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**Locations and types of ruptures involved in
the 2008 Sichuan Earthquake
inferred from SAR image matching**

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Abstract

We have detected detailed ground displacements in the proximity of the Longmen Shan fault zone (LMSFZ) by applying a SAR offset-tracking method in the analysis of the 2008 Sichuan earthquake. An elevation-dependent correction is indispensable for achieving sub-meter accuracy. A sharp displacement discontinuity with a relative motion of $\sim 1\text{-}2$ m appears over a length of 200 km along the LMSFZ, which demonstrates that the main rupture has proceeded on the Beichuan fault (BF) among several active faults composing the LMSFZ, and a new active fault is detected on the northeastward extension of the BF. The rupture on the BF is characterized by a right-lateral motion in the northeast, while in the southwest an oblique right-lateral thrust slip is suggested. In contrast to the northeast, where a major rupture proceeded on the BF only, in the southwest multiple thrust ruptures have occurred in the southeastern foot of the Pengguan massif.

Keywords: 2008 Sichuan Earthquake, ground displacement, SAR offset-tracking

1. Introduction

A catastrophic earthquake with a moment magnitude of 7.9 struck China's Sichuan area on 12 May 2008 [U.S. Geological Survey, 2008]. This earthquake could be associated with the reactivation of the Longmen Shan fault zone (LMSFZ), running over a length of 200 km along the western margin of the Sichuan Basin (Figure 1). The fault zone consists of several faults dipping to the northwest that have been constructed in the area subjected to contractional stress due to the eastward movement of the Tibetan plateau [Burchfiel *et al.*, 2008]. The faults have a complicated development in and around the Pengguan massif (PM) (Figure 1). Three major active faults, indicated by red lines in Figure 1 [Densmore *et al.*, 2007], are running near the epicenter (red star) located at the southwestern edge of the PM. The rupture is thought to proceed northeastward along the fault zone as inferred from the extent of the aftershocks reported by the U. S. Geological Survey (black dots in Figure 1), but it remains uncertain where and how the faults were involved in the seismic event [Burchfiel *et al.*, 2008]. The measurement of ground displacements around the epicentral area certainly plays a key role in answering these questions.

Satellite synthetic aperture radar (SAR) data can provide detailed and spatially comprehensive ground information. Interferometric SAR (InSAR) analysis has an

advantage of detecting ground deformation in a vast region with high precision [e.g., *Massonnet and Feigl*, 1998; *Bürgmann et al.*, 2000]. However, for the Sichuan event, the standard InSAR approach is not helpful in knowing the faults directly related to the seismic rupture. This is because the displacement amplitude near the fault zone was too large and a coherent loss area which spreads over an area ~ 200 km long and ~ 10 - 30 km wide prevents us from satisfactorily obtaining the ground deformation in the proximity of the fault zone. Thus, in order to reveal the unknown surface displacements, we conduct an offset-tracking procedure that enables us to robustly detect large ground deformation even in an incoherent area [*Michel et al.*, 1999; *Tobita et al.*, 2001; *Pathier et al.*, 2006]. While similar approaches can be taken with optical images [e.g., *Avouac et al.*, 2006], SAR images are advantageous because of the radar's all-weather detection capability.

2. SAR Data Analysis

We use ALOS/PALSAR data on the ascending orbital paths 471 to 476 (Table S1) for data analysis. The data obtained is strip-map imagery with an incidence angle of 38.7° . The analyzed areas are indicated by blue frames in the inset of Figure 1. We process the SAR data from a level-1.0 product using a software package Gamma

[Wegmüller and Werner, 1997]. After conducting coregistration between two images acquired before and after the mainshock, we divide the single-look SAR amplitude images into patches and calculate an offset between the corresponding patches by an intensity tracking method. This method is performed by cross-correlating samples of backscatter intensity of a master image with those of a slave image [Strozzi *et al.*, 2002]. We now employ a near-square search patch of 64×192 pixels (range $\times$ azimuth, ~ 480 m $\times$ 610 m on the ground) and a sampling interval of 36×108 pixels (~ 270 m $\times$ 340 m on the ground). The measured offset consists of two components: (1) displacement along the line of sight (range offset) and (2) a horizontal movement on the ground, parallel along the satellite track (azimuth offset).

Accuracy of offset measurement depends strongly on the quality of image coregistration, and thus, we carry out an additional operation that is not necessary over a flat area: a correction of stereoscopic effect [Michel *et al.*, 1999; Pathier *et al.*, 2006].

Besides the actual surface displacements, range offsets include artifact offsets over rugged terrain, resulting from a difference in foreshortening effect caused by separation between the satellite orbital tracks. Thus, we reduce the artifact by applying an elevation-dependent correction incorporating SRTM (NASA's Shuttle Radar Topography Mission) DEM data with a 3 arcsec resolution [Farr *et al.*, 2007].

3. Ground displacements for the 2008 Sichuan Earthquake

Figure 2 shows the estimated range offset field over the entire analyzed area. The offset analysis succeeded in mapping the ground displacement in the proximity of the fault zone, which was difficult to achieve from the standard InSAR approach. The most notable feature is a sharp color discontinuity over a length of about 200 km, across which the ground movement is in the opposite direction (solid line). The boundary extends continuously in the northeast-southwest orientation, running just along the LMSFZ. The movement in the northern side of the boundary is away from the satellite (red), while that in the southern side is towards the satellite (blue). A major displacement field is localized within a zone no wider than several tens of kilometers, and the amount of relative displacement near the boundary reaches up to ~ 2 m. It should now be noted that the displacement boundary is in very good agreement with the Beichuan fault (BF) trace (dotted line), suggesting that a main rupture proceeded on the BF. This is the first published information from satellite data that mentions the fault as a major contributor to the earthquake rupture. Further a noteworthy accomplishment is that we detected a new active fault. A displacement boundary clearly appears not only on the BF trace (dotted line) which has been identified as an active fault [Densmore et

al., 2007] but also on its northeastward extension (Figure 2). We stress that the location of the active fault, namely the BF, is confirmed not from the geomorphological observations but from the actual surface displacements associated with the earthquake.

The original azimuth offset field (Figure S1a) suffers from periodic offset patterns that are presumably ionosphere-related noises [Gray *et al.*, 2000; Wegmüller *et al.*, 2006]. In order to reduce the oscillatory noises, we apply a band-cut filter through which the wavenumbers corresponding to the noises are removed. The detailed method is shown in the auxiliary material (Figure S2). After filtering, in the azimuth offset field (Figure S1b), displacement boundaries turn out to be in the proximity of the fault zone (solid lines). However, the spatial pattern is slightly different from the range offset. In the northeast of the fault zone, the displacement boundary overlaps the BF (dotted line) and its extension in the same location as the range offset result, but in the southwest the boundary on the BF is not clear and a new discontinuity is identified at a different location. The rupture style of this seismic event seems to be more complex than that expected from the range offset field. In the following, we focus on the paths 472 and 474, in which the azimuth noises are suppressed well by filtering, and investigate more detailed features of the ruptures in the northeast and in the southwest.

The rupture type on the BF is different between in the northeast and in the

107 southwest. Figure 3a shows the offset field in the northeast of the BF (path 472) in
108 range (left) and azimuth (right). The range offset shows that the movement in the
109 northern side of the displacement boundary (solid line) is away from the satellite (red)
110 and that in the southern side is toward the satellite (blue), while for the azimuth offset
111 the movements in the northern and southern sides are toward (red) and opposite to
112 (blue) the satellite flight direction, respectively. Taking four simple fault types into
113 account; reverse, normal, right-lateral, and left-lateral motions, the fault motion that can
114 account for the observed displacement pattern is a right-lateral slip. The displacement
115 vector field, obtained from an assumption that the observed offsets are produced by only
116 horizontal movements on the ground, shows a nearly pure right-lateral fault motion of
117 approximately 2 m (Figure S3).

118 On the other hand, in the southwest on the BF (path 474) the fault motion is
119 slightly more complex (Figure 3b). A displacement boundary on the BF that is observed
120 from both the offset fields in the northeast (Figure 3a) is not distinct in the azimuth field
121 (Figure 3b, right), although the boundary is clearly visible in the range (solid line in
122 Figure 3b, left). This suggests that the observed displacement can be explained by
123 neither a pure dextral strike-slip nor a pure reverse fault motion. The preferred rupture
124 style is a composite motion including both right-lateral and reverse slip components.

The outline of the inferred fault motions is illustrated in Figure 4. The obtained offset field reveals that the main rupture occurred on the BF and its rupture style is characterized by a right-lateral motion in the northeast and an oblique right-lateral thrust slip in the southwest. The result inferred from the offset analysis has a good agreement with that of the preliminary teleseismic waveform analysis showing that the Sichuan event slipped right-laterally in the northeast and in the southwest with a right-lateral thrust component [e.g., *Nishimura and Yagi, 2008*]. The difference in the northeast and southwest fault-motion types may suggest that the BF is segmented at the north end of the PM, which corresponds to a geological boundary [*Burchfiel et al., 2008*].

In contrast to the northeastern part of the LMSFZ where the observations can be accounted for a relatively simple motion on the BF, the offset fields in the southwest are suggestive of complicated ground deformations. In the southeastern flank of the PM (inset of Figure 3b), a clear displacement boundary with a relative motion of nearly 1 m can be observed over a length of ~50 km in the azimuth offset field (arrows in right panel). A large offset of ~2 m moving toward the satellite, featured with arrows (inset in left panel of Figure 3b), appears around the corresponding area in the range offset field. The deforming area is obviously distant from the BF (dotted line in the inset), thus indicating occurrence of another regional rupture. The detected rupture is traced near

the Pengguan fault (PF) as mapped by *Densmore et al.* [2007] but does not completely correspond to the PF trace in the azimuth offset (dotted line in Figure 3b). Thus we do not argue at present to what extent the deformation is related to the PF.

The displacement pattern of the regional deformation is evidently different from that observed along the BF. In the azimuth field, the movement in the northern side of the boundary is opposite to the satellite flight direction (blue), while that in the southern side is toward the flight direction (red). Thus a convergent motion is suggested because a left-lateral motion is unlikely. There is no clear displacement discontinuity in the range, indicating no significant strike-slip motion but probably an upheaval. Thus we can say that the local deformation is of a reverse fault type. The offset result shows that in the southwest of the LMSFZ, multiple ruptures which are different fault motions but significantly include reverse slip components occurred (Figure 4).

On the other hand, there is no signal showing any significant ground deformations along the Wenchuan-Maowen fault (WMF), which is one of the major active faults composing the LMSFZ. The idea that no significant rupture occurred on the WMF is in harmony with the observation that the aftershocks mainly distribute around the BF rather than the WMF (Figure 1).

With the knowledge of the locations of main ruptures, we constructed a

preliminary fault model that consists of three rectangular faults with distributed slips, using an elastic half-space dislocation model [Okada, 1992] (Table S2). The synthetic range offset field obtained from our preliminary fault model reproduces roughly the observed spatial pattern on the offset field (Figure S4). In our model, a right-lateral motion is dominant in the northeast of the BF (Figure S5a; F1), while in the southwest an oblique slip consisting of right-lateral and reverse motions is assumed (Figure S5b; F2). The rupture locally occurring near the PF is explained by a nearly pure thrust motion (Figure S5c; F3). However, there still remain significant discrepancies between the observed and calculated offset fields. A joint analysis with InSAR-based far-field displacements data will allow us to infer more quantitative and detailed fault model, and is deferred for future studies.

4. Uncertainties and reliabilities of measured offsets

To confirm how much errors are included in the measured offset fields, we calculate offsets using two preseismic images for the paths 472 and 474, which have almost the same perpendicular baseline (B_{perp}) as the coseismic image pairs shown in Figure 3 (Table S1). Preseismic offset fields have little displacement (Figure S6), suggesting that the observed coseismic offsets exceed the potential error level

significantly. The offset values in the basin are nearly zero; the root mean square (RMS) values are estimated to be ~ 0.1 m for both paths, but those in the mountains are not necessarily zero, particularly in high elevation areas; the RMS values are ~ 0.3 and ~ 0.4 m, for the paths 472 and 474, respectively. This may be because a correction of stereoscopic effect is insufficient because of the large amount of topographic relief. It is further noted that the displacement boundaries shown in Figure 3 are not seen in the pre-seismic data. This justifies our claim that the observed boundaries are not artifact but actual surface changes associated with the seismic event.

The SAR images that include mountainous areas suffer strongly from foreshortening effects because of steep topography, which produces severe image distortion. Figure S7 shows the offset fields for the paths (a) 472 and (b) 474 calculated using the same pre-seismic data of Figure S6 but with no elevation-dependent correction. Considerable offset errors, locally in excess of 3 m, correlating with the elevation are included. Although in principle the offset error caused by a stereoscopic effect should affect high topographic areas only, the poor global image coregistration resultantly gives more errors to the basin area as well. This result demonstrates that an elevation-dependent correction is indispensable for mountainous environments for extracting surface changes with high precision.

5. Summary

We have successfully obtained ground displacements in the proximity of the Longmen Shan fault zone (LMSFZ) by an offset-tracking method, after applying an elevation-dependent correction. The outline of the inferred fault motions is illustrated in Figure 4 and the analysis results are summarized below.

(1) A sharp displacement boundary, across which the ground movement is in the opposite direction, appears northeastward over a length of 200 km along the LMSFZ.

(2) The main rupture has proceeded on the Beichuan fault (BF) among the several active faults composing the LMSFZ.

(3) A new active fault is detected on the northeastward extension of the BF.

(4) The rupture type is characterized by a right-lateral motion in the northeast on the BF, while an oblique slip consisting of right-lateral and reverse motions is suggested in the southwest.

(5) Multiple ruptures including significant reverse slips occurred in the southeastern foot of the Pengguan massif.

215

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Figure captions

Figure 1 Tectonic setting of the Longmen Shan fault zone. Topographic relief map is taken from the SRTM DEM data with a resolution of 3 arcsecs. Light green colored areas in the relief indicate lacks of DEM data. The epicenter is marked with a star, and the aftershocks that occurred from 12-19 May are shown as black dots [U.S. Geological Survey, 2008]. Solid lines are fault traces drawn after *Densmore et al.* [2007], and the red lines stand for the Wenchuan-Maowen, Beichuan, and Pengguan faults, which are active faults that compose the majority of the Longmen Shan fault zone. Analyzed regions are indicated by blue frames in the inset.

Figure 2 Displacement field in range component. Warm and cold colors represent displacements away from and toward the satellite, respectively. A sharp color discontinuity, across which the ground motion is in the opposite direction, is highlighted with a solid line, and is recognized to be just along the Beichuan fault trace indicated by a dotted line.

Figure 3 Displacement field for the paths (a) 472 and (b) 474 in (left) range and (right) azimuth components, respectively. In the range offset field, warm and cold colors

stand for displacements away from and toward the satellite, respectively, and in the azimuth toward and opposite the satellite flight direction, respectively. Displacement boundaries are highlighted with solid lines. Arrows in the inset of the range and azimuth fields features a deforming area located at the southeastern foot of the Pengguan massif. WMF: Wenchuan-Maowen fault, BF: Beichuan fault, PF: Pengguan fault, and PM: Pengguan massif.

Figure 4 Schematic fault motions inferred from the offset fields. BF: Beichuan fault and PM: Pengguan massif.

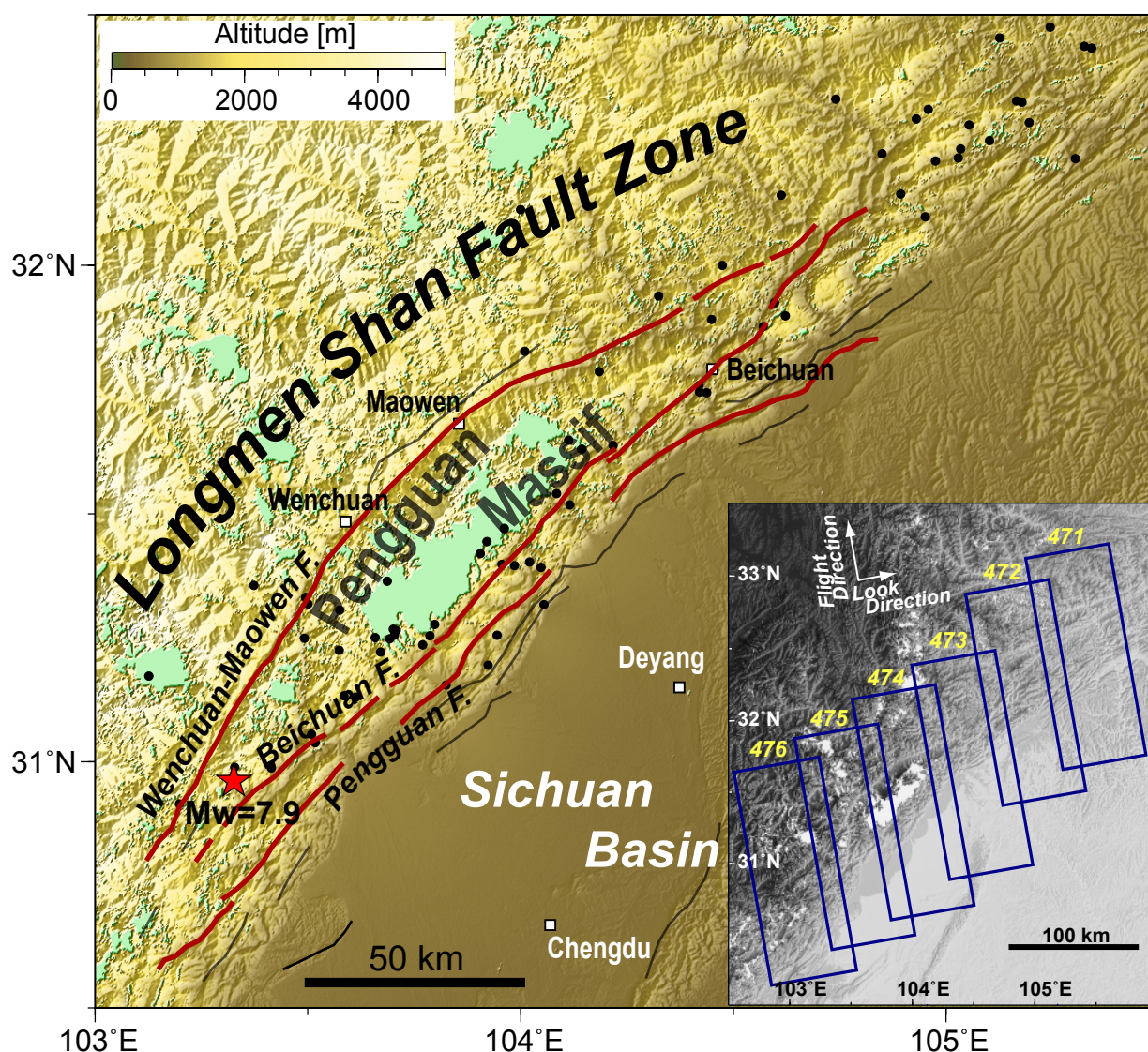


Figure 1 Tectonic setting of the Longmen Shan fault zone. Topographic relief map is taken from the SRTM DEM data with a resolution of 3 arcsecs. Light green colored areas in the relief indicate lacks of DEM data. The epicenter is marked with a star, and the aftershocks that occurred from 12-19 May are shown as black dots [U.S. Geological Survey, 2008]. Solid lines are fault traces drawn after Densmore et al. [2007], and the red lines stand for the Wenchuan-Maowen, Beichuan, and Pengguan faults, which are active faults that compose the majority of the Longmen Shan fault zone. Analyzed regions are indicated by blue frames in the inset.

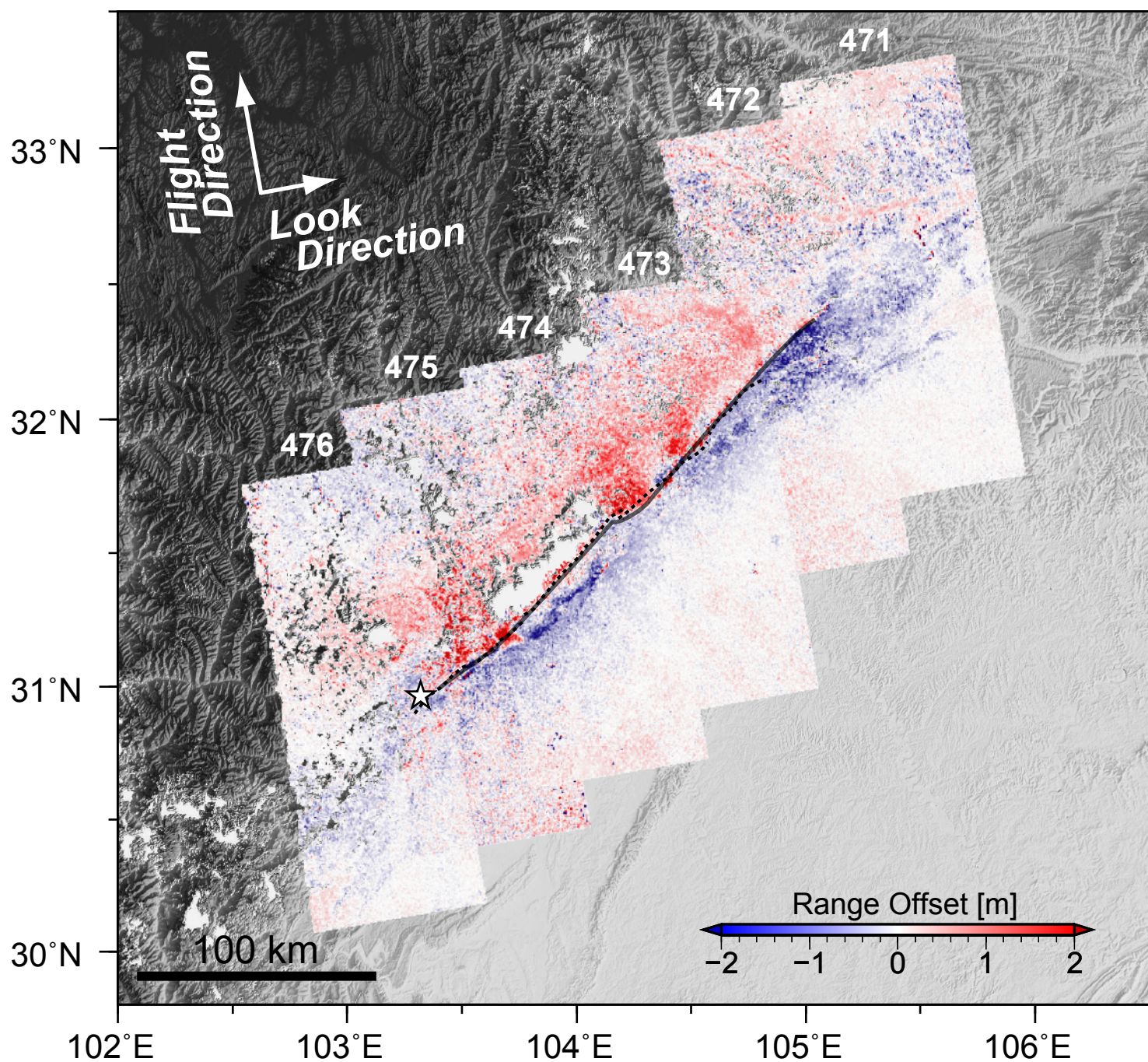


Figure 2 Displacement field in range component. Warm and cold colors represent displacements away from and toward the satellite, respectively. A sharp color discontinuity, across which the ground motion is in the opposite direction, is highlighted with a solid line, and is recognized to be just along the Beichuan fault trace indicated by a dotted line.

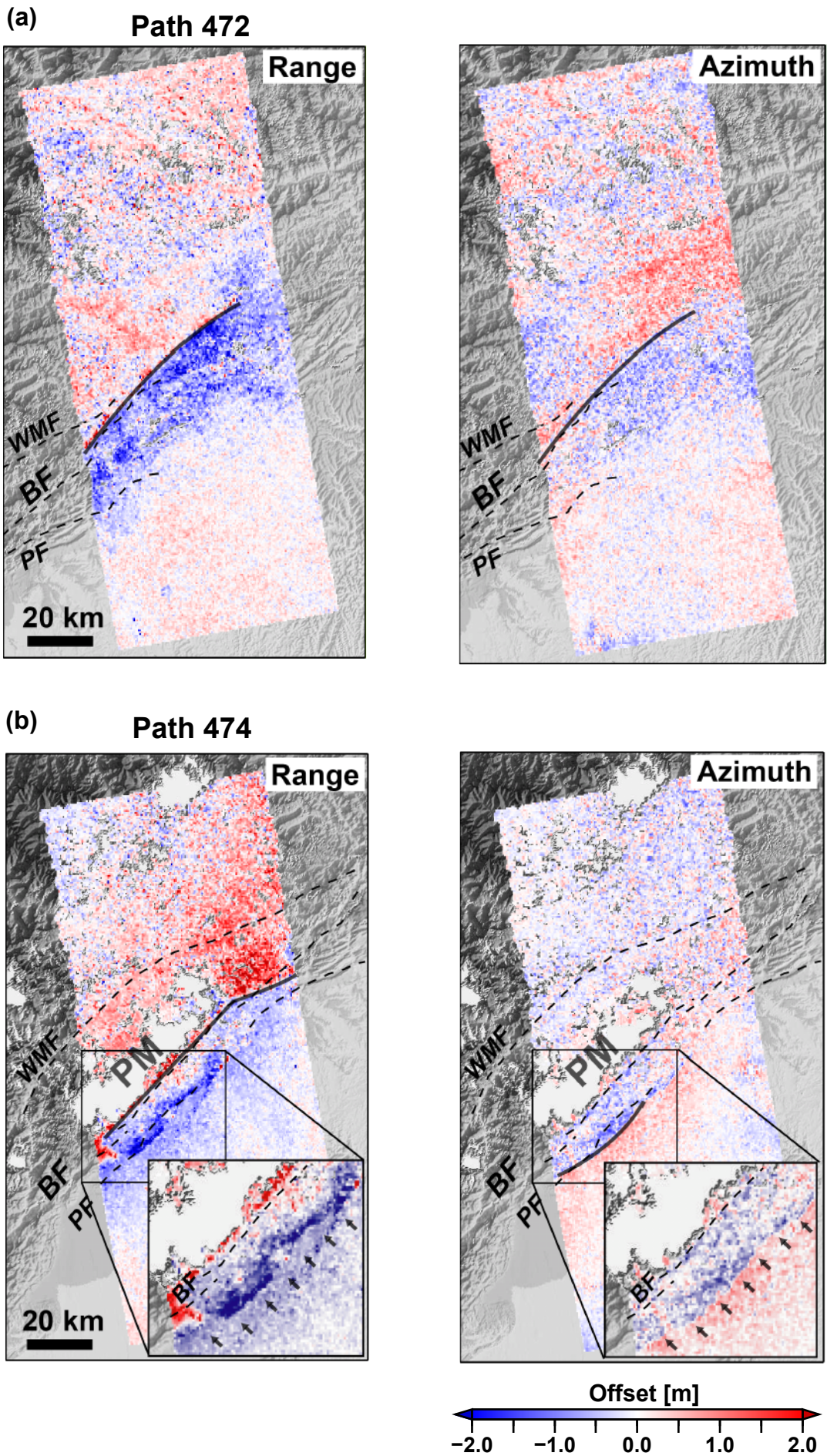


Figure 3 Displacement field for the paths (a) 472 and (b) 474 in (left) range and (right) azimuth components, respectively. In the range offset field, warm and cold colors stand for displacements away from and toward the satellite, respectively, and in the azimuth toward and opposite the satellite flight direction, respectively. Displacement boundaries are highlighted with solid lines. Arrows in the inset of the range and azimuth fields features a deforming area located at the southeastern foot of the Pengguan massif. WMF: Wenchuan-Maowen fault, BF: Beichuan fault, PF: Pengguan fault, and PM: Pengguan massif.

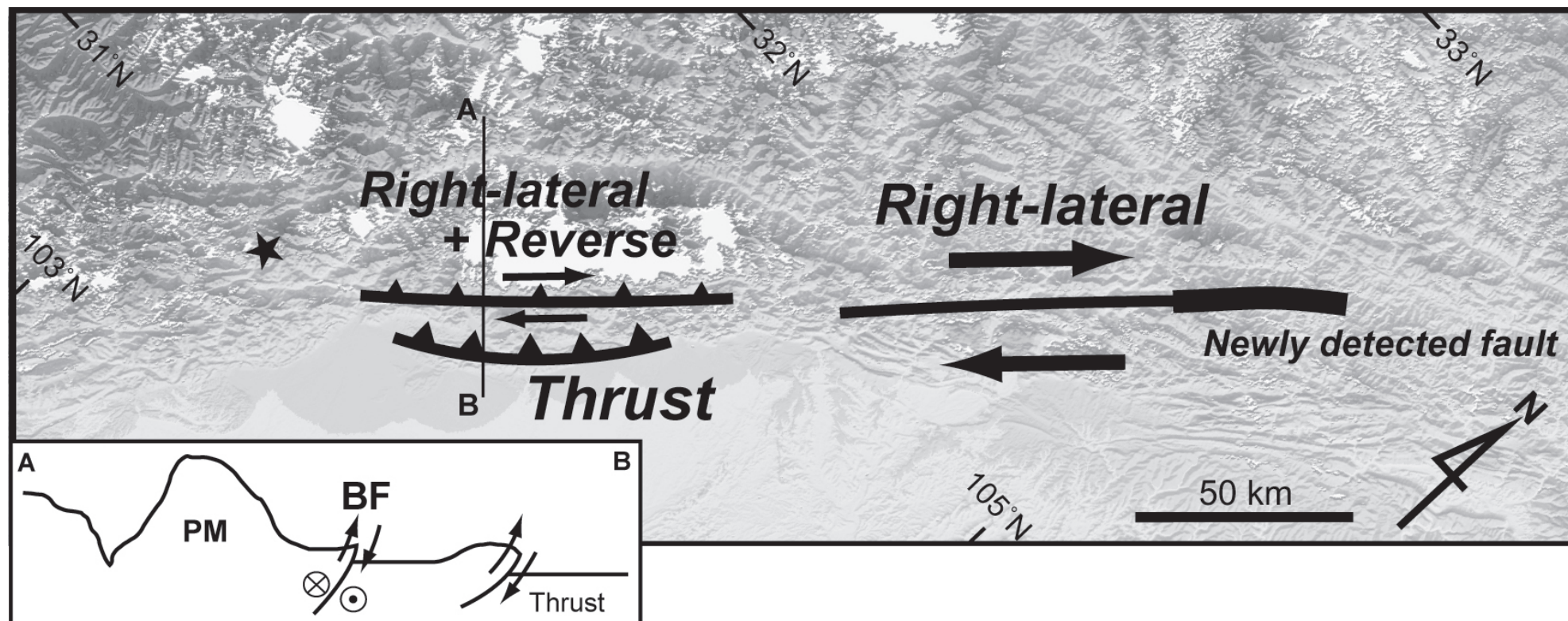


Figure 4 Schematic fault motions inferred from the offset fields. BF: Beichuan fault and PM: Pengguan massif.

Auxiliary Material Submission for Paper 2008GLXXXXXX

Locations and types of ruptures involved in the 2008 Sichuan Earthquake inferred from SAR image matching.

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Introduction

This auxiliary material contains (1) the azimuth offset fields covering the overall seismic areas, (2) a filtering method for reducing the azimuth streaks, (3) a horizontal displacement vector for the path 472, (4) a comparison between the observed and synthetic range offset fields, (5) slip distributions obtained from our preliminary fault model, (6) offset fields for preseismic data pairs for the paths 472 and 474, (7) preseismic offset fields without elevation-dependent correction for the paths 472 and 474, (8) the information list of acquired SAR data, and (9) list of parameters for the preliminary fault model. The offsets given in auxiliary materials are measured in the same manner as those shown in Figures 2 and 3.

Figure S1 is given in the JPEG file entitled “2008glXXXXXX-fs01.jpg”. The azimuth offset field suffers from periodic stripes with an amount of 4-5 m, which are presumably the ionosphere-related noise (Figure S1a) [Gray *et al.*, 2000; Wegmüller *et al.*, 2006]. In order to reduce the oscillatory noise, we applied a band-cut filter through which the

wavenumber corresponding to the noise is eliminated (Figure S1b). The filtering procedure we conducted is explained in Figure S2. Although the streaks are not completely removed in some images, displacement boundaries appear in the proximity of the fault zone (solid lines). This filtering approach does not solve the problem on the azimuth streaks substantially, but it is very useful to obtain the spatial pattern of surface changes practically.

Figure S2 is given in the JPEG file entitled “2008glXXXXXX-fs02.jpg”. We now explain the filtering procedure through which the azimuth streaks are reduced. The azimuth offset data of the path 472 is shown as an example in Figure S2. Figures S2a and S2d show the original observed azimuth offset field and its power spectrum in spatial frequency domain obtained by applying a 2D Fourier transform (FT), respectively. Strong powers are recognized at the four corners corresponding to a low-frequency domain, and among them the powers at the two corners, indicated by circles, correspond to the azimuth streaks. We cut the powers off and apply a 2D inverse Fourier transform (IFT) to the remaining frequency components (Figure S2e), and then we can obtain the azimuth offset field in which the oscillatory noise is suppressed (Figure S2b). A high-cut filter is further applied to avoid the aliasing effect and to obtain a spatially smoothed offset field by cutting high frequency noises, as the need arises. Figure S2c shows the modeled azimuth streaks that are constructed from the removed frequency components (Figure S2f). We emphasize that such a simple operation can conveniently provide us the actual ground displacement concealed under the azimuth streaks.

Figure S3 is given in the JPEG file entitled “2008glXXXXXX-fs03.jpg”. Taking into account that the fault motion for the path 472 is of strike-slip, displacement vectors are

drawn under an assumption that the observed offsets are produced only by horizontal movements on the ground. In order to stabilize the vector field, we average the offsets within a window of 30×30 pixels, which is set with a sampling interval of 30 pixels in the range and azimuth fields, respectively, and then vectorize them. The vector field is of a right-lateral fault motion, but it shows displacements that are strongly asymmetrical with much larger displacements on the southeast side of the fault than on the northwest side. Yellow arrows indicate the displacement vectors calculated using the data of the neighboring path (473). The displacement vector field including the data of path 473 shows a nearly symmetrical right-lateral motion. The displacements in the southeast side of the fault are consistent between the paths 472 and 473, but in the northwest side there is a large difference between the two. The data of path 473 was acquired one week after the earthquake occurrence, while that of path 472 was acquired one month after the earthquake (Table S1). This discrepancy in the displacement may suggest that other regional motions occurred in this mountainous area after the earthquake, which may not occur on the BF. The assumption of pure horizontal motion may not necessarily be applied to the data of path 472 in this area.

Figure S4 is given in the JPEG file entitled “2008glXXXXXX-fs04.jpg”. We construct a preliminary fault model that consists of three rectangular faults with distributed slips, using an elastic half-space dislocation model [Okada, 1992]. We separated the BF into two parts (F1 and F2) and additionally put a fault near the PF (F3). We now use the range offsets only for modeling. The fault parameters are not derived by any inversion techniques, but are manually assigned so that the calculated offset fields are largely consistent with the observations.

Figure S5 is given in the JPEG file entitled “2008glXXXXXX-fs05.jpg”.

Figure S6 is given in the JPEG file entitled “2008glXXXXXX-fs06.jpg”. The offset fields are calculated using the preseismic data pairs. The SAR scenes in Figure S6a were acquired on 16 December 2007 and 31 January 2008, and in Figure S6b on 19 January 2008 and 5 March 2008, respectively. The perpendicular baselines are almost the same as those for coseismic data pairs (Figure 3), which means that the stereoscopic effects are comparable with one another.

Figure S7 is given in the JPEG file entitled “2008glXXXXXX-fs07.jpg”. The offset fields for the paths 472 and 474 are calculated using the same preseismic data used for Figure S6 but with no elevation-dependent correction.

Table S1 is given in the text file entitled “2008glXXXXXX-ts01.txt”. The acquisition date, the observation modes handled by ALOS/PALSAR, and the perpendicular baseline are listed in Table S1. FBS and FBD indicate the fine-beam single polarization and the fine-beam dual polarization, respectively. Bperp stands for perpendicular baseline.

Table S2 is given in the text file entitled “2008glXXXXXX-ts02.txt”. Fault parameters of the preliminary model (Figures S4 and S5) are listed. The location of each fault is defined as its center. The strike is measured clockwise from the north.

1. 2008glXXXXXX-fs01.jpg (Figure S1) Azimuth offset fields for (a) original and (b) band-cut filtered images. Solid and dotted lines in Figure S1b indicate displacement boundaries across which the ground motion is in the opposite direction and a trace of the Beichuan fault [Densmore *et al.*, 2007], respectively.

2. 2008glXXXXXX-fs02.jpg (Figure S2) A filtering procedure for azimuth streaks. (a)

an original observed azimuth offset field, (b) a filtered azimuth offset field, (c) modeled azimuth streaks, and (d-f) power spectra corresponding to (a-c), respectively. Frequency components surrounded by circles correspond to the azimuth streaks.

3. 2008glXXXXXXXX-fs03.jpg (Figure S3) Horizontal displacement vectors for the area corresponding to the path 472. The vectors are calculated under an assumption that the observed offsets are produced only by horizontal movement on the ground. Red and yellow arrows represent the displacement vectors calculated using the data of paths 472 and 473, respectively. A solid line indicates the displacement boundary observed in offset fields.

4. 2008glXXXXXXXX-fs04.jpg (Figure S4) (a) Observed and (b) synthetic range offset fields. Frames indicate the fault positions projected onto the surface. Thick lines represent upper edge of each fault.

5. 2008glXXXXXXXX-fs05.jpg (Figure S5) Slip distributions for (a) F1, (b) F2, and (c) F3 shown in Figure S4, respectively. Arrows stand for the inferred slip vectors.

6. 2008glXXXXXXXX-fs06.jpg (Figure S6) Preseismic offset fields calculated by conducting an elevation-dependent correction. The SAR scenes were acquired on 16 December 2007 and 31 January 2008 for the path 472 (a) and on 19 January 2008 and 5 March 2008 for the path 474 (b).

7. 2008glXXXXXXXX-fs07.jpg (Figure S7) Same as Figure S6 but with no

elevation-dependent correction.

8. 2008glXXXXXXXX-ts01.txt (Table S1) ALOS/PALSAR images.

8.1 Column "Path", orbital track number.

8.2 Column "Date_Master", acquisition date of master image.

8.3 Column "Mode_Master", observation mode for master image.

8.4 Column "Date_Slave", acquisition date of slave image.

8.5 Column "Mode_Slave", observation mode for slave image.

8.6 Column "Bperp", m, perpendicular baseline.

9. 2008glXXXXXXXX-ts02.txt (Table S2) Fault parameters of our preliminary model shown in Figures S4 and S5.

9.1 Column "FLT", fault name.

9.2 Column "Lon", degrees, longitude of fault center.

9.3 Column "Lat", degrees, latitude of fault center.

9.4 Column "Depth", km, depth of fault center.

9.5 Column "Length", km, length of fault.

9.6 Column "Width", km, width of fault.

9.7 Column "Strike", degrees, strike direction of fault.

9.8 Column "Dip", degrees, dip angle of fault.

References:

Densmore, A. L., M. A. Ellis, Y. Li, R. Zhou, G. S. Hancock, and N. Richardson (2007),

145 Active tectonics of the Beichuan and Pengguan faults at the eastern margin of the
146 Tibetan Plateau, *Tectonics*, 26, doi:10.1029/2006TC001987.

147 Gray, A.L., D.E. Mattar, and G. Sofko (2000), Influence of ionospheric electron density
148 fluctuations on satellite radar interferometry, *Geophys. Res. Lett.*, 27, 1451-1454.

149 Okada, Y. (1992), Internal deformation due to shear and tensile faults in a half-space,
150 *Bull. Seismol. Soc. Am.*, 82, 1018-1040.

151 Wegmüller, U., C. Wener, T. Strozzi, and A. Wiesmann (2006), Ionospheric electron
152 concentration effects on SAR and INSAR, Proceeding of IGARSS 2006, Denver, 31
153 July-4 August.

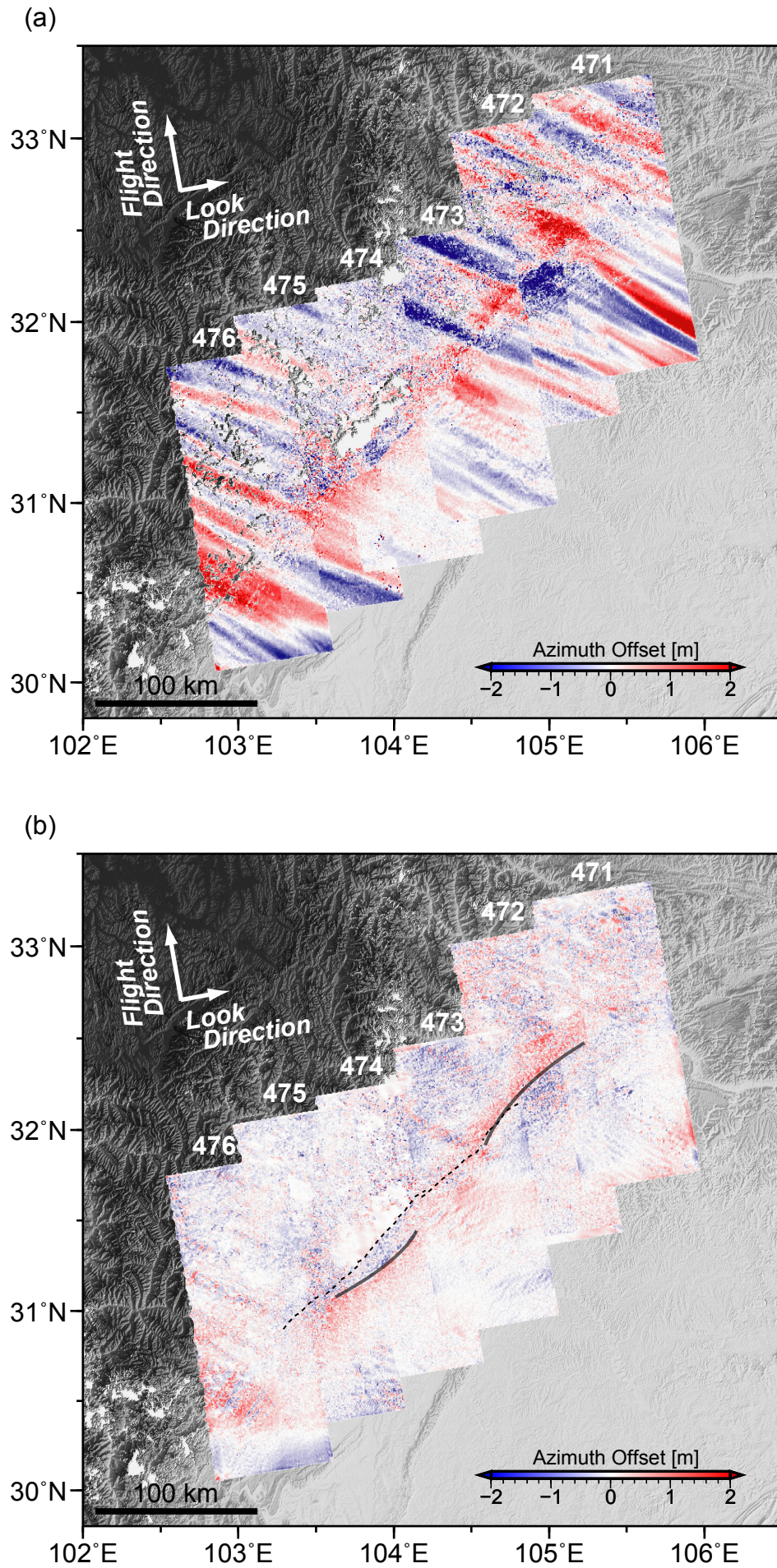


Figure S1 Azimuth offset fields for (a) original and (b) band-cut filtered images. Solid and dotted lines in Figure S1b indicate displacement boundaries across which the ground motion is in the opposite direction and a trace of the Beichuan fault [Densmore et al., 2007], respectively.

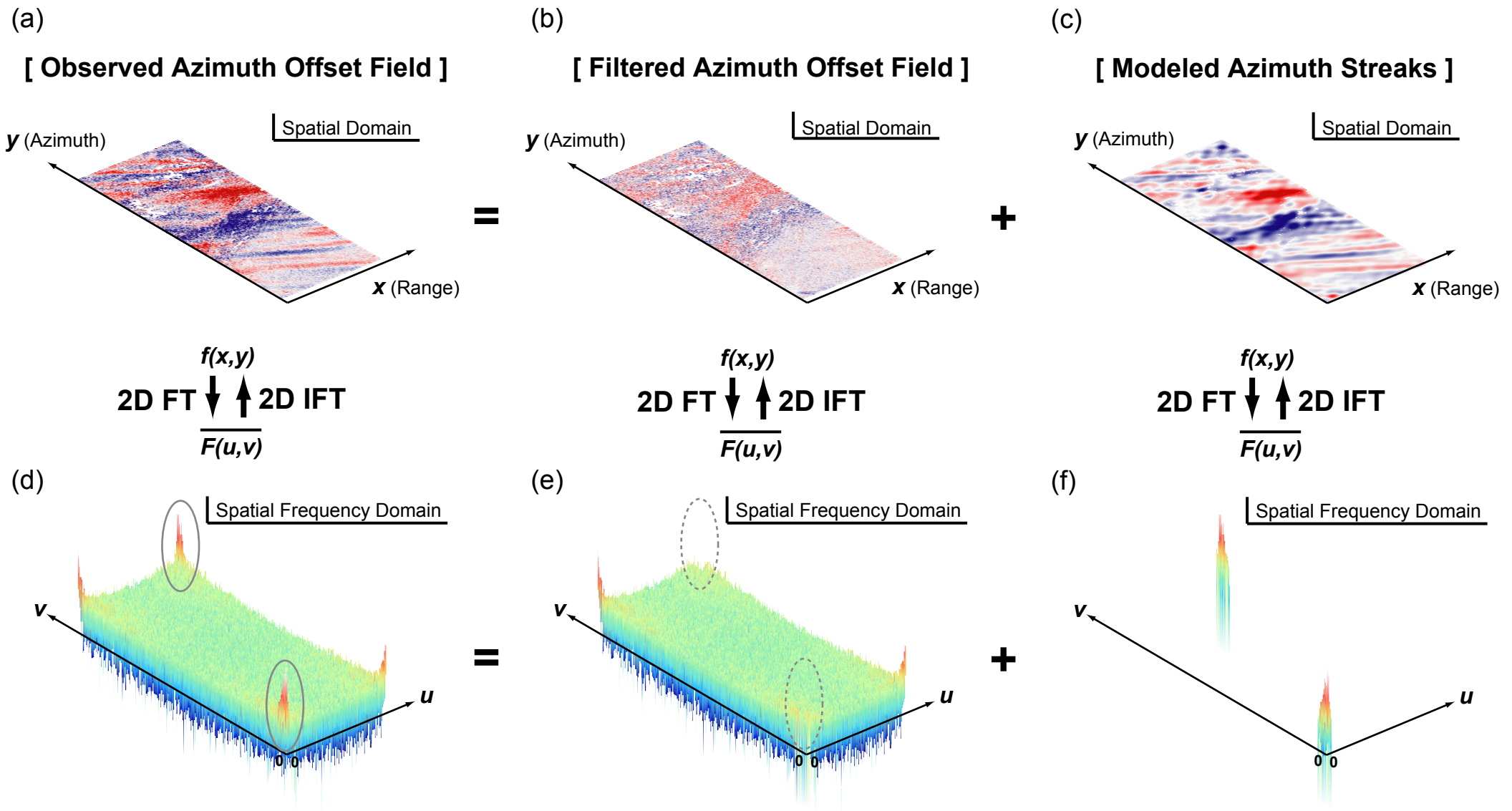


Figure S2 A filtering procedure for azimuth streaks. (a) an original observed azimuth offset field, (b) a filtered azimuth offset field, (c) modeled azimuth streaks, and (d-f) power spectra corresponding to (a-c), respectively. Frequency components surrounded by circles correspond to the azimuth streaks.

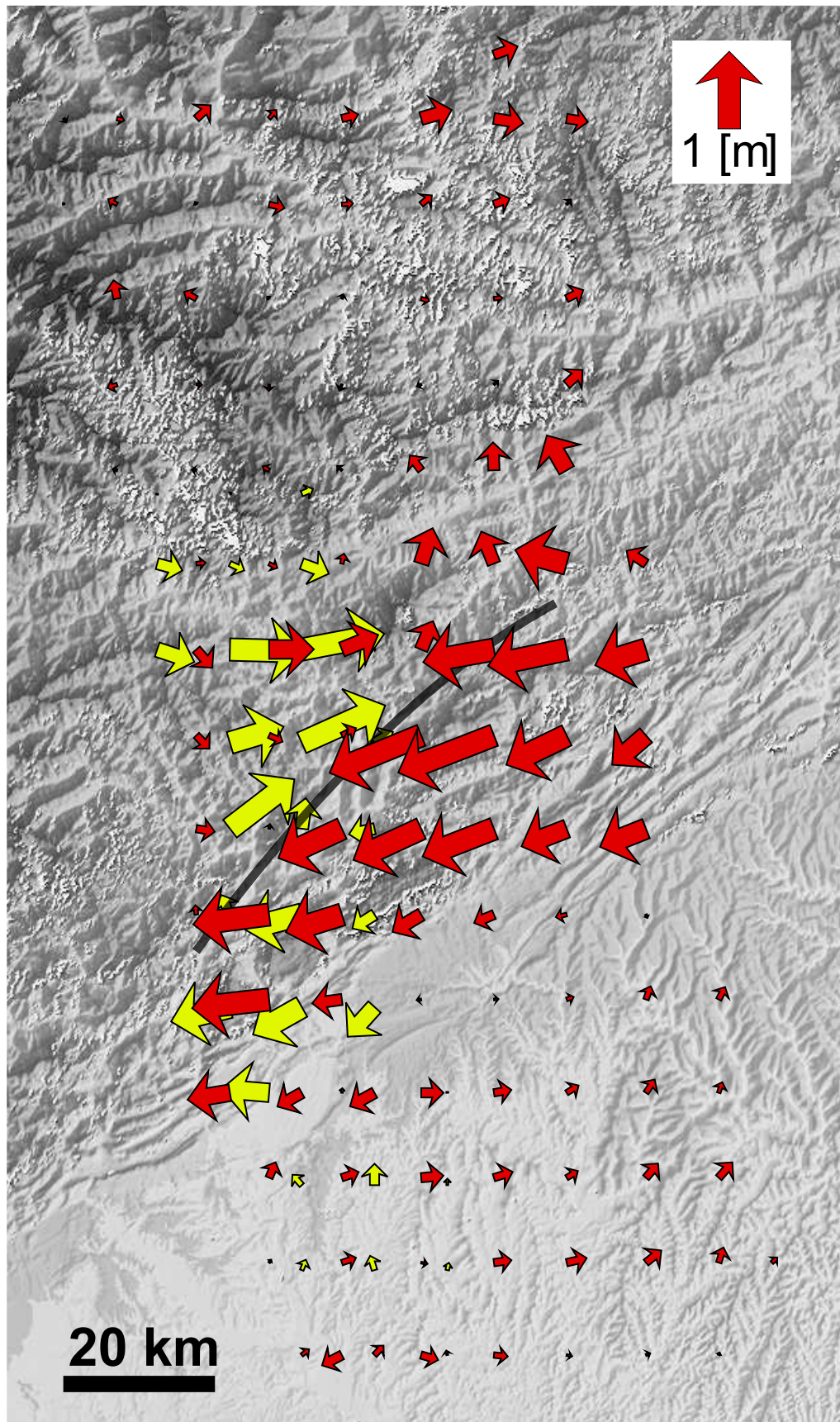


Figure S3 Horizontal displacement vectors for the area corresponding to the path 472. The vectors are calculated under an assumption that the observed offsets are produced only by horizontal movement on the ground. Red and yellow arrows represent the displacement vectors calculated using the data of paths 472 and 473, respectively. A solid line indicates the displacement boundary observed in offset fields.

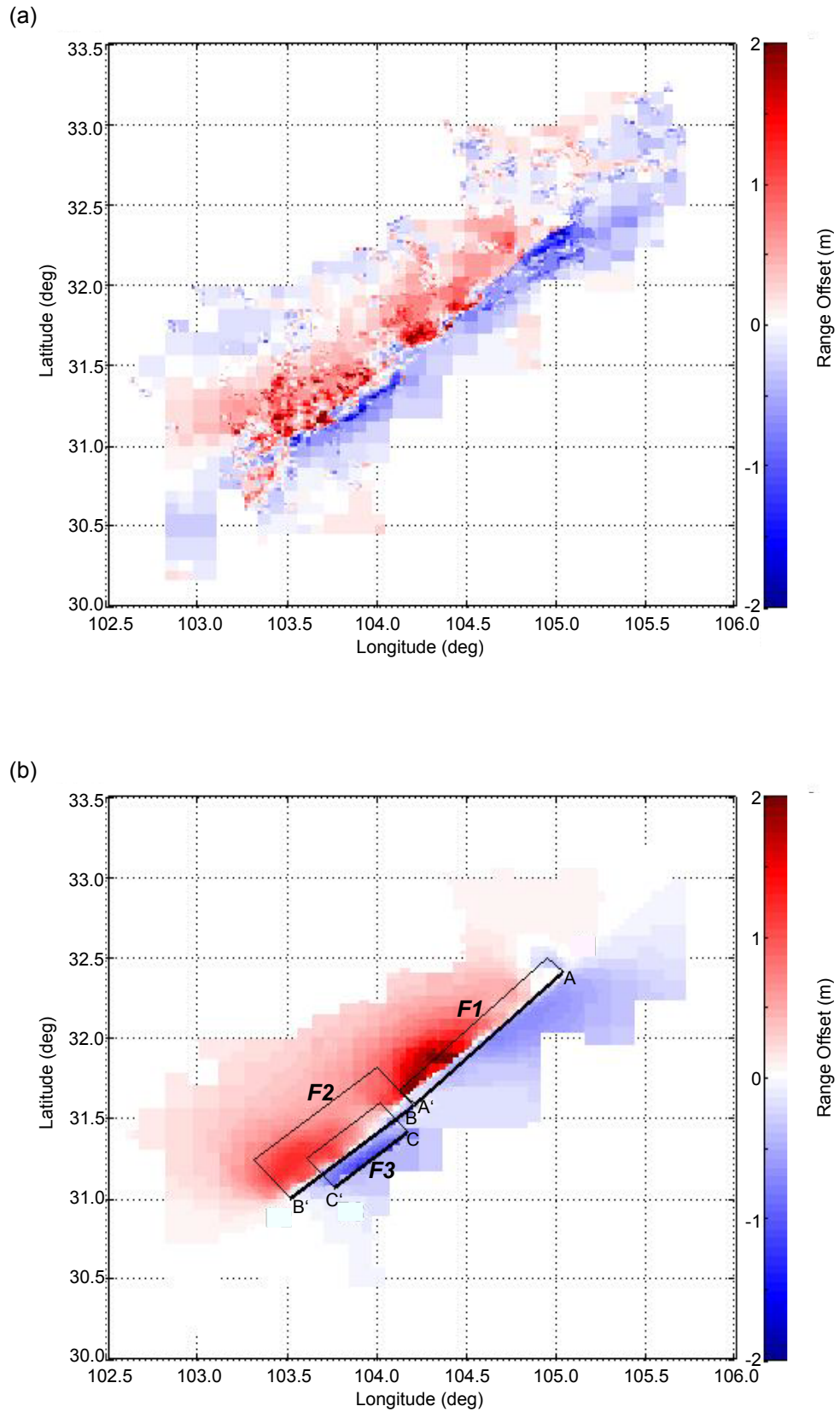


Figure S4 (a) Observed and (b) synthetic range offset fields. Frames indicate the fault positions projected onto the surface. Thick lines represent upper edge of each fault.

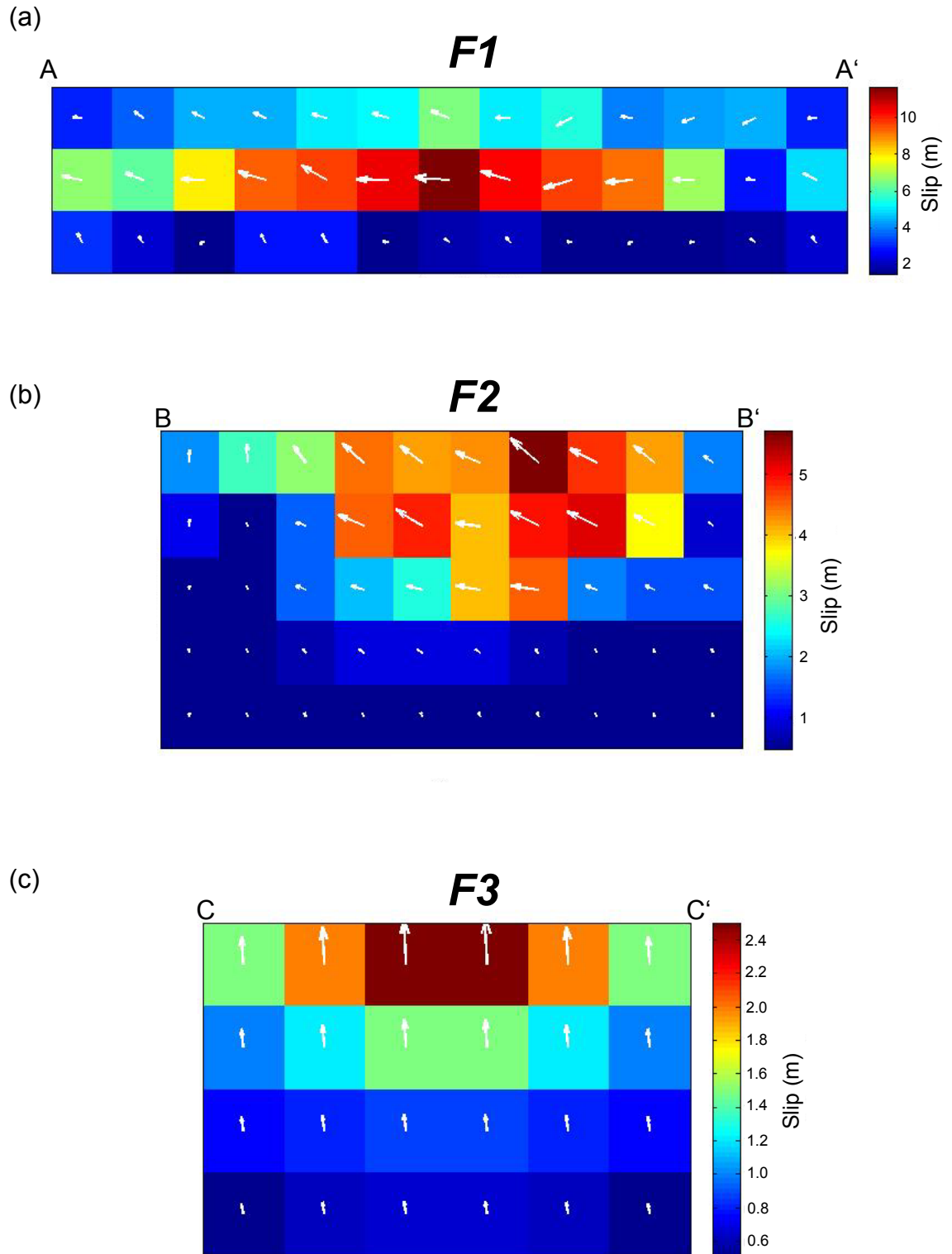
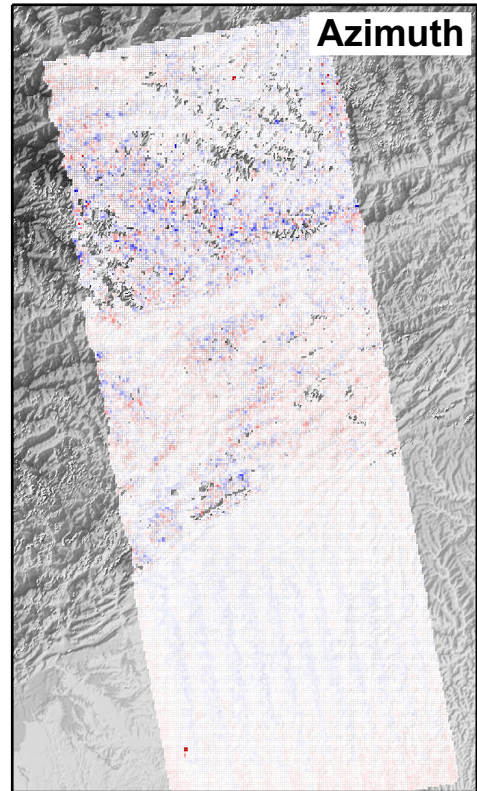
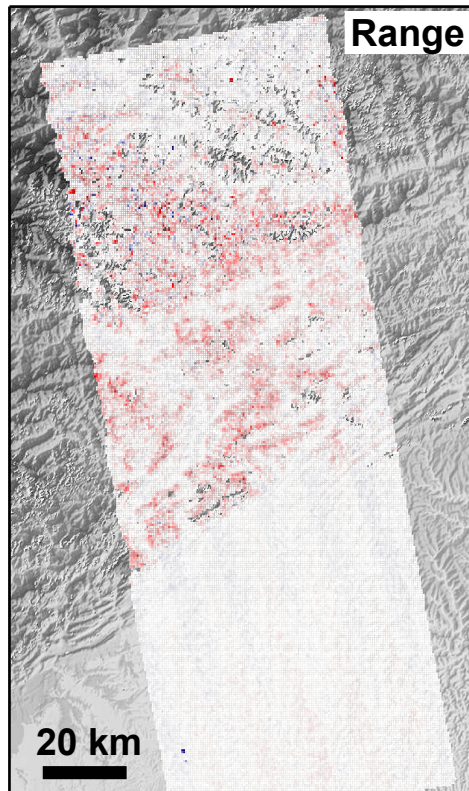


Figure S5 Slip distributions for (a) F1, (b) F2, and (c) F3 shown in Figure S4, respectively. Arrows stand for the inferred slip vectors.

(a)

Path 472



(b)

Path 474

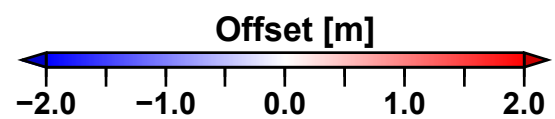
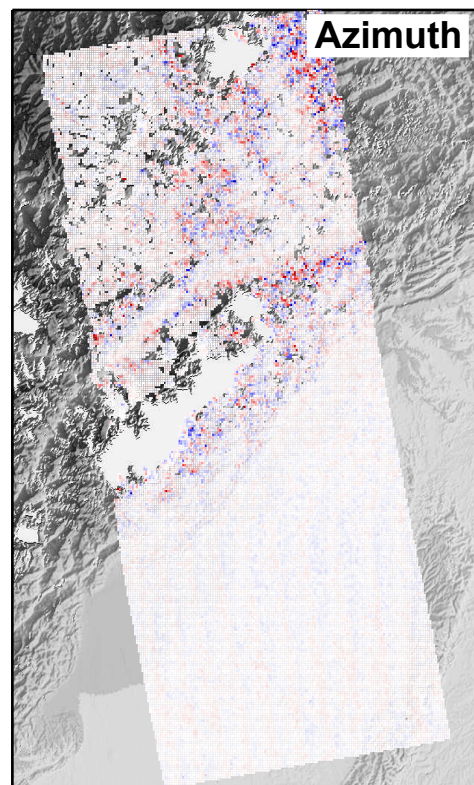
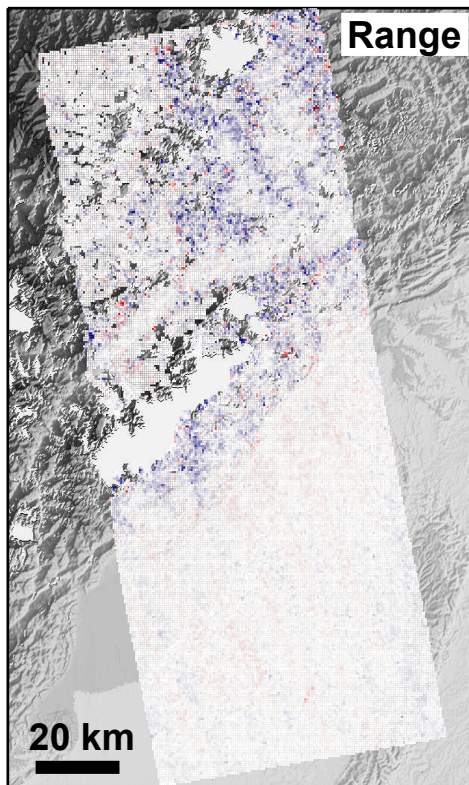
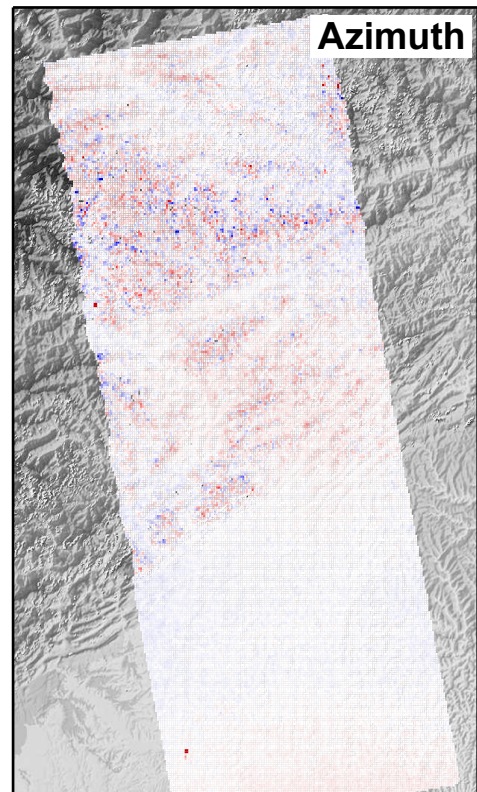
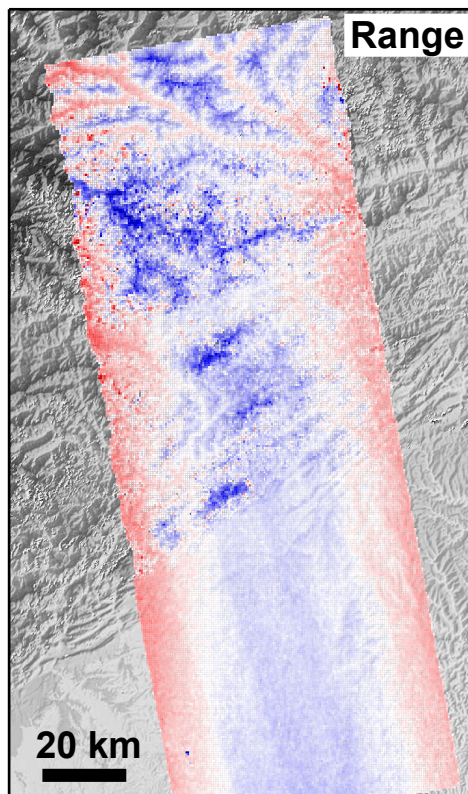


Figure S6 Preseismic offset fields calculated by conducting an elevation-dependent correction. The SAR scenes were acquired on 16 December 2007 and 31 January 2008 for the path 472 (a) and on 19 January 2008 and 5 March 2008 for the path 474 (b).

(a)

Path 472



(b)

Path 474

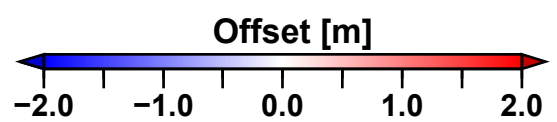
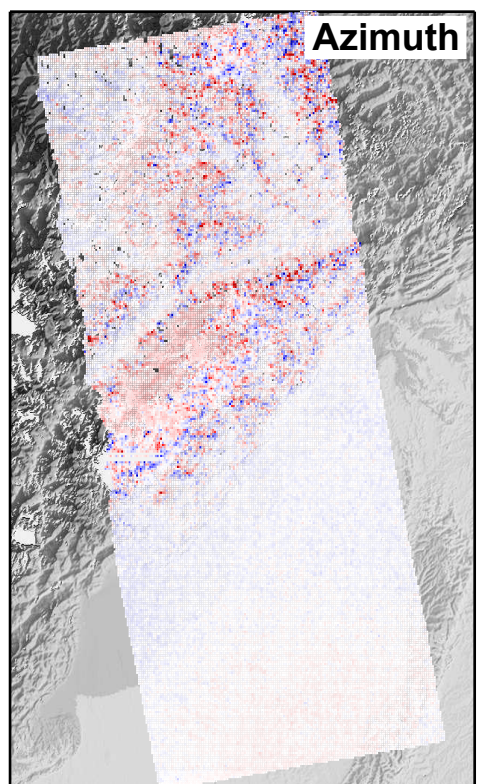
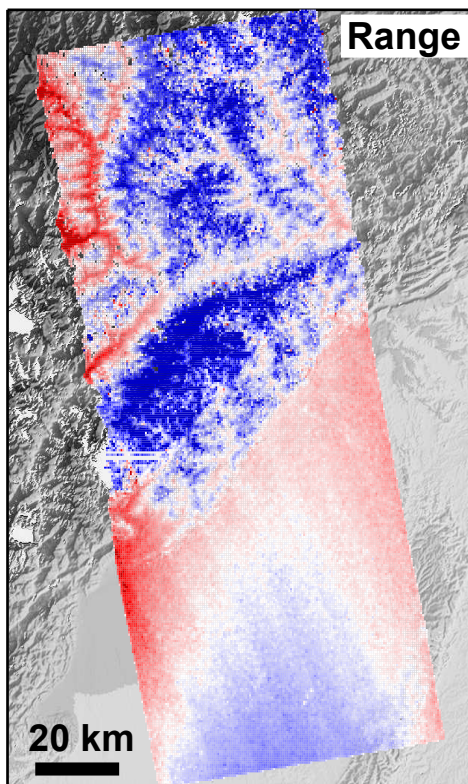


Figure S7 Same as Figure S6 but with no levation-dependent correction.