsoidal hump into the (ν, μ) -plane by the transformation (4.1) maps the elliptical domain into the interior of a fixed circle with unit radius (see Fig. 3). Assuming, therefore,

$$h = H_0(t)(1 - \nu^2 - \mu^2) \tag{4.7}$$

yields

$$H_0(t) = \frac{g}{2B} \left[g'' + 2 \frac{|\tilde{u}_c|}{\Xi h} g' \right], \tag{4.8}$$

$$= \frac{f}{2D} \left[f'' + \left(\frac{C}{|\tilde{u}_c|} + \frac{|\tilde{u}_c|}{\Xi h} \right) f' \right]. \tag{4.9}$$

If $H_0(t)$ were to be known, then these equations would form two second-order ordinary differential equations for $g(t)$ and $f(t)$, respectively. From the conservation of mass or volume, we can obtain this missing piece of information. Indeed, the total volume V is preserved:

$$\begin{aligned}
 V &= \iint_A h \, d\xi d\eta \\
 &= H_0 \iint_A (1 - \nu^2 - \mu^2) g f \, d\nu d\mu \\
 &= 2\pi H_0 g f \int_0^1 r(1 - r^2) \, dr \\
 &= \frac{\pi H_0 g f}{2}.
 \end{aligned}
 \tag{4.10}$$

With the initial volume V being prescribed, we thus have

$$h = \frac{2V}{\pi g f} (1 - \nu^2 - \mu^2), \quad \bar{h} = \frac{V}{\pi g f}. \tag{4.11}$$

The temporal evolution of the height is thus known, once f and g are determined. The differential equations for these follow from Equations (4.8) and (4.9), when H_0 , as determined from Equations (4.10) and \bar{h} are employed:

$$g'' + 2 \frac{|\bar{u}_c| \pi f g}{\bar{\Xi} V} g' = \frac{4BV}{\pi g^2 f}, \tag{4.12}$$

$$f'' + \left(\frac{C}{|\bar{u}_c|} + \frac{|\bar{u}_c| \pi f g}{\bar{\Xi} V} \right) f' = \frac{4DV}{\pi g f^2}.$$

There remains the corroboration that with the representations (4.3) the local mass balance in Equation (4.4)₁ is identically satisfied. This demonstration is routine and will be left to the reader. In summary, we must solve Equations (3.21)_{1,2} for the motion of the centre of mass together with Equations (4.12) for the longitudinal and transverse spreading rates. The integration of these equations must be performed with initial conditions

$$\begin{aligned}
 \bar{u}_c(t=0) &= \bar{u}_c^0, \\
 \bar{x}_c(t=0) &= \bar{x}_c^0,
 \end{aligned}
 \tag{4.13}$$

$$g(t=0) = g_0, \quad g'(t=0) = g'_0,$$

$$f(t=0) = f_0, \quad f'(t=0) = f'_0,$$

corresponding to a granular mass with initial length $2g_0[L_x]$ and initial width $2f_0[L_y]$.

4.2. Numerical integration

To complete the solution, the above initial value problem must be solved numerically. To this end, we transform the system of Equations (3.21)_{1,2} and (4.12) to standard form and use t instead of τ :

$$\begin{aligned}
 \frac{d\bar{x}_c}{dt} &= \bar{u}_c, \\
 \frac{d\bar{u}_c}{dt} &= \sin \zeta - C \operatorname{sgn}(\bar{u}_c) - \frac{\bar{u}_c |\bar{u}_c| \pi f g}{\bar{\Xi} V}, \\
 \frac{dg}{dt} &= g', \\
 \frac{dg'}{dt} &= \frac{4BV}{\pi g^2 f} - 2 \frac{|\bar{u}_c| \pi f g}{\bar{\Xi} V} g', \\
 \frac{df}{dt} &= f', \\
 \frac{df'}{dt} &= \frac{4DV}{\pi g f^2} - \left(\frac{C}{|\bar{u}_c|} + \frac{|\bar{u}_c| \pi f g}{\bar{\Xi} V} \right) f'.
 \end{aligned}
 \tag{4.14}$$

In writing down these equations, we have also used the relation $V = \bar{h} \pi f g$. Furthermore, the coefficients B, C and D are defined in Equations (2.17). We note the following properties of the system (4.14):

- (i) Without the viscous sliding term ($\bar{\Xi} \rightarrow \infty$), the motion of the centre of mass of the pile decouples from that of the deformation.
- (ii) There can never be a rigid-body motion (for which g and f are constant) even for finite $\bar{\Xi}$. This follows from Equations (4.14)_{4,6}. A rigid-body motion would require $g' = f' = dg'/dt = df'/dt = 0$, but the statements $df'/dt = 0$ and $dg'/dt = 0$ are not possible with $f' = 0$ since the righthand sides of Equations (4.14)_{4,6} are positive when $f' = 0, g' = 0$.
- (iii) The centre of mass motion is never steady unless $\bar{\Xi} \rightarrow \infty$.
- (iv) When the bed-friction angle, δ , varies linearly between the avalanche front and trailing edge, then Equation (4.14)₄ has on its righthand side an additional term $-\Delta C/2$. Without a similar term also arising in Equation (4.14)₆, rigid-body and steady motions still cannot exist.

We thus have proved that *within the context of similarity solutions, our model equations that incorporate Coulomb-type and Voellmy-type resistive drag cannot exhibit rigid-body motion. Likewise, the centre of mass cannot reach a finite steady speed. Alternatively, when the Voellmy drag is ignored, rigid-body motions are still not possible, but a steady, finite velocity can be reached in this case.* These results contrast with earlier results of Hutter and Nohguchi (1990), who demonstrated for chute flows existence of rigid-body motions and finite steady speeds. The reason for the difference in behaviour is the influence of the sidewise spreading. Equations (4.14) contain the coefficients B, C and D , which depend on the Earth pressure coefficients $k_{\text{actpass}}^{x,y}$ that are given by

$$\begin{aligned}
 k_{\text{actpass}}^x &= \begin{cases} k_{\text{act}}^x, & \text{for } g' > 0, \\ k_{\text{pass}}^x, & \text{for } g' < 0, \end{cases} \\
 k_{\text{actpass}}^y &= \begin{cases} k_{\text{act}}^y, & \text{for } f' > 0, \\ k_{\text{pass}}^y, & \text{for } f' < 0. \end{cases}
 \end{aligned}
 \tag{4.15}$$

Therefore, depending upon whether the flow is extending or compressing, these coefficients take different values. Equations (4.14) must be solved subject to the initial conditions (4.13). A straight-forward integration, however, is not possible when $\bar{u}_c^0 = 0$, because the system is singular at such early times. For those early times, a power-series solution must be sought when $\bar{u}_c^0 = 0$. This solution reads

$$\left. \begin{aligned}
 \bar{x}_c &= \bar{x}_c^0 + u_1 \frac{t^2}{2} + u_3 \frac{t^4}{24} + \dots, \\
 \bar{u}_c &= u_1 t + u_3 \frac{t^3}{6} + \dots, \\
 g &= g_0 + g_0' t + g_2 \frac{t^2}{2} + g_3 \frac{t^3}{6} + \dots, \\
 f &= f_0 + f_2 \frac{t^2}{2} + \dots,
 \end{aligned} \right\} \text{as } t \rightarrow 0 \tag{4.16}$$

in which

$$\begin{aligned}
 u_1 &= \sin \zeta - C, \\
 u_3 &= -2 \frac{\pi}{\Xi V} (\sin \zeta - C)^2 g_0 f_0, \\
 g_2 &= \frac{4BV}{\pi g_0^2 f_0}, \\
 g_3 &= \frac{8BV}{\pi g_0^3 f_0} - \frac{2\pi}{\Xi V} g_0 f_0 g_0' u_1, \\
 f_2 &= \frac{4DV}{\pi g_0 f_0^2 (1 + (C/u_1))}.
 \end{aligned}
 \tag{4.17}$$

Notice that this early time solution starts with a centre of mass velocity at rest and with $f'(t=0) = f'_0 = 0$. This latter condition is necessary for consistency; usual initial conditions require $f'_0 = 0$ and $g'_0 = 0$. Notice also that the short-time solution in Equations (4.16) breaks down when $\Xi \rightarrow \infty$. This case is physically unrealistic anyhow and will not be considered.

5. RESULTS

In the construction of similarity solutions, it was assumed that $\bar{u} \geq 0$ everywhere within the moving pile. At early times, in a motion starting from rest ($\bar{u}_c = 0$: no centre of mass velocity, $f'_0 = 0, g'_0 = 0$: neither sidewise nor longitudinal initial spreading), this condition could be violated unless $g' < \bar{u}_c$. In view of Equations (4.16) and (4.17), this condition reads

$$\lim_{t \rightarrow 0} \left(\frac{g'}{\bar{u}_c} \right) = \frac{4BV}{\pi f_0 g_0^2 (\sin \zeta - C)} < 1, \quad (g'_0 = 0) \tag{5.1}$$

or when invoking Equations (2.17) and (4.10)

$$\epsilon_x \leq \frac{1}{2k_{act}^x} \frac{g_0}{H_0} (\tan \zeta - \tan \delta). \tag{5.2}$$

H_0 and g_0 are the initial values of the dimensionless depth and longitudinal semi-span, both of which are usually chosen as 1. In a motion along a plane-inclined surface, the flow cannot be compressing; so only k_{act}^x applies. Values for k_{act}^x that depend on δ and ϕ are between 0.3 and 2.0 for $0 \leq \delta \leq 35^\circ$ and $0 \leq \phi \leq 35^\circ$. Generally, k_{act}^x decreases with increasing ϕ . More generally, the right-hand side of inequality (5.2) grows with increasing ϕ and ζ .

Values of ϵ_x that do conform with inequality (5.2) are conservative because they imply that the condition $u > 0$ is *nowhere* violated. One may relax this condition a bit and tolerate, at very early times and close to the upper margin, a violation. For $\zeta > 40^\circ$ and $\delta < 30^\circ$, condition (5.2) is almost always satisfied when $\epsilon_x < 0.5$. We shall present results only for values of ϵ_x for which inequality (5.2) is obeyed.

Usual pile geometries have $[L_y] \leq [L_x]$. We express this by the following orderings:

$$\epsilon_y = \epsilon_x^\alpha, \quad \epsilon_{xy} = \epsilon_x^{1-\alpha}, \quad \frac{\epsilon_y}{\epsilon_{xy}} = \epsilon_x^{2\alpha-1}$$

with $0 \leq \alpha \leq 1$; $\alpha = 1$ means that $[L_x] = [L_y]$, and $\alpha > 0.5$ is requested for the approximation to be valid. Most ensuing computations will be performed for $\alpha = 1$. Physically, it is not important what values are chosen for the parameters *once* the approximations have been carried through. However, if these parameters are chosen at will, then it is only meaningful to compare quantities with physical dimensions, for the non-dimensionalized quantities do then not stay in the proportions the physical

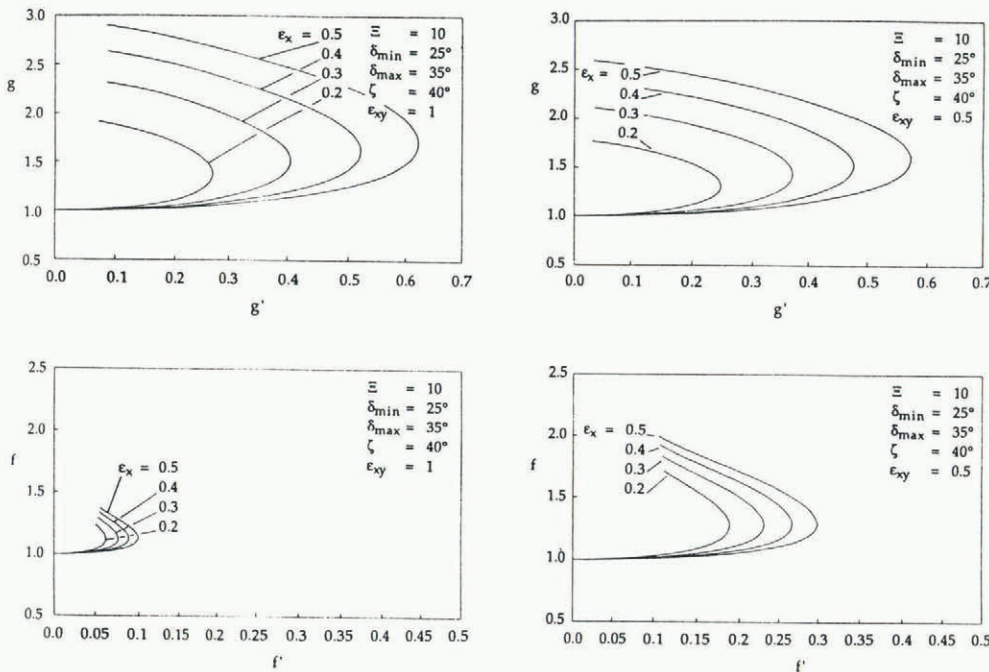


Fig. 4. Phase diagrams g versus g' or f versus f' plotted for an angle of internal friction $\phi = 35^\circ$ and the physical parameters as shown in the insets and for the indicated values of the aspect ratios $\epsilon_x, \epsilon_{xy}$. The left panels show $g(g')$ (top) and $f(f')$ (bottom) for $\epsilon_{xy} = 1$ and hence $\epsilon_y = \epsilon_x$, the panels on the right show the same for $\epsilon_{xy} = 0.5$. The inclination angle is $\zeta = 40^\circ$.

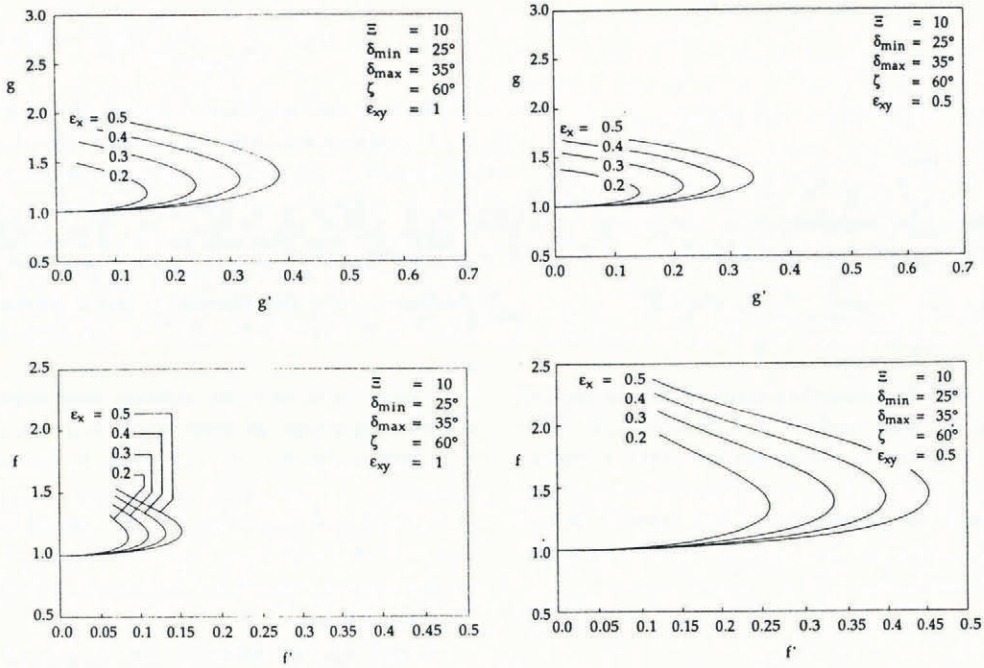


Fig. 5. Same as Figure 4 but now for an inclination angle $\zeta = 60^\circ$.

quantities will. We shall make clear which choices are taken.

In the graphical representation of the results, two types of figures are presented. The first type is a kind of phase diagram in which g is plotted against g' and f against f' for various values of the parameters. In the second type of figure, the spreads g or f or the velocity \bar{u}_c and position \bar{x}_c of the centre of mass are plotted as functions of dimensionless time, again for various values of the parameters involved.

In all following computations, the angle of internal friction was $\phi = 35^\circ$. Figure 4 shows phase-plane

trajectories for g and f , respectively, for a slope angle $\zeta = 40^\circ$, $\epsilon_{xy} = 1$ (left panels) and $\epsilon_{xy} = 0.5$ (right panels), a Voellmy coefficient $\Xi = 10$ and bed-friction angles $\delta_{\text{front}} = 35^\circ, \delta_{\text{rear}} = 25^\circ$, and the aspect ratio ϵ_x as indicated. All phase-space trajectories start at $g_0 = f_0 = 1$ and $g'_0 = f'_0 = 0$. Figure 5 shows the same for a slope angle $\zeta = 60^\circ$.

It can be seen that both the longitudinal and the sideways spreading depend strongly upon the variation of the geometry parameters ϵ_x (and ϵ_{xy}). The general tendency is the same as was observed when $\Xi = 0$ and $\Delta\delta = \delta_{\text{front}} - \delta_{\text{rear}} = 0$ (see Hutter and others (in press b)

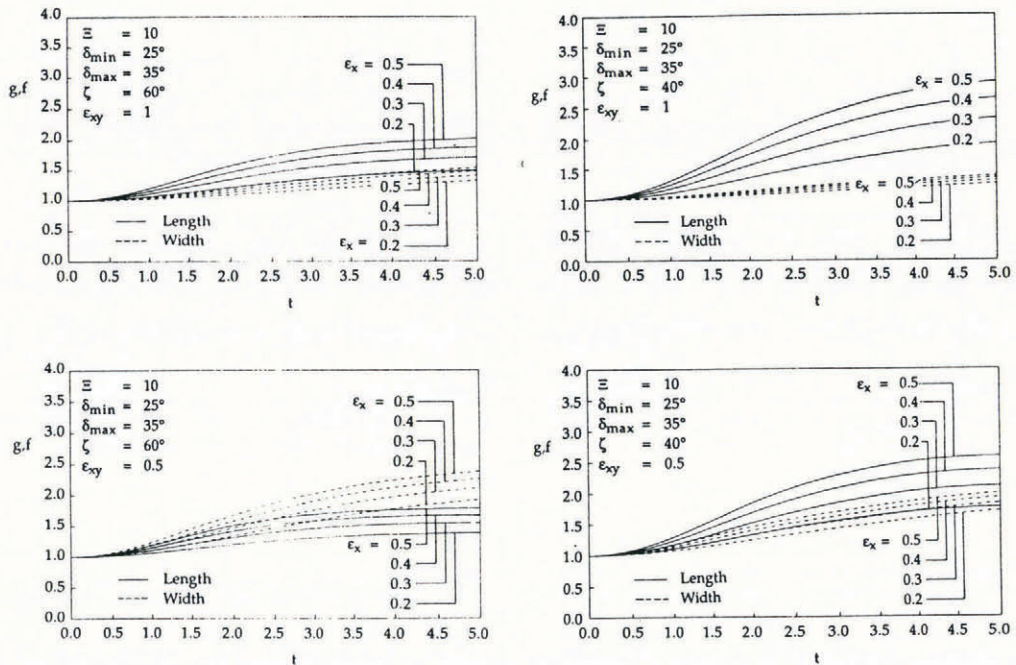


Fig. 6. Temporal evolutions (dimensionless) of the semi-spreads $g(t)$ (solid) and $f(t)$ (dashed) calculated for $\phi = 35^\circ$ and the physical parameters as shown in the insets. Panels on the left are for $\zeta = 60^\circ$, those on the right for $\zeta = 40^\circ$; top panels are for $\epsilon_{xy} = 1$, those on the bottom for $\epsilon_{xy} = 0.5$.

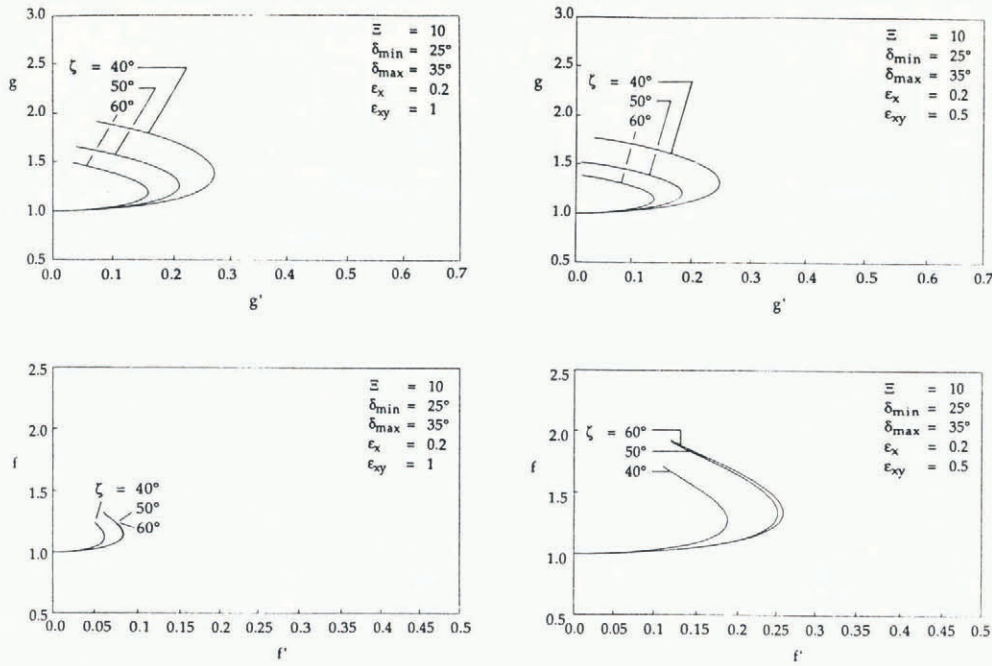


Fig. 7. Same as Figure 4, for $\phi = 35^\circ$ and the physical parameters shown in the insets; however, now the aspect ratios $\epsilon_x, \epsilon_{xy}$ are held fixed and the slope angle ζ is varied as indicated. The left panels show the result for $\epsilon_{xy} = 1$, those on the right for $\epsilon_{xy} = 0.5, \epsilon_x = 0.2$.

for comparison); longitudinal spreadings are bigger than sidewise spreadings and their rates are larger when ϵ_x is increased. Only for $\epsilon_{xy} = 0.5$ and steep slopes ($\zeta = 60^\circ$; Figure 5, right panels) sidewise spreading is bigger than longitudinal spreading. Thus, even though the condition $\epsilon_{xy} = O(1)$ must be fulfilled in order that the model equations are valid, the values of the initial aspect ratios are important for the geometry of the evolving masses. Granular masses which develop from relatively compact geometries spread faster than shallow masses do, and this spreading is generally larger in the longitudinal than in the transverse direction (exception again $\zeta = 60^\circ$,

$\epsilon_{xy} = 0.5$). Figure 6, which shows temporal evolutions of $g(t)$ and $f(t)$ for $\zeta = 60^\circ$ (left panels) and $\zeta = 40^\circ$ (right panels) indicates this very clearly. Depending on the value of the slope angle ζ and the aspect ratios $\epsilon_x, \epsilon_{xy}$, a granular pile developing from a circular shape will develop into an ellipse whose major semi-axis is either in the longitudinal (regular case) or transverse ($\zeta = 60^\circ, \epsilon_{xy} = 0.5$) direction.

One distinctive feature of the phase diagrams not exhibited when $\Delta\delta = 0$ and $\Xi = 0$ is the fact that the trajectories $g(g')$ and $f(f')$ “bend towards the ordinate”. Thus, g' and f' reach an absolute maximum whose value

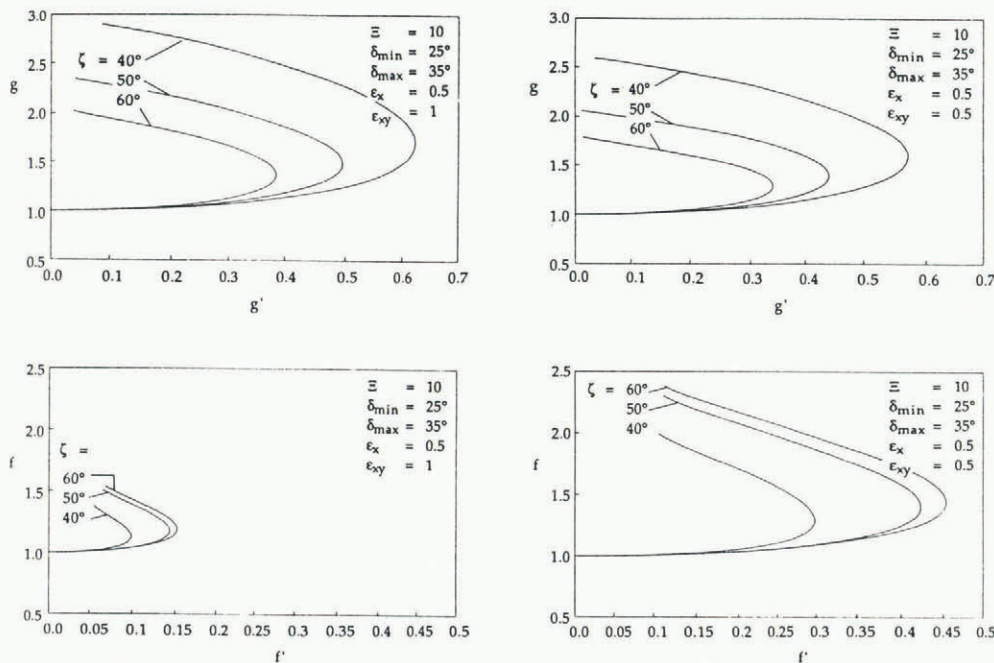


Fig. 8. Same as Figure 7 but now for $\epsilon_x = 0.5$.

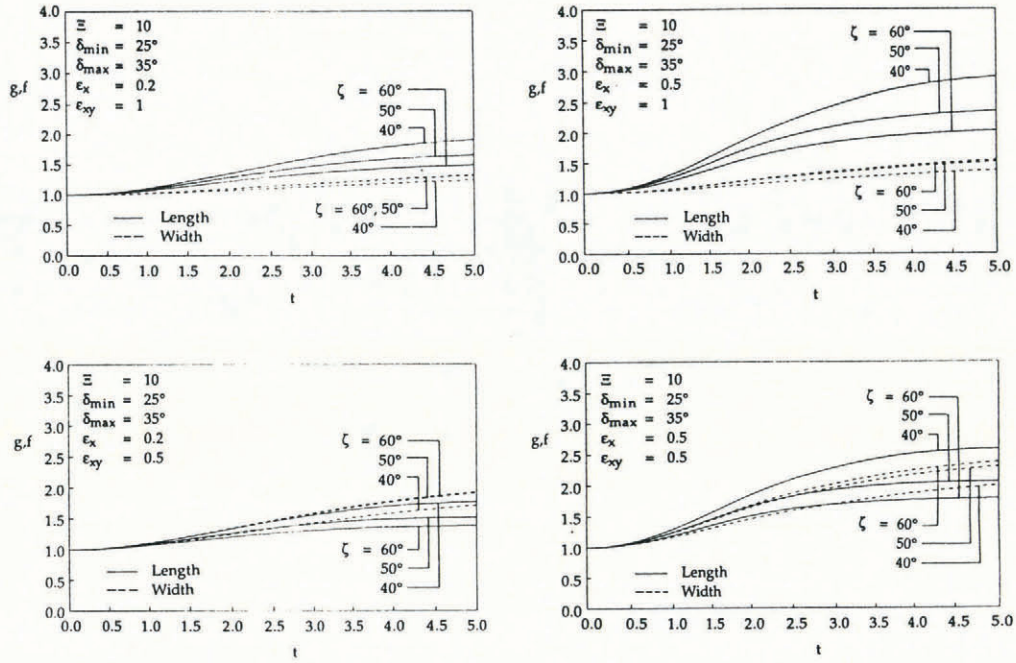


Fig. 9. Same as Figure 6, but now the inclination angle ζ is varied. Parameters as shown in the insets. Left panels are for $\epsilon_x = 0.2$, those on the right for $\epsilon_x = 0.5$.

depends on ϵ_x and ϵ_{xy} . Figures 4 and 5 indicate that the g - and f -trajectories bend and approach the ordinate and do this faster, the smaller ϵ_x is. If both trajectories reach the ordinate simultaneously, then the motion is of rigid-body type: the granular avalanche would then move like a rigid body. Figures 4, 5 and 6 show that such a rigid-body motion is approached but not reached. Deformations are, however, smaller the smaller ϵ_x is. Thus, flat and shallow piles deform less than do compact piles. Interesting to observe, and physically obvious, is the fact that sidewise spreading is significant only when $\epsilon_y = \epsilon_x / \epsilon_{xy}$ is relatively large. For $\epsilon_y = 1$ ($\epsilon_x = 0.5$,

$\epsilon_{xy} = 0.5$), the spreading is primarily in the transverse direction and actual ratios quickly reach small values. So, whereas in these instances the prerequisites of the model equations are not satisfied initially, the evolution is such that they are more and more fulfilled in the course of motion.

Similar qualitative behaviour is also seen in Figures 7 and 8 where phase diagrams $g(g')$ and $f(f')$ are shown for $\epsilon_{xy} = 1$ (left panels) and $\epsilon_{xy} = 0.5$ (right panels) and two different values of ϵ_x ($\epsilon_x = 0.2$, Fig. 7; $\epsilon_x = 0.5$, Fig. 8) parameterized for slope angles $\zeta = 40^\circ, 50^\circ$ and 60° . According to the figures, the steeper the slope is the

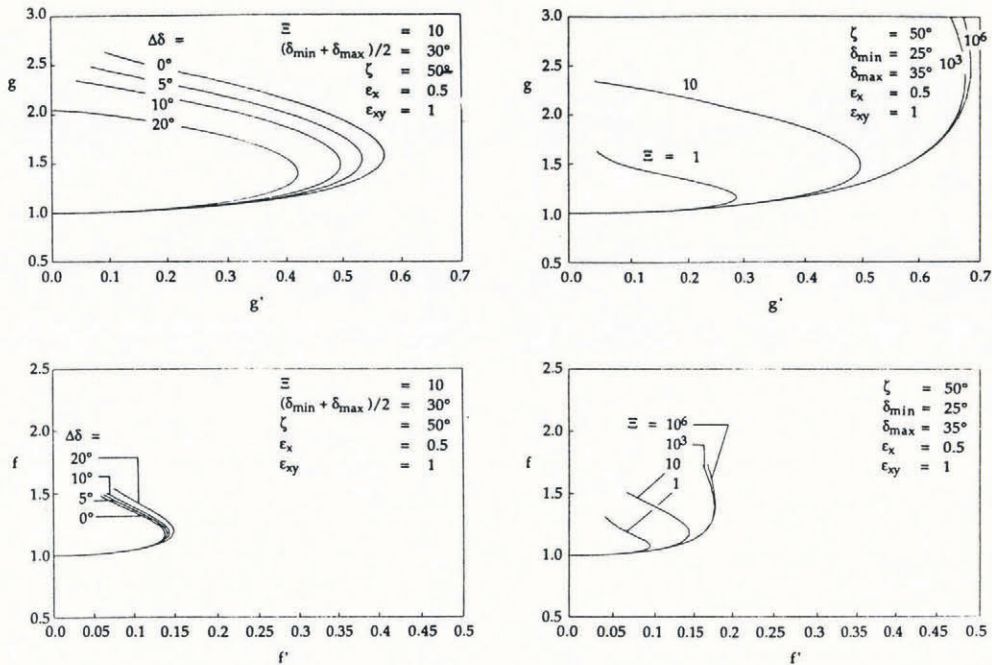


Fig. 10. Phase diagrams g versus g' and f versus f' plotted for an internal friction angle $\phi = 35^\circ$ and the physical parameters as shown in the insets. In the left panels $\Delta\delta = \delta_{\text{front}} - \delta_{\text{rear}}$ is varied, while $\Xi = 10$ is held fixed. In right panels $\Delta\delta = 10^\circ$ while Ξ is varied. Calculations are for $\epsilon_x = \epsilon_y = 0.5$.

smaller (larger) will the longitudinal (transverse) spreading become. Moreover, the aspect ratios have again a significant influence on the spreading rates. When $\epsilon_{xy} = 1$, longitudinal spreading is clearly larger than sidewise spreading; when $\epsilon_{xy} = 0.5$, this is not necessarily so as can be clearly inferred from the lower two panels in Figure 9. Not surprising is also the fact that the spreading generally decreases with increasing slope angle.

Of particular interest is the quantitative analysis of the similarity solution to the granular-avalanche problem when, first the bed-friction angle is varied from the front to the rear end of the moving pile or, secondly, the Voellmy drag coefficient is changed. Figures 10 and 11 collect results in this regard. In the left panels of Figure 10 the g and f trajectories are shown where the Voellmy coefficient Ξ is held constant ($\Xi = 10$), but $\Delta\delta = \delta_{\text{front}} - \delta_{\text{rear}}$ is varied between 0° and 20° . It is seen that both longitudinal and sidewise spreadings are somewhat affected, but that the qualitative behaviour is unchanged. The trajectories bend over towards the ordinate axis. Generally, the smaller $\Delta\delta$ is the less will the longitudinal spreading be inhibited. The sidewise behaviour is opposite but less pronounced.

Quite contrary to this behaviour of the moving and deforming granular pile is its response to variations in the Voellmy parameter Ξ (right panels in Figure 10). Both, the longitudinal and the transverse spreadings are affected by the amount of viscous-type friction; however, the influence to the sidewise spreading is less dramatic. The computations for the graphs in Figure 10 have been done for $\epsilon_x = 0.5$ and $\epsilon_{xy} = 1$. For smaller values of ϵ_x (but $\epsilon_{xy} = 1$), these effects are less pronounced, but for larger values they are enhanced. Similarly, a decrease of ϵ_{xy} enhances these effects. This can be seen, in parts in Figure 11, which shows the temporal evolutions of the semi-spreads for several values

of $\Delta\delta$ (left panels) and Ξ (right panels) and value $(\epsilon_x, \epsilon_{xy}) = (0.5, 1)$ (top) and $(\epsilon_x, \epsilon_{xy}) = (0.5, 0.5)$ (bottom). An increase in $\Delta\delta$, which corresponds to a more efficient ploughing, hinders a spreading though not dramatically.

A variation of Ξ (by several powers) affects the spreading, both longitudinally and transversely, in a considerable manner. This behaviour is understandable, if one considers the governing differential Equations (4.14). The Voellmy term affects the spreadings in Equations (4.14)_{2,4,6} and its influence becomes vanishingly small when $\Xi \rightarrow \infty$. In Figure 11 (right panels) the role played by the Voellmy drag on the evolution of the spreadings $g(t)$ and $f(t)$ is very clearly seen.

As can be surmised from the dependence of the equations of motion on Ξ , there must also be a strong dependence of the centre of mass motion on the Voellmy coefficient. Figure 12 provides evidence for this. In the top two panels, the centre of mass position, \bar{x}_c , in the lower panels the centre of mass velocity, \bar{u}_c , (both dimensionless) are plotted against dimensionless time, t , for various values of the parameter Ξ . For very large Ξ values ($\Xi > 1000$), the centre of mass velocity is essentially linear in time and its position grows quadratically. These results are an important corroboration of our earlier calculations which were performed without the Voellmy term. With growing viscosity (decreasing Ξ values), the growth of the centre of mass velocity is more and more reduced. The velocity can even go through an absolute maximum and decrease afterwards, so that a decelerating motion of the centre of mass is possible, in principle. The fact that the graphs for $\epsilon_{xy} = 1$ and $\epsilon_{xy} = 0.5$ hardly differ is an indication that the centre of mass motion is only little affected by the amount of spreading. On the other hand, that the amount of spreading crucially depends on both the dry and viscous drag behaviours, demonstrates

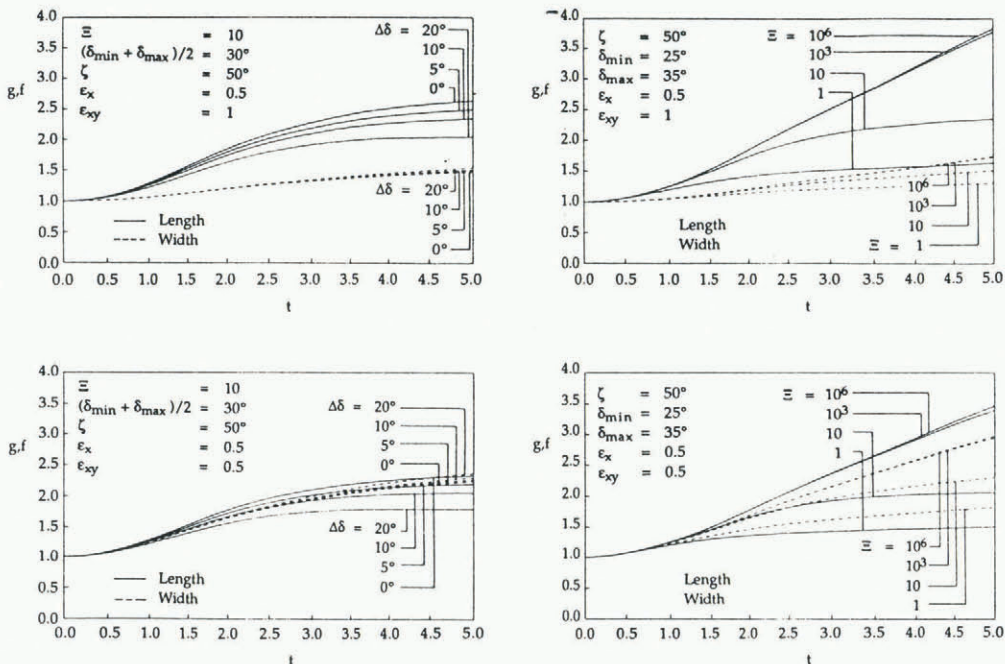


Fig. 11. Temporal evolutions of the semi-spreads $g(t)$ (solid) and $f(t)$ (dashed), calculated for $\phi = 35^\circ$ and the physical parameters shown in the insets. In the left panels $\Delta\delta = \delta_{\text{front}} - \delta_{\text{rear}}$ is varied while Ξ is held fixed, in the right panels it is the reverse. $(\epsilon_x, \epsilon_{xy}) = (0.5, 1)$ (top) and $(\epsilon_x, \epsilon_{xy}) = (0.5, 0.5)$ (bottom).

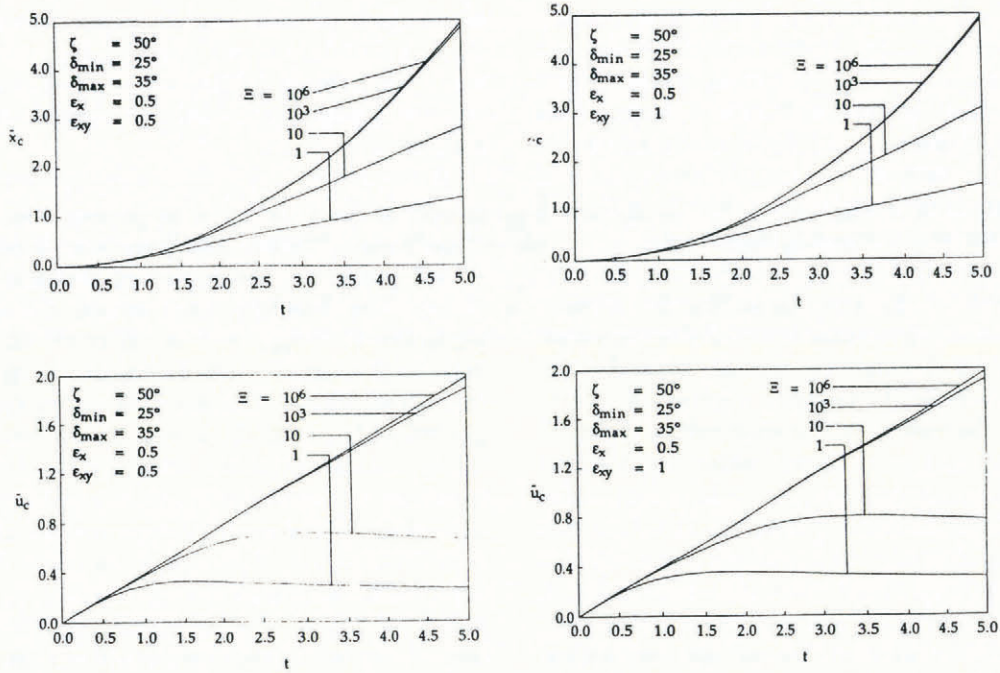


Fig. 12. Temporal evolution of the centre of mass position, \bar{x}_c and its velocity, \bar{u}_c (both dimensionless) calculated for $\phi = 35^\circ$ and the physical parameters shown in the insets, when the Voellmy coefficient is varied. Computations have been done for $(\epsilon_x, \epsilon_{xy}) = (0.5, 0.5)$ (left) and $(\epsilon_x, \epsilon_{xy}) = (0.5, 1.0)$ (right).

the superiority of the present model over the classical avalanche models due to Voellmy, Salm and others.

6. CLOSING REMARKS

In this paper, we have been concerned with the motion of a finite mass of a granular material down an inclined plane that is released from rest and may freely spread in the longitudinal and transverse directions as it moves down its track. The granular mass was treated as a cohesionless Coulomb-like continuum with a basal friction law in which the shear traction is additively composed of a Coulomb-like drag and a viscous drag proportional to the squared velocity. We believe that such a model is a valid one for the study of the dynamics of flow avalanches for the following reasons: the common mass point or hydraulic models that are used incorporate physically essentially the same complexity as this one, except that this one allows for the variation of the frictional drag within the avalanche. Mathematically or geometrically, this model has greater flexibility, as it permits a longitudinal and a sideways spreading, both of which are not present in the Voellmy, Salm, Perla, etc. models. Thus, this model is more general than the former, yet incorporates essentially the same physics.

We used the depth-averaged equations of Hutter and others (in press b). These equations are scaled, and dimensionless spatially two-dimensional evolution equations are derived for the distribution of the avalanche depth and depth averages of the velocity field of which analytical solutions are not likely to be found. In an approximate treatment, however, at least semi-analytical solutions were determined. To find these solutions, the motion was split into the motion of the centre of gravity of the pile plus a deformation from it. Such a decomposition

cannot be achieved without additional ad hoc assumptions. The semi-analytical solutions are the so-called similarity solutions, and they enjoy the property of preserving the shape. For their existence, the granular avalanche must start from an initial shape in the form of an elliptical or circular paraboloid, and this shape will be preserved during the motion, and only its aspect ratio will change. For the existence of such similarity solutions, however, additional mild assumptions are needed: the earth-pressure coefficient must be assumed to be constant and the sidewise variation of the bed-friction angle must be ignored. Moreover, the construction of the similarity solutions is based on the formal separation of the evolution equations into one set governing the motion of the centre of mass and another governing the deformation. In general, this separation is not possible; however, with the approximations incorporated, two equation sets are obtained which are integrable. We explicitly listed the restrictions to make the reader aware about the limitations of these solutions.

The physical parameters that govern the model are the slope angle ζ , the initial depth to length ratio of the pile, ϵ_x , the ratio of the width to the length, ϵ_{xy} , necessarily of order unity, the internal angle of friction ϕ , the basal-friction angle δ and the drag coefficient of the viscous sliding law Ξ , called the Voellmy parameter. The effect of these parameters on the spreading rate of the granular pile that develops from a circular paraboloid led to results which may be summarized as follows:

- (i) For the two-dimensional spreading, there exists no rigid-body motion, i.e. the length and width of the elliptical pile will always vary in time, no matter what the values of the bed-friction angle and the coefficient of viscous drag are.
- (ii) As an immediate consequence of the above

statement, there can be no steady motion of the centre of mass of the granular pile. However, for values of the dimensionless viscous-drag coefficient $\Xi < 10^3$ the centre of mass motion is greatly affected by the value of Ξ . The motion can be accelerating or decelerating, or oscillate between the two.

- (iii) The geometry of the pile depends on both the variation of the bed-friction angle with position (linear variations of $\tan \delta$ in the long direction were studied) as well as on the value of the viscous-drag coefficient. However, the aspect ratio of the moving pile is influenced more by the latter than by the former.

In which way are these solutions useful to the avalanche dynamicist? First, they are capable of providing physical insight into the behaviour of a deforming finite mass of snow, in a way previous models did not. For small aspect ratios ϵ_x and ϵ_y , the model indicates small deformation. Under such conditions, a rigid-body assumption for the motion of a finite mass of snow does not seem to be too drastic a simplification. On the other hand, viscous sliding affects the deformation (spreading) considerably while ploughing does less. These inferences are qualitative and provide physical insight.

Secondly, this model could be used in actual computations along curved avalanche paths to predict deposition and its areal extent. While such a procedure will certainly be useful, we do not believe it to be accurate. In fact, laboratory experiments on the motion of a finite mass of a cohesionless granular material down an inclined plane show that similarity solutions are not reproduced. Plan views of moving granular avalanches that develop from a circular geometry rather develop into tear-drop shapes and along curved beds the shapes are even more complicated. One may therefore think that our solutions are of little use. We do not think so, as our solutions provide physical insight into the *basic mechanisms* of the motion and spreading of a granular pile; the model does have diagnostic value. For a prognostic use of the governing equations, however, integration from more general initial configurations are needed. Such studies are under way.

REFERENCES

- Alean, J. 1984. Untersuchungen über Entstehungsbedingungen und Reichweiten von Eislawinen. *Eid. Tech. Hochschule, Zürich. Versuchsanst. Wasserbau, Hydrol. Glaziol. Mitt.* 74.
- Alean, J. 1985. Ice avalanche activity and mass balance of a high altitude hanging glacier in the Swiss Alps. *Ann. Glaciol.*, **6**, 248–249.
- Beghin, P., E.J. Hopfinger and R.E. Britter. 1981. Gravitational convection from instantaneous sources on inclined boundaries. *J. Fluid Mech.*, **107**, 407–422.
- Greve, R. and K. Hutter. 1993. Motion of a granular avalanche in a convex and concave curved chute: experiments and theoretical predictions. *Phil. Trans. R. Soc. Lond., Ser. A*, **342**, 573–600.
- Gubler, H. 1987. Measurements and modelling of snow avalanche speeds. *International Association of Hydrological Sciences Publication* 162 (Symposium at Davos 1986 — *Avalanche Formation, Movement and Effects*), 405–420.
- Gubler, H. 1989. Comparison of three models of avalanche dynamics. *Ann. Glaciol.*, **13**, 82–89.
- Gubler, H. Unpublished. Messungen an Fließlawinen. Eidgenössisches Institut für Schnee- und Lawinenforschung (EISLF), Weissfluhjoch/Davos, Switzerland. Interner Bericht 600, 1981.
- Gubler, H. and M. Hiller. 1984. The use of microwave FMCW radar in snow and avalanche research. *Cold Reg. Sci. Technol.*, **9**(2), 109–119.
- Hermann, F. and K. Hutter. 1991. Laboratory experiments on the dynamics of powder-snow avalanches in the run-out zone. *J. Glaciol.*, **37**(126), 281–295.
- Hopfinger, E.J. 1983. Snow avalanche motion and related phenomena. *Annu. Rev. Fluid Mech.*, **15**, 47–76.
- Hopfinger, E.J. and P. Beghin. 1980. Buoyant clouds appreciably heavier than the ambient fluid on sloping boundaries. In *IAHR. International Association for Hydraulic Research. Second International Symposium on Stratified Flows, Trondheim*, 495–504.
- Hsü, K. 1975. On sturzstroms — catastrophic debris streams generated by rockfalls. *Geol. Soc. Am. Bull.*, **86**, 129–140.
- Hsü, K. 1978. Albert Heim: Observations on landslides and relevance to modern interpretations. In Voight, B., ed. *Rockslides and avalanches, 1. Natural phenomena*. Amsterdam, etc., Elsevier, 69–93.
- Hutter, K. 1992. Lawinendynamik, eine Übersicht. *Schweizer Ingenieur und Architekt*, **13**, 259–269.
- Hutter, K. and T. Koch. 1991. Motion of a granular avalanche in an exponentially curved chute: experiments and theoretical predictions. *Philos. Trans. R. Soc. London, Ser. A*, **334**, 93–138.
- Hutter, K. and Y. Nohguchi. 1990. Similarity solutions for a Voellmy model of snow avalanches with finite mass. *Acta Mech.*, **82**, 99–127.
- Hutter, K. and S.B. Savage. 1988a. Avalanche dynamics: the motion of a finite mass of gravel down a mountain side. *Proceedings of the 5th International Symposium on Landslides, July 10–15, 1989, Lausanne, Switzerland*. In Bonnard, C., ed., 691–697.
- Hutter, K. and S.B. Savage. 1988b. Granular avalanches. Theory and laboratory experiments. A review of new developments. *Internationale Symposium, Interpret 1988, Graz. Tagungs-publikation*, Vol.3, 251–266.
- Hutter, K., T. Koch, C. Plüss and S.B. Savage. In press a. Dynamics of avalanches of granular materials from initiation to runout. Part II. Laboratory experiments. *Acta Mech.*
- Hutter, K., N. Siegel, S.B. Savage and Y. Nohguchi. In press b. Two dimensional spreading of a granular avalanche down an inclined plane. Part I. Theory. *Acta Mech.*
- Lang, T.E. and M. Martinelli, Jr. 1979. Application of numerical transient fluid dynamics to snow avalanche flow. Part II. Avalanche modeling and parameter error evaluation. *J. Glaciol.*, **22**(86), 117–126.
- Lang, R.M., B.R. Leo and K. Hutter. 1989. Flow characteristics of an unconstrained, non-cohesive, granular medium down an inclined curved surface. *Ann. Glaciol.*, **13**, 146–153.

- Nohguchi, Y., K. Hutter and S.B. Savage. 1989. Similarity solutions for granular avalanches of finite mass with variable bed friction and rigid body motion. *Continuum Mechanics and Thermodynamics*, **1**, 239–265.
- Norem, H. and K. Kristensen. 1988. *The Ryggfjonn project. Avalanche data from the winter 1987/88*. Oslo, Norwegian Geotechnical Institute. (Report 58120-12.)
- Norem, H., K. Kristensen and K. Tronstad. 1986. *The Ryggfjonn project. Avalanche data from the winter 1984/85*. Oslo, Norwegian Geotechnical Institute. (Report 58120-8.)
- Norem, H., K. Kristensen and K. Tronstad. 1988. *The Ryggfjonn project. Avalanche data from the winter 1986/87*. Oslo, Norwegian Geotechnical Institute. (Report 58120-10.)
- Perla, R. I. and M. Martinelli, Jr. 1978. Avalanche handbook. *U.S. Dep. Agric. For. Serv. Agric. Handb.* 489.
- Perla, R. I., T. T. Cheng and D. M. McClung. 1980. A two-parameter model of snow avalanche motion. *J. Glaciol.*, **26**(94), 197–207.
- Salm, B. 1966. Contribution to avalanche dynamics. *International Association of Scientific Hydrology Publication* 69 (Symposium at Davos 1969 — *Scientific Aspects of Snow and Ice Avalanches*), 199–214.
- Salm, B. 1968. On nonuniform, steady flow of avalanching snow. *International Association of Scientific Hydrology Publication* 79 (General Assembly of Bern 1967 — *Snow and Ice*), 19–29.
- Savage, S. B. and K. Hutter. 1989. The motion of a finite mass of granular material down a rough incline. *J. Fluid Mech.*, **199**, 177–215.
- Savage, S. B. and K. Hutter. 1991. Dynamics of avalanches of granular materials from initiation to runout. Part I. Analysis. *Acta Mech.*, **86**, 201–223.
- Savage, S. B. and Y. Nohguchi. 1988. Similarity solutions for avalanches of granular materials down curved beds. *Acta Mech.*, **75**, 153–174.
- Scheiwiller, T. 1986. Dynamics of powder snow avalanches. (Ph.D. thesis, Eidgenössische Technische Hochschule, Zürich.)
- Scheiwiller, T. and K. Hutter. 1982. Lawindynamik. Übersicht über Experimente und theoretische Modelle von Fliess- und Staublawinen. *Eid. Tech. Hochschule, Zürich. Versuchsanst. Wasserbau, Hydrol. Glaziol. Mitt.* 58.
- Scheiwiller, T., K. Hutter and F. Hermann. 1987. Dynamics of powder snow avalanches. *Annales Geophysicae*, **5B**(6), 569–588.
- Tochon-Danguy, J.-C. 1977. Étude des courants de gravité sur forte pente avec application aux avalanches poudreuses. (Thèse, Université de Grenoble.)
- Tochon-Danguy, J.-C. and E. J. Hopfinger. 1975. Simulation of the dynamics of powder avalanches. *International Association of Hydrological Sciences Publication* 114 (Symposium at Grindelwald 1974 — *Snow Mechanics*), 369–380.
- Vila, J. P. 1987. La prévision des vagues produites par la chute d'une avalanche dans une retenue. *International Association of Hydrological Sciences* 162 (Symposium at Davos 1986 — *Avalanche Formation, Movement and Effects*), 509–518.
- Voellmy, A. 1955. Über die Zerstörungskraft von Lawinen. *Schweiz. Bauztg.*, **73**, 159–162, 212–217, 246–249, 280–285.

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