



HOKKAIDO UNIVERSITY

Title	A semi-analytical solution for the positive degree-day model with stochastic temperature variations
Author(s)	Calov, Reinhard; Greve, Ralf
Description	Correspondence
Citation	Journal of Glaciology, 51(172), 173-175 https://doi.org/10.3189/172756505781829601
Issue Date	2005-01
Doc URL	https://hdl.handle.net/2115/34536
Rights	© 2005 International Glaciological Society
Type	journal article
File Information	Calov_Greve_2005_JGlaac.pdf



Correspondence

A semi-analytical solution for the positive degree-day model with stochastic temperature variations

The degree-day model is a parameterization for the melt rate of snow and ice at the surface of an ice sheet or glacier. It is a simple, empirical relation which states that the melt rate is proportional to the surface-air temperature excess above 0°C (e.g. Braithwaite and Olesen, 1989; Hock, 2003). The physical basis of this and related temperature-based melt-index methods was examined by Braithwaite (1995).

It was suggested by Braithwaite (1984) that the number of positive degree-days, PDD, could be calculated from the normal probability distribution around the long-term monthly mean temperatures. A form, which is based on that and is widely used in current ice-dynamic models, was proposed by Reeh (1991). It reads

$$\text{PDD} = \frac{1}{\sigma\sqrt{2\pi}} \int_0^A dt \int_0^\infty dT T \exp\left(-\frac{(T - T_{ac}(t))^2}{2\sigma^2}\right), \quad (1)$$

where t is the time, T the surface-air temperature, T_{ac} the annual temperature cycle (both in °C) and σ the standard deviation of the temperature from the annual cycle, which accounts for the daily cycle and further, stochastic temperature variations. It is often assumed that T_{ac} varies sinusoidally over time,

$$T_{ac}(t) = T_{ma} + (T_{mj} - T_{ma}) \cos \frac{2\pi t}{A}, \quad (2)$$

where T_{mj} is the mean July (January) surface-air temperature on the extratropical Northern (Southern) Hemisphere, T_{ma} is the mean annual surface-air temperature, and $A = 1$ year. Nonetheless, any other representation of T_{ac} can also be used as a basis for the calculation of the PDD integral (1).

In a number of ice-sheet models (e.g. Huybrechts and others, 1991; Van de Wal and Oerlemans, 1997; Tarasov and Peltier, 1999; Marshall and others, 2000; Charbit and others, 2002) the double integral in Equation (1) is computed numerically. Since this must be done for each gridpoint separately, it requires a considerable amount of computing time. Further, the inevitable cut-off of the upper limit of the temperature integral at some finite value, T_{max} , influences the accuracy of the results. Therefore, Janssens and Huybrechts (2000) proposed an approximation method for solving the temperature integral by fitting an exponential function.

Here, we demonstrate that the temperature integral can be evaluated fully analytically. To the best of our knowledge, this analytical solution has not been applied to the PDD temperature integral before. A similar procedure was carried out by Roe and Lindzen (2001) in the context of determining the precipitation rate over ice sheets. However, they applied an unnecessary absolute value in their equation, they did not show the integration substitutions which lead to the resulting expression and they did not evaluate the integral boundaries of the exponential function. Here, we present the complete derivation and the resulting equations for the temperature integral of the statistical positive degree-day model.

We start by adding the expression

$$\int_0^\infty dt T_{ac} \exp\left(-\frac{(T - T_{ac})^2}{2\sigma^2}\right) - \int_0^\infty dt T_{ac} \exp\left(-\frac{(T - T_{ac})^2}{2\sigma^2}\right) = 0$$

to the temperature integral in Equation (1). This yields

$$\text{PDD} = \frac{1}{\sigma\sqrt{2\pi}} \int_0^A dt \left[\int_0^\infty dT (T - T_{ac}) \exp\left(-\frac{(T - T_{ac})^2}{2\sigma^2}\right) + \int_0^\infty dT T_{ac} \exp\left(-\frac{(T - T_{ac})^2}{2\sigma^2}\right) \right]. \quad (3)$$

The substitutions $u = (T - T_{ac})^2/2\sigma^2$, $du = (T - T_{ac}) dT/\sigma^2$ and $v = (T - T_{ac})/\sqrt{2}\sigma$, $dv = dT/\sqrt{2}\sigma$ provide

$$\begin{aligned} \text{PDD} &= \frac{1}{\sigma\sqrt{2\pi}} \int_0^A dt \left[\int_{T_{ac}^2/2\sigma^2}^\infty du \sigma^2 \exp(-u) \right. \\ &\quad \left. + \int_{-T_{ac}/\sqrt{2}\sigma}^\infty dv \sqrt{2}\sigma T_{ac} \exp(-v^2) \right] \\ &= \frac{1}{\sigma\sqrt{2\pi}} \int_0^A dt \left\{ [-\sigma^2 \exp(-u)]_{T_{ac}^2/2\sigma^2}^\infty \right. \\ &\quad \left. + \sqrt{2}\sigma T_{ac} \int_{-T_{ac}/\sqrt{2}\sigma}^\infty dv \exp(-v^2) \right\} \\ &= \int_0^A dt \left[\frac{\sigma}{\sqrt{2\pi}} \exp\left(-\frac{T_{ac}^2}{2\sigma^2}\right) \right. \\ &\quad \left. + \frac{T_{ac}}{\sqrt{\pi}} \int_{-T_{ac}/\sqrt{2}\sigma}^\infty dv \exp(-v^2) \right]. \quad (4) \end{aligned}$$

With the definition of the complementary error function

$$\text{erfc}(x) = 1 - \text{erf}(x) = \frac{2}{\sqrt{\pi}} \int_x^\infty \exp(-\tilde{x}^2) d\tilde{x} \quad (5)$$

(e.g. Press and others, 1996, section 6.2; $\text{erf}(x)$ denotes the conventional error function) this can be rewritten as

$$\text{PDD} = \int_0^A dt \left[\frac{\sigma}{\sqrt{2\pi}} \exp\left(-\frac{T_{ac}^2}{2\sigma^2}\right) + \frac{T_{ac}}{2} \text{erfc}\left(-\frac{T_{ac}}{\sqrt{2}\sigma}\right) \right]. \quad (6)$$

This representation has the great advantage that the temperature integral of the original Equation (1) has vanished. Therefore, it can be numerically evaluated much faster and yields more precise results, because the upper limit ∞ need not be cut off at some finite value. The error function is not a standard built-in function of Fortran 90 or C. However, it is implemented in a number of modern Fortran and C compilers as well as in the packages MATLAB and Mathematica. Alternatively, it can easily be computed by the routines given by Press and others (1996) or similar sources. Here, we employ the subroutine `erfc` by Press and others (1996, section 6.2) with a fractional error everywhere less than 1.2×10^{-7} .

We investigate briefly the gain in accuracy and reduction in computing time to solve the new semi-analytical representation for PDD, Equation (6), compared to the fully numerical (Equation (1)) and other representations. Let us

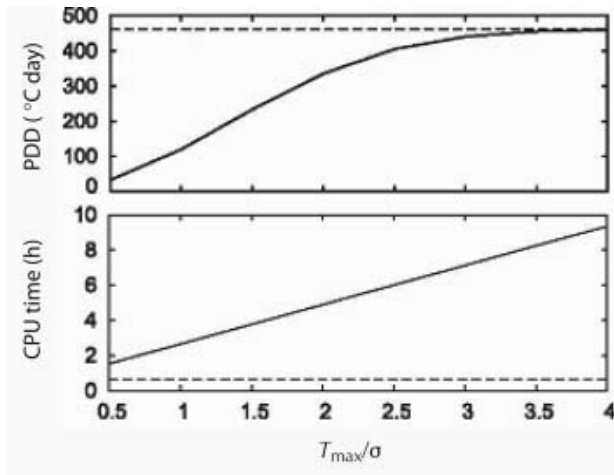


Fig. 1. PDD values and CPU times (10^9 computations) for the numerical solution of Equation (1) as a function of $T_{\max}/\sigma$ (solid lines) and Equation (6) (dashed lines). For details see main text.

consider a typical situation for the ablation zone in south Greenland with $T_{\text{ma}} = -10^\circ\text{C}$, $T_{\text{mj}} = 5^\circ\text{C}$ and $\sigma = 5^\circ\text{C}$. The time integrals in Equations (1) and (6) are solved with a monthly time-step. In addition, the temperature integral in Equation (1) is computed by the trapezoidal rule with a temperature step of $\Delta T = 0.5^\circ\text{C}$ and a cut-off temperature $T_{\max}$ varying between 0.5σ and 4σ .

The results are shown in Figure 1. CPU (computing) times refer to computations carried out with a Fortran 90 program, compiled with the Intel Fortran Compiler Version 8 and run on a 3.4 GHz Pentium-4 PC under SuSE LINUX 9.0. They are given for 10^9 computations, which is the approximate number required for a simulation of the Greenland ice sheet with 20 km resolution over 250 000 years (two glacial–interglacial cycles) with a time-step of 2.5 years. It is evident that our new method of computing the PDD integral is far more efficient than the fully numerical method. In order to be sufficiently accurate, the latter requires at least $T_{\max} = 3\sigma$, and for this cut-off temperature it consumes more than ten times more CPU time than the solution of the semi-analytical expression (6).

Table 1 gives some more detailed information about the accuracies and CPU times of the different methods. In addition to our representation of the complementary error function by the routine *erfcc* (Press and others, 1996), we have considered the rational approximation by Abramowitz (1970, section 7.1.25), here denoted as *erfcr*, which has an absolute error $\leq 2.5 \times 10^{-5}$. This approximation was also used by Roe and Lindzen (2001), but they did not give the values for the parameters. Further, we have implemented the approximate solution of the temperature integral in Equation (1) by Janssens and Huybrechts (2000) (here called *approx*). There is only a tiny difference in accuracy between the PDD values resulting from the *erfcc* and the *erfcr* representations of the complementary error function. The error due to the numerical integration over the year outweighs the differences in accuracy of the representations of the complementary error function. The accuracy of the approximation by Janssens and Huybrechts (2000) is clearly lower than that of both other methods, but is certainly tolerable in view of the relatively crude assumptions of the positive degree-day method itself (e.g. the approximation of the synoptic processes by the normal probability distribution

Table 1. Accuracy (with respect to *erfcc*) and relative CPU time (*erfcc* = 1) for different methods of solving the PDD integral. Details are given in the main text. See also Figure 1

$T_{\max}/\sigma$	$\frac{\Delta \text{PDD}}{\text{PDD}}$	CPU time
1	−74.01%	4.12
2	−27.28%	7.61
3	−4.438%	11.07
4	−0.335%	14.54
<i>erfcc</i>	(Reference)	
<i>erfcr</i>	−0.002%	0.80
<i>approx.</i>	+1.982%	1.14

(Braithwaite, 1984)). The computing time for the method with *erfcc* is slightly longer than that with *erfcr*, because of the higher-order representation of the complementary error function in the former. Interestingly, the approximation by Janssens and Huybrechts (2000) is, at least in our implementation, about 14% slower despite its lower accuracy. This is presumably because three different functions (*abs*, *exp*, *max*) must be evaluated in this approximation.

Since the complementary error function is available in nearly every compiler, or, alternatively, can be easily implemented by several tested subroutines with low computational time and high accuracy, we strongly encourage ice-sheet and glacier modellers to make use of this improved method.

Epilogue

Braithwaite’s (1984) introduction of the normal distribution in order to reduce the degree-day calculations to monthly rather than daily data reduced data requirements by a factor of 30. The methods presented in this work save more than 90% computing time. Who is going to make the next saving?

ACKNOWLEDGEMENT

We thank R.J. Braithwaite and R.C. Warner for their constructive comments on the original version of this correspondence.

Potsdam Institute for Climate Impact Research (PIK), PO Box 601203,
D-14412 Potsdam, Germany

Reinhard CALOV

Institute of Low Temperature Science,
Hokkaido University,
Sapporo 060-0819, Japan

Ralf GREVE

14 February 2005

REFERENCES

Abramowitz, M. and I.A. Stegun. 1970. *Handbook of mathematical functions with formulas, graphs, and mathematical tables*. New York, Dover Publications.
Braithwaite, R.J. 1984. Calculation of degree-days for glacier-climate research. *Z. Gletscherkd. Glazialgeol.*, **20**, 1–20.

- Braithwaite, R.J. 1995. Positive degree-day factors for ablation on the Greenland ice sheet studied by energy-balance modelling. *J. Glaciol.*, **41**(137), 153–160.
- Braithwaite, R.J. and O.B. Olesen. 1989. Calculation of glacier ablation from air temperature, West Greenland. In Oerlemans, J., ed. *Glacier fluctuations and climatic change*. Dordrecht, etc., Kluwer Academic Publishers, 219–233.
- Charbit, S., C. Ritz and G. Ramstein. 2002. Simulations of Northern Hemisphere ice-sheet retreat: sensitivity to physical mechanisms involved during the last deglaciation. *Quat. Sci. Rev.*, **21**(1–3), 243–265.
- Hock, R. 2003. Temperature index melt modelling in mountain areas. *J. Hydrol.*, **282**(1–2), 104–115.
- Huybrechts, P., A. Letréguilly and N. Reeh. 1991. The Greenland ice sheet and greenhouse warming. *Global Planet. Change*, **3**(4), 399–412.
- Janssens, I. and P. Huybrechts. 2000. The treatment of meltwater retardation in mass-balance parameterizations of the Greenland ice sheet. *Ann. Glaciol.*, **31**, 133–140.
- Marshall, S.J., L. Tarasov, G.K.C. Clarke and W.R. Peltier. 2000. Glaciological reconstruction of the Laurentide ice sheet: physical processes and modelling changes. *Can. J. Earth Sci.*, **37**(5), 769–793.
- Press, W.H., S.A. Teukolsky, W.T. Vetterling and B.P. Flannery. 1996. *Numerical recipes in FORTRAN 77: the art of scientific computing. Second edition*. Cambridge, Cambridge University Press.
- Reeh, N. 1991. Parameterization of melt rate and surface temperature on the Greenland ice sheet. *Polarforschung*, **59**(3), [1989], 113–128.
- Roe, G.H. and R.S. Lindzen. 2001. The mutual interaction between continental-scale ice sheets and atmospheric stationary waves. *J. Climate*, **14**(7), 1450–1465.
- Tarasov, L. and W.R. Peltier. 1999. The impact of thermo-mechanical ice sheet coupling on a model of the 100 kyr ice-age cycle. *J. Geophys. Res.*, **104**(D8), 9517–9545.
- Van de Wal, R.S.W. and J. Oerlemans. 1997. Modelling the short-term response of the Greenland ice sheet to global warming. *Climate Dyn.*, **13**(10), 733–744.