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Abstract

Daytime and nighttime fine aerosol ($\text{PM}_{2.5}$) samples were collected during a haze episode in January 2013 within the urban area of Chengdu, southwest China. Aerosol samples were analyzed for low-molecular-weight homologous dicarboxylic acids, oxocarboxylic acids and α -dicarbonyls, as well as organic carbon and elemental carbon. Concentration ranges of diacids, oxoacids, and α -dicarbonyls were $1400\text{--}5250\text{ ng m}^{-3}$, $272\text{--}1380\text{ ng m}^{-3}$, and $88\text{--}220\text{ ng m}^{-3}$, respectively. Molecular distributions of diacids (mean $3388 \pm 943\text{ ng m}^{-3}$) were characterized by a predominance of oxalic acid (C_2 ; $1373 \pm 427\text{ ng m}^{-3}$), followed by succinic (C_4), terephthalic (tPh), and phthalic (Ph) acids. Such high levels of tPh and Ph were different from those in other Asian cities where malonic acid (C_3) is the second or third highest species, mostly owing to significant emissions from coal combustion and uncontrolled waste incineration. High contents of diacids, oxoacids and α -dicarbonyls were detected on hazy days, suggesting an enhanced emission and/or formation of these organics during such a weather condition. Concentrations of unsaturated aliphatic diacids (e.g., maleic acid) and phthalic acids were higher in nighttime than in daytime. Good positive correlations of C_2 with C_3 , C_4 , ketomalonic (kC_3), pyruvic (Pyr), and glyoxylic (ωC_2) acids in daytime suggest secondary production of C_2 via the photooxidation of longer-chain diacids and ωC_2 . This study demonstrated that both primary emissions and secondary production are important sources of dicarboxylic acids and related compounds in atmospheric aerosols in the Sichuan Basin.

Keywords: dicarboxylic acids; oxocarboxylic acids; α -dicarbonyls; urban aerosols; Sichuan Basin

1. Introduction

In recent years, increasing population, industrial and agricultural development, and urbanization have resulted in serious air pollution around the world (Akimoto, 2003; Chen and Kan, 2008; Shah et al., 2013). Fine particulate matters with an aerodynamic diameter less than 2.5 μm ($\text{PM}_{2.5}$) are of particular interest in China because of the frequent severe haze events across the country (Cao et al., 2013; Ho et al., 2011; Sun et al., 2006). These fine particles can easily penetrate a human's lungs causing health problems, and also lead to adverse environmental effects such as visibility deterioration, radiative forcing, and climate change (Huebert et al., 2003; Nel, 2005; Osborne and Lambert, 2014; Wang et al., 2009c), which depend on their chemical components and characteristics (Beckerman et al., 2013; Ebel et al., 2005; Perrone et al., 2013; Schlesinger, 2007). Although the sulfate, nitrate, and inorganic carbon compounds in aerosols have received much attention, organic compounds including dicarboxylic acids (diacids) and related polar compounds are important components of atmospheric particulate matters (Alier et al., 2014; Kawamura and Sakaguchi, 1999; Nirmalkar et al., 2014).

Diacids and related polar compounds are widely present in the urban, rural, polar and marine aerosols, while their molecular distributions and variable concentrations are critical important to explore their sources and to understand related photochemical reactions (Fu et al., 2013; Ho et al., 2007; Huang et al., 2005; Kawamura and Ikushima, 1993; Kawamura et al., 2013; Kundu et al., 2010; Miyazaki et al., 2009a). Diacids are directly emitted to the atmosphere from natural and anthropogenic primary sources, and/or produced by secondary atmospheric chemical reactions (Kawamura et al., 2013; Narukawa et al., 1999; Warneck, 2003).

Studies on the molecular distributions of diacids have been conducted in urbanized regions of China such as the Yangtze River Delta (Cao et al., 2013; Wang and Kawamura, 2005; Wang et al., 2009b), Beijing-Tianjin-Hebei region, Guanzhong Plain (Cheng et al., 2013; Ho et al., 2010; Huang et al., 2005; Wang et al., 2007; Yang et al., 2011), and Pearl River Delta (Ho et al.,

2011; Li and Yu, 2010; Ma et al., 2010; Miyazaki et al., 2009b); however, little is known about their levels in the Sichuan Basin.

Chengdu, the largest megacity in southwest China, is located in the west of the Sichuan Basin. Owing to its urbanization and the rapid increase in number of vehicles and energy consumption, severe air pollution has occurred in Chengdu (Tao et al., 2011; Wang et al., 2004). The bowl-shaped basin topography with low wind speeds and stable atmospheric conditions makes Chengdu an ideal site, particularly in the winter, for investigating the molecular compositions of diacids during severe haze episodes. In the current study, we report the diurnal variability of diacids, oxocarboxylic acids (oxoacids), and α -dicarbonyls in PM_{2.5} collected in Chengdu during wintertime. Then, we compare their distributions, concentrations and compositions during different weather conditions to examine the sources and transformation mechanisms in a basin region.

2. Material and methods

2.1. Sampling site

Located in the western Sichuan Basin, southwest China, Chengdu City covers a land size of 12,121 km² with an urban area of 283.9 km², while the population of the city exceeded 11.7 million by the end of 2012 (Fig. 1; <http://www.cdstats.chengdu.gov.cn/list.asp?ClassID=020703>). The Sichuan Basin is a bowl-shaped basin covering 200,000 km² and is completely encircled by mountains that are 1500–4500 m above sea level (Li et al., 2003). The combination of a high evaporation rate and cold wind blowing from the Tibetan Plateau leads to constant production of clouds in the basin and results in overcast days for most of the year. Such clouds also prohibit the exchange of air inside the basin with the ambient atmosphere. Thus, the anthropogenic pollutants emitted into the air tend to remain within the basin (Deng et al., 2012; Lei et al., 1997). Meanwhile, haze episodes in Chengdu are enhanced by the development of an inversion layer

(<1000 m above ground level) and low wind speeds over the Sichuan Basin during the winter, accelerating the accumulation and secondary formation of pollutants emitted from local sources such as heat and power plants, home stoves, and automobiles (Lei et al., 1997). The Chinese government has promoted the auto industry in China, which has become the world's second-largest vehicle market based on unit sales, following the United States. According to Chengdu Economic and Social Development Statistics Bulletin, the number of vehicles in Chengdu exceeded 3.13 million by the end of 2012, which makes Chengdu second to Beijing for the number of vehicles in China.

Because of the basin topography, wind speed rarely exceeded that of a calm wind or breeze in Chengdu City during the sampling period, so wind direction and speed could be ignored (<http://www.weather.com.cn/weather/101270101.shtml>). Ambient temperatures ranged from 4.0–9.0°C (mean 6.3°C) in daytime, and from 1.0–8.0°C (mean 3.9°C) in nighttime. Relative humidity ranged from 51–86% (mean 72%) in daytime and from 70–90% (mean 83%) in nighttime (<http://sc.weather.com.cn/qxfw/index.shtml>).

2.2. Sample collection

Fifteen daytime (12 h, 7 a.m. to 7 p.m.) and seventeen nighttime (12 h, 7 p.m. to next day 7 a.m.) PM_{2.5} samples were collected from January 7–23, 2013 in Chengdu (30°40'40"N, 104°08'30"E), southwest China, except during the daytime of January 22 because of an external power outage. A low-volume air sampler (ME-APS100, Beijing, China) was placed on the rooftop (15 m above ground level) of a four-story building on the campus of Chengdu University of Technology, which is located in the northeastern part of Chengdu City (512 m a.s.l., see Fig. 1). The samples were collected onto a pre-baked (450°C for 12 h) quartz fiber filter (Ø47 mm) on a day/night basis with an airflow rate of 16.7 L min⁻¹. Field blanks were collected at the beginning and end of the campaign by mounting the blank filter onto the sampler for 10 min without sucking air. After sampling, all the samples were sealed within an aluminum foil bag and stored at –18°C

prior to analysis.

2.3. Analytical procedures

Filter samples were analyzed for water-soluble diacids, oxoacids, and α -dicarbonyls using the improved method of Kawamura and Ikushima (1993) and Kawamura (1993). Filter aliquots were extracted under ultrasonication with organic-free pure water (5 ml $\times$ 3). The water extracts were adjusted to pH 8.5–9.0 using a 0.05 M KOH solution and then concentrated down to ca. 0.5 ml in a pear-shaped flask by using a rotary evaporator under a vacuum and dried in a nitrogen stream (Mkoma and Kawamura, 2013). The dried samples were then reacted with 14% BF₃/*n*-butanol at 100°C for 30 min. During this procedure, carboxyl functional group was derivatized to butyl ester, and aldehyde and oxo groups were derivatized to dibutoxy acetal (Kawamura, 1993). The derivatives were extracted with 5 ml of *n*-hexane after adding 3 ml of pure water and 0.2 ml of acetonitrile, and the latter made the excess butanol more effectively transfer into the aqueous phase. The hexane layer was further washed with pure water (3 ml $\times$ 2). The extracts were dried by using a rotary evaporator and a nitrogen blow-down system, and then dissolved in 50 μ l of *n*-hexane.

The butyl esters and acetals were determined using an Agilent 6890 gas chromatograph (GC) equipped with a split/splitless injector, a fused silica capillary column (HP-5, 0.2 mm ID $\times$ 25 m $\times$ 0.52 μ m film thickness), and a flame ionization detector (FID). Peaks and GC retention times were compared with those of authentic standards to identify the compounds. Diacids and related compounds were confirmed by mass spectral analysis using a GC-MS system. Recovery experiments were performed by spiking free diacids onto precombusted quartz fiber filters. Results showed that the recoveries were better than 85% for oxalic acid and 90% for malonic, succinic, and adipic acids. Duplicate analysis of the field samples showed that the analytical error of this method was less than 10% for major species. Concentrations of diacids and related compounds reported here were corrected for the blanks but not for recoveries.

Organic carbon (OC) and elemental carbon (EC) concentrations were measured using a Sunset Lab EC/OC Analyzer following the Interagency Monitoring of Protected Visual Environments (IMPROVE) thermal evolution protocol. Typically, a 1.54 cm² punch (Ø14 mm) of the filter was placed in a quartz tube inside the thermal desorption chamber of the analyzer, and then stepwise heating was applied (Wang and Kawamura, 2005). Duplicate analyses of filter samples for OC and EC showed uncertainties of ±10%.

3. Results and discussion

3.1. Variations of OC and EC in Chengdu aerosols

EC is a product of incomplete combustion of residential coal, motor vehicle fuel, and biomass, while OC results from primary anthropogenic sources and secondary formation via chemical reactions in the atmosphere (Hegde and Kawamura, 2012). Concentration ratios of OC/EC in the Chengdu samples were 2.22±0.51 in daytime, which were slightly higher than those (2.05±0.51) in nighttime, suggesting an enhanced production of OC in daytime. Moreover, the high abundances of EC in nighttime, which may originate from the open burning of municipal wasters including plastics with large emission of terephthalic acid (a point will be discussed later), also contribute to the relative low OC/EC ratios. The OC/EC ratios were similar to those obtained in Beijing (1.9) (Jung et al., 2009b), Guangzhou (2.0, 1.8) (Ho et al., 2011; Jung et al., 2009a), Shanghai (2.3) (Ye et al., 2003), China, and Seoul, Korea (1.8) (Kim et al., 2007). Saarikoski et al. (2008) observed an OC/EC ratio of 6.6 for biomass burning and 0.71 for vehicular emissions, whereas Sandradewi et al. (2008) found values of 7.3 and 1.1 for these two sources, respectively. Cachier et al. (1989) and Watson et al. (2001) also reported that OC/EC ratio is lower for vehicle exhaust (about 1.1), and much higher for the emissions from coal combustion and biomass burning (2.7 and 9.0, respectively). Emissions from coal and biomass burning may contain more organic pollutants that should result in higher OC/EC ratios than vehicular exhausts.

OC/EC ratios in Chengdu are site specific but result from mixed emissions from the

rapidly growing number of automobiles, and burning of coal and biomass. Domestic cooking and heating in the Sichuan Basin can emit a high amount of organic pollutants. The accumulation of pollutants under steady weather conditions may enhance the formation of secondary organic aerosols, resulting in high OC/EC ratios, along with enhanced primary emission from coal and biomass burning, and vehicular exhaust.

3.2. Molecular distributions of diacids and related compounds

A homologous series of α , ω -dicarboxylic acids (C_2 – C_{11}), oxocarboxylic acids (ωC_2 – ωC_9 and pyruvic acid), and α -dicarbonyls (glyoxal and methylglyoxal) were measured in the $PM_{2.5}$ samples of Chengdu. Their concentrations together with average and median values are summarized in Table 1. Concentrations of total diacids ranged from 1410–5250 $ng\ m^{-3}$, with an average of $3450 \pm 995\ ng\ m^{-3}$ in daytime and $3330 \pm 920\ ng\ m^{-3}$ in nighttime, respectively. The average values were higher than those from the Xi'an winter haze sample (2885 $ng\ m^{-3}$) (Cheng et al., 2013), New Delhi (2875 $ng\ m^{-3}$) (Miyazaki et al., 2009a), Nanjing (1684 $ng\ m^{-3}$) (Wang et al., 2002), Hong Kong (1655 $ng\ m^{-3}$) (Li and Yu, 2010), Beijing (760 $ng\ m^{-3}$) (Ho et al., 2010), Ulaanbaatar (536 $ng\ m^{-3}$) (Jung et al., 2010), Chennai (673 $ng\ m^{-3}$) (Pavuluri et al., 2010), urban Tokyo (438 $ng\ m^{-3}$) (Kawamura and Yasui, 2005), and 3–4 times larger than those reported in 14 Chinese cities (904 $ng\ m^{-3}$) (Ho et al., 2007). Because these samples were obtained at urban sites in winter, the highest concentrations of total diacids found in Chengdu would be attributed to serious air pollution associated with specific topography of the Sichuan Basin.

Oxalic acid (C_2) was detected as the most abundant diacid species ($1370 \pm 427\ ng\ m^{-3}$), followed by succinic (C_4 , $431 \pm 147\ ng\ m^{-3}$), and terephthalic acids (tPh, $391 \pm 151\ ng\ m^{-3}$) (Table 1). On average, concentrations of C_2 – C_4 diacids accounted for 60% of total diacids. The average concentration of C_2 is lower than those at New Delhi ($1430 \pm 990\ ng\ m^{-3}$) (Miyazaki et al., 2009a) and Xi'an (winter haze time) ($1769 \pm 450\ ng\ m^{-3}$) (Cheng et al., 2013), but substantially higher than those observed at other urban sites in Asia such as Tokyo (270–1350 $ng\ m^{-3}$) (Kawamura and

Ikushima, 1993; Kawamura and Yasui, 2005), Chennai (435 ng m⁻³) (Pavuluri et al., 2010), Beijing (449 ng m⁻³) (Ho et al., 2010), Hong Kong (464 ng m⁻³) (Ho et al., 2011), and 14 Chinese cities (558 ng m⁻³) (Ho et al., 2007).

Concentrations of total oxoacids ranged from 272–1380 ng m⁻³, with an average of 750 ng m⁻³ (Table 1). Glyoxylic acid (ωC₂) (average: 430 ± 193 ng m⁻³) was found as a dominant oxoacid, followed by pyruvic acid (Pyr) and 4-oxobutanoic acid (ωC₄). The OH oxidation of glyoxal, glycolic acid, methylglyoxal, and acetic acid use the intermediate, ωC₂, and result in oxalic acid (Kroll et al., 2005). The predominance of ωC₂ in oxoacids was also observed at the Chinese urban sites (Ho et al., 2007). Oxoacids have been regarded as intermediates during the oxidation of monoacids and other precursors in the atmosphere, then resulting in diacids (Kawamura and Ikushima, 1993; Kawamura et al., 1996a).

In the present study, average concentrations of two α-dicarbonyls, glyoxal (Gly, 71.1 ng m⁻³) and methylglyoxal (MeGly, 72.1 ng m⁻³) were almost the same in Chengdu (Table 1). Their concentrations were similar with those in Xi'an (45 ng m⁻³ and 71 ng m⁻³) (Cheng et al., 2013) and Hong Kong (35.8 ng m⁻³ and 92.2 ng m⁻³) (Li and Yu, 2010), but much higher than those reported in 14 urban cities of China (3.02 ng m⁻³ and 14.5 ng m⁻³) (Ho et al., 2007), and other Asian cities such as New Delhi (18 ng m⁻³ and 25 ng m⁻³) (Miyazaki et al., 2009a), Ulaanbaatar (14 ng m⁻³ and 20 ng m⁻³) (Jung et al., 2010), Chennai (5.78 ng m⁻³ and 6.36 ng m⁻³) (Pavuluri et al., 2010), and urban Tokyo (21.5 ng m⁻³ and 24.1 ng m⁻³) (Kawamura and Yasui, 2005).

α-Dicarbonyls (Gly and MeGly) might be the gas-phase oxidation products of volatile organic compounds (VOCs) including benzene, toluene, xylene (Volkamer et al., 2001), ethylene (Ervens et al., 2004), isoprene (Zimmermann and Poppe, 1996), and terpene (Fick et al., 2004), but may act as precursors of secondary organic aerosols via heterogeneous processes (Jang et al., 2002; Kroll et al., 2005; Liggitto et al., 2005). In general, these oxoacids and carbonyls are secondarily produced via atmospheric photooxidation of various organic precursors and/or primarily formed

by fossil fuel combustion and biomass burning, and further oxidized into saturated diacids.

3.3. Temporal variations of diacids, oxoacids, and α -dicarbonyls

3.3.1. Day and night variations

Concentrations of diacids were higher in daytime than nighttime, except for unsaturated aliphatic diacids (M, F, and mM) and aromatic diacids (Ph, iPh, and tPh) (Figs. 2 and 3). The day- and night-time differences and temporal variations in diacids seem to be closely associated with the sources followed by photochemical oxidation (Hegde and Kawamura, 2012; Miyazaki et al., 2009a). Generally, higher concentrations of most diacids may be attributed to photochemical production of diacids and/or their emissions by fossil fuel combustion and cooking activities, which occurs more frequently in the daytime.

As the dominant species of diacids, oxalic acid (C_2) can be primarily generated by fossil fuel combustion and biomass burning, and secondarily formed by the oxidation of VOCs and other organic precursors in the gas phase and/or aerosol phase (Kawamura et al., 2005; Kawamura et al., 1996b). Higher concentrations and fractions of oxalic acids in daytime than nighttime might result from the differences in net production processes of C_2 between day and night: more effective production of C_2 and/or less significant loss of C_2 in daytime than in nighttime. It is possible that rates of production and loss rates of C_2 are different for daytime and nighttime. During the nighttime, oxidation of organics may be less efficient because of the absence of sunlight and lack of H_2O_2 and OH production (Buxton et al., 1997; Carlton et al., 2007; Charbouillot et al., 2012; Kawamura et al., 1996a; Singh et al., 1985; Suh et al., 2003). Positive correlations between C_2 and precursor compounds such as ωC_2 , Pyr, C_3 , and C_4 as well as kC_3 , were found especially in daytime, implying that secondary production of C_2 occurs through photochemical degradation of its precursors.

Maleic acid (M) may be produced by the ring opening of aromatic hydrocarbons such as toluene and benzene (Kawamura and Sakaguchi, 1999; Kawamura et al., 1996a). Ho et al. (2006)

reported that high concentrations of toluene observed at Hong Kong were one of the most important sources of methylmaleic acid (mM). It has been suggested that *cis* maleic acid can isomerize to *trans* isomer (fumaric acid) during photochemical transformations (Kawamura and Ikushima, 1993). Higher concentrations of maleic acid than fumaric acid may demonstrate that the Chengdu aerosols contained relatively fresh oxidation products of aromatic hydrocarbons emitted from local pollution sources, and these unsaturated aliphatic diacids were formed in an early stage of photochemical oxidation (Kawamura and Ikushima, 1993). Therefore, the higher concentrations of M, F, and mM in nighttime can be explained by the accumulation of the oxidation products of aromatic hydrocarbons, which are produced during daytime. However, we cannot exclude the possibility that aromatic hydrocarbons are oxidized at night to produce M, F, and mM.

Terephthalic and phthalic acids, which originate from the oxidation of polycyclic aromatic hydrocarbons, were found to be higher in nighttime than in daytime, similar to that of EC. Such a diurnal distribution may be attributed to different precursors and transformation processes of these aromatic diacids in Chengdu aerosols. Unlike the saturated C₂–C₄ diacids, unsaturated and branched-chain diacids are generated via photochemical oxidation of more specific precursors such as aromatic hydrocarbons (Fisseha et al., 2004; Kawamura et al., 1996a) and methylcycloalkenes (Grosjean and Fung, 1984). Singh et al. (1985) reported diurnal variations of aromatic hydrocarbons such as benzene, toluene, ethylbenzene, *o*-, *m*-, and *p*-xylenes, and 3- and 4-ethyltoluenes in twelve US cities, with maxima at night and early morning hours and minima in the late morning and early afternoon. In addition, uncontrolled burning of plastic bottles and bags and waste incineration that frequently happened at nighttime would be responsible for high concentrations of tPh during the night (Simoneit et al., 2005; Kawamura and Pavuluri, 2010), while the accumulation of primary emission and secondary product generated during the day also resulted in high Ph in nighttime. Rather, a positive correlation ($R^2=0.53$,

P<0.05) was observed between tPh and EC at nighttime, suggesting that uncontrolled burning of plastic bottles and bags is a one of the significant sources of EC in Chengdu during nighttime.

3.3.2. Enhanced formation of oxalic acid in hazy days within the Sichuan Basin

Local emissions of diacids and their precursors may be significant in Chengdu because it is a densely populated urban center with many industries including petrochemicals and thermal power plants. Several factors played an important role to form and strengthen the organic acids in Chengdu, such as the special topography of the Sichuan Basin, the height of the planetary boundary layer (PBL), and wind patterns changing infrequently during wintertime. Therefore, relatively high concentrations of total diacids, oxoacids, and α -dicarbonyls were detected in a haze episode during the winter sampling period.

Generally, the air pollution index (API) was higher in haze episodes than other times, and the temporal trend of API was somewhat similar with that of concentrations of total diacids and related compounds (Fig. 4). Aqueous phase chemistry in aerosol/cloud/fog droplets is also important in the production of diacids (Miyazaki et al., 2009a; Warneck, 2003). Since sunlight irradiation was weak and the relative humidity was quite high in hazy days, one possible explanation for the observed increases in the relative concentrations and fractions of C_2 is its secondary formation in the aqueous phase of aerosols. This is supported by high concentrations of pyruvic acid (Pyr) and glyoxylic acid (ωC_2) obtained during the haze episode (Figs. 5 and 6). C_2 is known to be formed in the aqueous phase by OH oxidation of various precursors such as longer-chain diacids, pyruvic acid (Pyr), and glyoxylic acid (ωC_2) (Carlton et al., 2007; Kawamura et al., 1996a).

It was reported that SO_4^{2-} fraction dramatically increased with the pollution level at urban sites in China (Jung et al., 2009b). Previous research observed similarity in size distributions of SO_4^{2-} and C_2 , suggesting a common source (Crahan et al., 2004; Guo et al., 2010; Huang et al., 2006; Miyazaki et al., 2010; Wang et al., 2012; Yao et al., 2003). The relationship

between C_2 and SO_4^{2-} and its linear correlation coefficient were adopted to interpret the C_2 formation processes mainly through aqueous phase reactions (Yu et al., 2005). A good correlation between C_2 and SO_4^{2-} was observed at urban cities in East Asia ($R^2 > 0.69$) (Yu et al., 2005) and in New Delhi, India ($R^2 > 0.65$) (Miyazaki et al., 2009a; Pavuluri et al., 2010), suggesting aqueous phase reactions in cloud/fog as an important pathway for the formation of secondary organic aerosols. In the present study, strong correlations between C_2 and SO_4^{2-} ($r = 0.74$ in daytime and 0.78 in nighttime; Fig. 7), ωC_2 and SO_4^{2-} ($r = 0.76$ in daytime and 0.89 in nighttime) were observed, suggesting that the photochemical production of C_2 via aqueous phase reactions under high relative humidity conditions was pronounced during the winter in Chengdu. In addition, the slope of the regression line drawn for C_2 and SO_4^{2-} in the daytime samples (14.8) is greater by 14% than that in the nighttime samples (13.1) (Fig. 7). The greater slope together with higher concentrations of C_2 in daytime suggests more effective production of C_2 in daytime.

3.4. Molecular compositions and concentration ratios: Unique features of basin aerosols

3.4.1. High concentrations of selected diacids, C_9 , Ph, tPh, and ωC_2

Azelaic acid (C_9): Among $C_5 - C_{11}$ diacids examined in urban $PM_{2.5}$ in Chengdu, azelaic acid (C_9) was much more abundant than others, indicating its unique source (Table 1 and Fig. 3). Azelaic acid was abundant in fine particles from cooking emissions in China (He et al., 2004; Zhao et al., 2007). In addition, degradation of unsaturated fatty acids (from cooking emissions) in the atmosphere can also lead to secondary formation of azelaic acid (Kawamura et al., 1996a; Robinson et al., 2006; Rogge et al., 1991). Therefore, high concentrations of azelaic acid resulted directly and indirectly from cooking emissions in Chengdu.

Terephthalic acid (tPh): Three phthalic acids including *o*-, *m*-, and *p*-isomers were detected. The isomer distribution was characterized by the predominance of terephthalic acid (tPh), followed by phthalic acid, and isophthalic acid. Interestingly, tPh acid in the Chengdu aerosols is as abundant as succinic acid (C_4) in nighttime, and terephthalic and phthalic acids are at higher

concentrations than malonic acid (C_3) in daytime (Table 1). The tPh acid was detected as the third dominant species in the Chengdu aerosols (Table 1 and Fig. 3), while such a high abundance of tPh has rarely been reported in atmospheric aerosols from other regions. The average concentration of tPh acid in Chengdu (391 ng m^{-3}) is higher than that observed in urban Xi'an (250 ng m^{-3} , tPh being the second highest species) (Cheng et al., 2013), Ulaanbaatar (130 ng m^{-3} , tPh being the dominant species) (Jung et al., 2010), about ten times higher than the 45.6 ng m^{-3} reported in wintertime aerosols of Chennai, India (Pavuluri et al., 2010), Beijing (32.4 ng m^{-3}) (Ho et al., 2010), Guangzhou (31.8 ng m^{-3}) (Ho et al., 2011), and Hong Kong (19.9 ng m^{-3}) (Ho et al., 2011). Since tPh is a major chemical component of polyester (polyethylene terephthalate) fiber and plastic bags/bottles, the burning of plastic bottles and shopping bags may emit a significant amount of tPh (Fu et al., 2010; Kawamura and Pavuluri, 2010). In addition, waste incinerators containing plastics frequently occur in open surroundings without any emission control. High concentrations of tPh obtained in Chengdu are different from those in other Asian cities, because they are emitted from uncontrolled burning of waste materials including plastic bottles and bags within the surrounding rural area.

Phthalic acid (Ph): Phthalic acid is one of the most abundant species in the Chengdu samples (Table 1 and Fig. 3), as well as in other Chinese megacities such as Xi'an (Cheng et al., 2013), Guangzhou (Ho et al., 2011), and Beijing (Ho et al., 2010). Organic aerosols in megacities of China might be related to significant emissions from coal combustion, which is not treated effectively in rural areas. Ph is directly emitted from coal combustion and/or generated by atmospheric degradation of aromatic hydrocarbons such as naphthalene (Bunce et al., 1997; Kawamura and Ikushima, 1993). In addition, high concentrations of naphthalene were observed in some cities in China (Lee et al., 2001; Liu et al., 2001; Wang et al., 2009a). The high relative abundances of Ph in total diacids were also reported in urban aerosols in Sapporo (Aggarwal and Kawamura, 2008), in aerosols from fourteen Chinese megacities (Ho et al., 2007), and in primary

auto exhaust from near a roadside and tunnel in Hong Kong (Ho et al., 2006). Ph in the Chengdu aerosols was higher than those in other cities, which can be explained by high levels of precursors including PAHs emitted from vehicular exhausts and/or coal burning in the Sichuan Basin.

Glyoxylic acid (ω C₂): Concentrations of ω C₂ in the Chengdu aerosols (Table 1 and Fig. 6b) were also among the highest in comparison to previous studies; they were approximately three to ten times higher than those from New Delhi (121 ng m⁻³) (Miyazaki et al., 2009a) and other Chinese sites (37.8 ng m⁻³) (Ho et al., 2007). C₂ is partly produced via the oxidation of ω C₂ that is derived from MeGly and Gly (Carlton et al., 2007; Cheng et al., 2013; Rinaldi et al., 2011); the levels of these α -dicarbonyls in Chengdu aerosols were also higher than those in other urban regions in China and India (Ho et al., 2007; Miyazaki et al., 2009a). Biogenic and anthropogenic VOCs can react with oxidants to produce MeGly and Gly, which can be hydrated in the aqueous phase and is further oxidized to result in ω C₂, acetic acid, hydrated ω C₂, and oxalic acid (Lim et al., 2005). Such high concentrations of ω C₂ in Chengdu aerosols indicate that ω C₂ is one of the important precursors of oxalic acid.

3.4.2. Correlation analysis and the ratios of C₃/C₄, C₆/C₉, and Ph/C₉

The correlation coefficients of selected species (C₂, C₃, C₄, kC₃, Pyr, ω C₂, and ω C₄) were examined in daytime and nighttime, respectively (Fig. 8). Other than direct vehicular emission, photochemical processes could also control the atmospheric concentrations of these species. For instance, the most abundant oxocarboxylic acid (ω C₂) can be further oxidized to C₂; a good correlation was found between ω C₂ and C₂ ($r = 0.87$, $P < 0.01$ in daytime and $r = 0.80$, $P < 0.01$ in nighttime). Malonic (C₃) and succinic (C₄) acids can be oxidized to C₂ via the breakdown of intermediates such as ketomalonic acid (kC₃) (Kawamura and Ikushima, 1993); thus, strong correlations were often observed among C₂, C₃, and C₄ in this study (Fig. 8).

The C₃/C₄ ratio can be used as an indicator of enhanced photochemical formation of diacids in the atmosphere (Kawamura and Ikushima, 1993). Malonic acid (C₃) is derived from the

incomplete combustion of fossil fuels or from the secondary atmospheric production. C_3/C_4 ratios (0.25–0.44, average 0.35) observed in vehicular exhaust were lower than those in normal atmospheric aerosol (0.56–2.9, average 1.6), due to C_3 being thermally less stable than C_4 in the high-temperature combustion process (Kawamura and Ikushima, 1993). However, the C_3/C_4 ratios in particles from secondary origin were significantly higher than that range. An average C_3/C_4 ratio above 3.0 was detected in the remote Pacific atmosphere, where these diacids mostly resulted from photochemical oxidation of aerosols in which C_3 was abundant (Kawamura and Sakaguchi, 1999). The ratios of C_3/C_4 in Chengdu $PM_{2.5}$ were on average 0.50 ± 0.13 , with a range of 0.01–0.76 (Fig. 9a), which suggests that these diacids mainly originate from the vehicular exhaust. Compared with other megacities (Table 2), the C_3/C_4 ratios in Chengdu aerosols were similar to those in Beijing (0.56) (Ho et al., 2010), Xi'an (0.59) (Cheng et al., 2013), 14 Chinese cities (0.51) (Ho et al., 2007), and New Delhi (0.59) (Miyazaki et al., 2009a), but lower than those in Chennai (1.36) (Pavuluri et al., 2010), Hong Kong (1.24) (Ho et al., 2006), and Guangzhou (0.72) (Ho et al., 2011) where secondary marine aerosols are enhanced. The C_3/C_4 ratios obtained in this study are lower than or similar to the values reported for direct emissions of vehicular exhaust (Kawamura and Ikushima, 1993), suggesting an important contribution from vehicular emissions in Chengdu.

C_6 and Ph are produced by the atmospheric oxidation of anthropogenic cyclic hexene and aromatic hydrocarbons such as naphthalene, respectively (Kawamura and Sakaguchi, 1999), whereas C_9 is derived from biogenic unsaturated fatty acids (e.g., oleic and linoleic acids) which originated from biogenic emissions (marine and terrestrial plants) and/or from anthropogenic cooking activities (Kawamura and Ikushima, 1993; Robinson et al., 2006). Previous studies reported that C_6/C_9 and Ph/ C_9 ratios could be used as potential indicators for the source strength of anthropogenic and biogenic precursors transformed to aerosol diacids (Kawamura and Yasui, 2005). The mean value of the C_6/C_9 ratio (0.66 for the day and night samples; Table 2; Fig. 9b) in

our study is about 2–3 times lower than that for biomass burning aerosols. However, the C_6/C_9 ratio is similar to that for urban Chinese and Tokyo aerosols (Ho et al., 2007; Kawamura and Yasui, 2005), suggesting that adipic (C_6) acid is less commonly derived from biomass burning and production of C_9 through the oxidation of biogenic unsaturated fatty acids and is possibly higher due to domestic cooking in Chengdu.

We obtained an average Ph/ C_9 ratio of 2.36 (Table 2 and Fig. 9c) that is comparable to those (2.23) from urban areas of Chinese megacities (Ho et al., 2010), but higher than those in urban areas of New Delhi and Chennai, India (Miyazaki et al., 2009a; Pavuluri et al., 2010), Ulaanbaatar, Mongolia (Jung et al., 2010), and Tokyo, Japan (Kawamura and Yasui, 2005). This comparison implies that the contribution of Ph from other anthropogenic sources is more significant than from biomass burning. Kawamura and Kaplan (1987) reported that diesel fuel vehicular exhaust showed a Ph/ C_6 ratio of 6.58, which was higher than that from gasoline fuel vehicles (2.05). An averaged Ph/ C_6 ratio of 3.70 ± 0.9 obtained in this study (Table 2 and Fig. 9d), is higher than those from the urban areas in India (1.38 in New Delhi, 2.92 in Chennai) (Miyazaki et al., 2009a; Pavuluri et al., 2010) and in Japan (1.68 in Tokyo) (Kawamura and Yasui, 2005). Because most passenger cars in Chengdu are gasoline fuel vehicles, this study suggested that naphthalenes and other PAHs from coal burning may be the important source of Ph in Chengdu.

4. Conclusions

Homologous dicarboxylic acids, oxoacids and α -dicarbonyls (glyoxal and methylglyoxal), as well as OC and EC were detected in aerosol samples collected during winter 2012 at an urban site in Chengdu, Sichuan Basin in southwest China. Average concentrations of oxalic acid (C_2 , 1373 ng m^{-3}), terephthalic acid (tPh, 391 ng m^{-3}), glyoxylic acid (ωC_2 , 430 ng m^{-3}), and total diacids (3388 ng m^{-3}) were significantly higher than was reported at other urban sites in Asia.

C₂ was the most abundant dicarboxylic acid, followed by succinic, terephthalic, and phthalic acids. The average concentrations and mass fractions of C₂ were higher in daytime (1439 ng m⁻³ and 41.7% of total diacids) than in nighttime (1315 ng m⁻³ and 38.9% of total diacids). Positive correlations of C₂ with C₃, C₄, and ωC₂ were better in daytime than in nighttime, suggesting that significant secondary production of C₂ occurred in daytime via the oxidation of both longer-chain diacids and ωC₂ in the Sichuan Basin during wintertime. An aqueous phase production of C₂ might be important in hazy days under the condition of the basin topography of Chengdu. Relatively low C₃/C₄ ratios (0.50 ± 0.13) and high concentrations of tPh and Ph were found in this study, suggesting that primary emissions from vehicular exhaust, coal combustions and waste incinerations are important emission sources for polar organic compounds in the Sichuan Basin.

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References

Aggarwal SG, Kawamura K (2008) Molecular distributions and stable carbon isotopic compositions of dicarboxylic acids and related compounds in aerosols from Sapporo, Japan: Implications for photochemical aging during long - range atmospheric transport. J Geophys Res 113, D14301

447 Akimoto H (2003) Global air quality and pollution. *Science* 302, 1716-1719

448 Alier M, Osto MD, Lin Y-H, Surratt JD, Tauler R, Grimalt JO, van Drooge BL (2014) On the origin of
449 water-soluble organic tracer compounds in fine aerosols in two cities: the case of Los Angeles and
450 Barcelona. *Environ Sci Pollut Res* 21, 11649-11660

451 Beckerman BS, Jerrett M, Serre M, Martin RV, Lee S-J, van Donkelaar A, Ross Z, Su J, Burnett RT (2013)
452 A hybrid approach to estimating national scale spatiotemporal variability of PM_{2.5} in the contiguous
453 United States. *Environ Sci Technol* 47, 7233-7241

454 Bunce NJ, Liu L, Zhu J, Lane DA (1997) Reaction of naphthalene and its derivatives with hydroxyl
455 radicals in the gas phase. *Environ Sci Technol* 31, 2252-2259

456 Buxton GV, Malone TN, Salmon GA (1997) Oxidation of glyoxal initiated by OH in oxygenated aqueous
457 solution. *J Chem Soc, Faraday Trans* 93, 2889-2891

458 Cachier H, Brémond M-P, Buat-Ménard P (1989) Carbonaceous aerosols from different tropical biomass
459 burning sources. *Nature* 340, 371-373

460 Cao JJ, Zhu CS, Tie XX, Geng FH, Xu HM, Ho S, Wang GH, Han YM, Ho KF (2013) Characteristics and
461 sources of carbonaceous aerosols from Shanghai, China. *Atmos Chem Phys* 13, 803-817

462 Carlton AG, Turpin BJ, Altieri KE, Seitzinger S, Reff A, Lim H-J, Ervens B (2007) Atmospheric oxalic
463 acid and SOA production from glyoxal: Results of aqueous photooxidation experiments. *Atmos Environ* 41,
464 7588-7602

465 Charbouillot T, Gorini S, Vyard G, Parazols M, Brigante M, Deguillaume L, Delort A-M, Mailhot G (2012)
466 Mechanism of carboxylic acid photooxidation in atmospheric aqueous phase: Formation, fate and reactivity.
467 *Atmos Environ* 56, 1-8

468 Chen B, Kan H (2008) Air pollution and population health: a global challenge. *Environ Health Prev Med*
469 13, 94-101

470 Cheng C, Wang G, Zhou B, Meng J, Li J, Cao J, Xiao S (2013) Comparison of dicarboxylic acids and
471 related compounds in aerosol samples collected in Xi'an, China during haze and clean periods. *Atmos*
472 *Environ* 81, 443-449

473 Crahan KK, Hegg D, Covert DS, Jonsson H (2004) An exploration of aqueous oxalic acid production in the
474 coastal marine atmosphere. *Atmos Environ* 38, 3757-3764

475 Deng LQ, Qian J, Liao RX, Tong HJ (2012) Pollution characteristics of atmospheric particulates in
476 chengdu from August to September in 2009 and their relationship with meteorological conditions. *China*
477 *Environ Sci* 32, 1433-1438

478 Ebelt ST, Wilson WE, Brauer M (2005) Exposure to ambient and nonambient components of particulate
479 matter: a comparison of health effects. *Epidemiology* 16, 396-405

480 Ervens B, Feingold G, Frost GJ, Kreidenweis SM (2004) A modeling study of aqueous production of
481 dicarboxylic acids: 1. Chemical pathways and speciated organic mass production. *J Geophys Res* 109,

482 D15205

483 Fick J, Nilsson C, Andersson B (2004) Formation of oxidation products in a ventilation system. *Atmos*
 484 *Environ* 38, 5895-5899

485 Fisseha R, Dommen J, Sax M, Paulsen D, Kalberer M, Maurer R, Höfler F, Weingartner E, Baltensperger U
 486 (2004) Identification of organic acids in secondary organic aerosol and the corresponding gas phase from
 487 chamber experiments. *Anal Chem* 76, 6535-6540

488 Fu PQ, Kawamura K, Pavuluri CM, Swaminathan T, Chen J (2010) Molecular characterization of urban
 489 organic aerosol in tropical India: contributions of primary emissions and secondary photooxidation. *Atmos*
 490 *Chem Phys* 10, 2663-2689

491 Fu PQ, Kawamura K, Usukura K, Miura K (2013) Dicarboxylic acids, ketocarboxylic acids and glyoxal in
 492 the marine aerosols collected during a round-the-world cruise. *Mar Chem* 148, 22-32

493 Grosjean D, Fung K (1984) Hydrocarbons and carbonyls in Los Angeles air. *J Air Pollut Control Assoc* 34,
 494 537-543

495 Guo S, Hu M, Wang ZB, Slanina J, Zhao YL (2010) Size-resolved aerosol water-soluble ionic
 496 compositions in the summer of Beijing: implication of regional secondary formation. *Atmos Chem Phys* 10,
 497 947-959

498 He LY, Hu M, Huang XF, Yu BD, Zhang YH, Liu DQ (2004) Measurement of emissions of fine particulate
 499 organic matter from Chinese cooking. *Atmos Environ* 38, 6557-6564

500 Hegde P, Kawamura K (2012) Seasonal variations of water-soluble organic carbon, dicarboxylic acids,
 501 ketocarboxylic acids, and α -dicarbonyls in Central Himalayan aerosols. *Atmos Chem Phys* 12, 6645-6665

502 Ho KF, Lee SC, Cao JJ, Kawamura K, Watanabe T, Cheng Y, Chow JC (2006) Dicarboxylic acids,
 503 ketocarboxylic acids and dicarbonyls in the urban roadside area of Hong Kong. *Atmos Environ* 40,
 504 3030-3040

505 Ho KF, Cao JJ, Lee SC, Kawamura K, Zhang RJ, Chow JC, Watson JG (2007) Dicarboxylic acids,
 506 ketocarboxylic acids, and dicarbonyls in the urban atmosphere of China. *J Geophys Res* 112, D22S27

507 Ho KF, Lee SC, Ho SSH, Kawamura K, Tachibana E, Cheng Y, Zhu T (2010) Dicarboxylic acids,
 508 ketocarboxylic acids, α - dicarbonyls, fatty acids, and benzoic acid in urban aerosols collected during the
 509 2006 Campaign of Air Quality Research in Beijing (CAREBeijing - 2006). *J Geophys Res* 115, D19312

510 Ho KF, Ho S, Lee SC, Kawamura K, Zou SC, Cao JJ, Xu HM (2011) Summer and winter variations of
 511 dicarboxylic acids, fatty acids and benzoic acid in PM 2.5 in Pearl Delta River Region, China. *Atmos Chem*
 512 *Phys* 11, 2197-2208

513 Huang X, Hu M, He L, Tang X (2005) Chemical characterization of water-soluble organic acids in PM2.5
 514 in Beijing, China. *Atmos Environ* 39, 2819-2827

515 Huang XF, Yu JZ, He LY, Yuan ZB (2006) Water - soluble organic carbon and oxalate in aerosols at a
 516 coastal urban site in China: Size distribution characteristics, sources, and formation mechanisms. *J*

517 Geophys Res 111, D22212

518 Huebert BJ, Bates T, Russell PB, Shi G, Kim YJ, Kawamura K, Carmichael G, Nakajima T (2003) An
 519 overview of ACE - Asia: Strategies for quantifying the relationships between Asian aerosols and their
 520 climatic impacts. J Geophys Res 108, 8633

521 Jang M, Czoschke NM, Lee S, Kamens RM (2002) Heterogeneous atmospheric aerosol production by
 522 acid-catalyzed particle-phase reactions. Science 298, 814-817

523 Jung J, Lee H, Kim YJ, Liu XG, Zhang YH, Gu JW, Fan SJ (2009a) Aerosol chemistry and the effect of
 524 aerosol water content on visibility impairment and radiative forcing in Guangzhou during the 2006 Pearl
 525 River Delta campaign. J Environ Manage 90, 3231-3244

526 Jung J, Lee H, Kim YJ, Liu XG, Zhang YH, Hu M, Sugimoto N (2009b) Optical properties of atmospheric
 527 aerosols obtained by in situ and remote measurements during 2006 Campaign of Air Quality Research in
 528 Beijing (CAREBeijing - 2006). J Geophys Res 114, D00G02

529 Jung JS, Tsatsral B, Kim YJ, Kawamura K (2010) Organic and inorganic aerosol compositions in
 530 Ulaanbaatar, Mongolia, during the cold winter of 2007 to 2008: Dicarboxylic acids, ketocarboxylic acids,
 531 and α - dicarbonyls. J Geophys Res 115, D22203

532 Kawamura K, Kaplan IR (1987) Motor exhaust emissions as a primary source for dicarboxylic acids in Los
 533 Angeles ambient air. Environ Sci Technol 21, 105-110

534 Kawamura K (1993) Identification of C₂-C₁₀ ω -oxocarboxylic acids, pyruvic acid, and C₂-C₃ α -dicarbonyls
 535 in wet precipitation and aerosol samples by capillary GC and GC/MS. Anal Chem 65, 3505-3511

536 Kawamura K, Ikushima K (1993) Seasonal changes in the distribution of dicarboxylic acids in the urban
 537 atmosphere. Environ Sci Technol 27, 2227-2235

538 Kawamura K, Kasukabe H, Barrie LA (1996a) Source and reaction pathways of dicarboxylic acids,
 539 ketoacids and dicarbonyls in arctic aerosols: One year of observations. Atmos Environ 30, 1709-1722

540 Kawamura K, Steinberg S, Kaplan IR (1996b) Concentrations of monocarboxylic and dicarboxylic acids
 541 and aldehydes in southern California wet precipitations: Comparison of urban and nonurban samples and
 542 compositional changes during scavenging. Atmos Environ 30, 1035-1052

543 Kawamura K, Sakaguchi F (1999) Molecular distributions of water soluble dicarboxylic acids in marine
 544 aerosols over the Pacific Ocean including tropics. J Geophys Res 104, 3501-3509

545 Kawamura K, Imai Y, Barrie LA (2005) Photochemical production and loss of organic acids in high Arctic
 546 aerosols during long-range transport and polar sunrise ozone depletion events. Atmos Environ 39, 599-614

547 Kawamura K, Yasui O (2005) Diurnal changes in the distribution of dicarboxylic acids, ketocarboxylic
 548 acids and dicarbonyls in the urban Tokyo atmosphere. Atmos Environ 39, 1945-1960

549 Kawamura K, Pavuluri C (2010) New Directions: Need for better understanding of plastic waste burning as
 550 inferred from high abundance of terephthalic acid in South Asian aerosols. Atmos Environ 44, 5320-5321

551 Kawamura K, Tachibana E, Okuzawa K, Aggarwal SG, Kanaya Y, Wang ZF (2013) High abundances of

552 water-soluble dicarboxylic acids, ketocarboxylic acids and α -dicarbonyls in the mountaintop aerosols over
 553 the North China Plain during wheat burning season. *Atmos Chem Phys* 13, 8285-8302

554 Kim HS, Huh JB, Hopke PK, Holsen TM, Yi SM (2007) Characteristics of the major chemical constituents
 555 of PM_{2.5} and smog events in Seoul, Korea in 2003 and 2004. *Atmos Environ* 41, 6762-6770

556 Kroll JH, Ng NL, Murphy SM, Varutbangkul V, Flagan RC, Seinfeld JH (2005) Chamber studies of
 557 secondary organic aerosol growth by reactive uptake of simple carbonyl compounds. *J Geophys Res* 110,
 558 D23207

559 Kundu S, Kawamura K, Andreae T, Hoffer A, Andreae M (2010) Molecular distributions of dicarboxylic
 560 acids, ketocarboxylic acids and α -dicarbonyls in biomass burning aerosols: implications for photochemical
 561 production and degradation in smoke layers. *Atmos Chem Phys* 10, 2209-2225

562 Lee S, Ho K, Chan L, Zielinska B, Chow JC (2001) Polycyclic aromatic hydrocarbons (PAHs) and
 563 carbonyl compounds in urban atmosphere of Hong Kong. *Atmos Environ* 35, 5949-5960

564 Lei HC, Tanner PA, Huang MY, Shen ZL, Wu YX (1997) The acidification process under the cloud in
 565 southwest China: observation results and simulation. *Atmos Environ* 31, 851-861

566 Li Y, Allen PA, Densmore AL, Qiang X (2003) Evolution of the Longmen Shan foreland basin (western
 567 Sichuan, China) during the Late Triassic Indosinian orogeny. *Basin Res* 15, 117-138

568 Li YC, Yu JZ (2010) Composition profile of oxygenated organic compounds and inorganic ions in PM_{2.5}
 569 in Hong Kong. *Environ Chem* 7, 338-349

570 Liggio J, Li SM, McLaren R (2005) Reactive uptake of glyoxal by particulate matter. *J Geophys Res* 110,
 571 D10304

572 Lim HJ, Carlton AG, Turpin BJ (2005) Isoprene forms secondary organic aerosol through cloud processing:
 573 Model simulations. *Environ Sci Technol* 39, 4441-4446

574 Liu Y, Zhu L, Shen X (2001) Polycyclic aromatic hydrocarbons (PAHs) in indoor and outdoor air of
 575 Hangzhou, China. *Environ Sci Technol* 35, 840-844

576 Ma S, Peng Pa, Song J, Zhao J, He L, Sheng G, Fu J (2010) Stable carbon isotopic compositions of organic
 577 acids in total suspended particles and dusts from Guangzhou, China. *Atmos Res* 98, 176-182

578 Miyazaki Y, Aggarwal SG, Singh K, Gupta PK, Kawamura K (2009a) Dicarboxylic acids and water -
 579 soluble organic carbon in aerosols in New Delhi, India, in winter: Characteristics and formation processes.
 580 *J Geophys Res* 114, D19206

581 Miyazaki Y, Kondo Y, Shiraiwa M, Takegawa N, Miyakawa T, Han S, Kita K, Hu M, Deng ZQ, Zhao Y
 582 (2009b) Chemical characterization of water - soluble organic carbon aerosols at a rural site in the Pearl
 583 River Delta, China, in the summer of 2006. *J Geophys Res* 114, D14208

584 Miyazaki Y, Kawamura K, Sawano M (2010) Size distributions and chemical characterization of water -
 585 soluble organic aerosols over the western North Pacific in summer. *J Geophys Res* 115, D23210

586 Mkoma SL, Kawamura K (2013) Molecular composition of dicarboxylic acids, ketocarboxylic acids,

587 α -dicarbonyls and fatty acids in atmospheric aerosols from Tanzania, East Africa during wet and dry
588 seasons. *Atmos Chem Phys* 13, 2235-2251

589 Narukawa M, Kawamura K, Takeuchi N, Nakajima T (1999) Distribution of dicarboxylic acids and carbon
590 isotopic compositions in aerosols from 1997 Indonesian forest fires. *Geophys Res Lett* 26, 3101-3104

591 Nel A (2005) Air pollution-related illness: effects of particles. *Science* 308, 804-806

592 Nirmalkar J, Deb MK, Deshmukh DK, Tsai YI, Verma SK (2014) Molecular markers in ambient aerosol in
593 the Mahanadi Riverside Basin of eastern central India during winter. *Environ Sci Pollut Res*,
594 DOI:10.1007/s11356-11014-13416-11354

595 Osborne JM, Lambert FH (2014) The missing aerosol response in twentieth-century mid-latitude
596 precipitation observations. *Nature Climate Change*

597 Pavuluri CM, Kawamura K, Swaminathan T (2010) Water - soluble organic carbon, dicarboxylic acids,
598 ketoacids, and α - dicarbonyls in the tropical Indian aerosols. *J Geophys Res* 115, D11302

599 Perrone M, Gualtieri M, Consonni V, Ferrero L, Sangiorgi G, Longhin E, Ballabio D, Bolzacchini E,
600 Camatini M (2013) Particle size, chemical composition, seasons of the year and urban, rural or remote site
601 origins as determinants of biological effects of particulate matter on pulmonary cells. *Environ Pollut* 176,
602 215-227

603 Rinaldi M, Decesari S, Carbone C, Finessi E, Fuzzi S, Ceburnis D, O'Dowd CD, Sciare J, Burrows JP,
604 Vrekoussis M (2011) Evidence of a natural marine source of oxalic acid and a possible link to glyoxal. *J*
605 *Geophys Res* 116, D16204

606 Robinson AL, Subramanian R, Donahue NM, Bernardo-Bricker A, Rogge WF (2006) Source
607 apportionment of molecular markers and organic aerosol. 3. Food cooking emissions. *Environ Sci Technol*
608 40, 7820-7827

609 Rogge WF, Hildemann LM, Mazurek MA, Cass GR, Simoneit BR (1991) Sources of fine organic aerosol.
610 1. Charbroilers and meat cooking operations. *Environ Sci Technol* 25, 1112-1125

611 Saarikoski S, Timonen H, Saarnio K, Aurela M, Järvi L, Keronen P, Kerminen V-M, Hillamo R (2008)
612 Sources of organic carbon in fine particulate matter in northern European urban air. *Atmos Chem Phys* 8,
613 6281-6295

614 Sandradewi J, Prevot AS, Szidat S, Perron N, Alfarra MR, Lanz VA, Weingartner E, Baltensperger U (2008)
615 Using aerosol light absorption measurements for the quantitative determination of wood burning and traffic
616 emission contributions to particulate matter. *Environ Sci Technol* 42, 3316-3323

617 Schlesinger RB (2007) The health impact of common inorganic components of fine particulate matter
618 (PM_{2.5}) in ambient air: a critical review. *Inhal Toxicol* 19, 811-832

619 Shah AS, Langrish JP, Nair H, McAllister DA, Hunter AL, Donaldson K, Newby DE, Mills NL (2013)
620 Global association of air pollution and heart failure: a systematic review and meta-analysis. *The Lancet* 382,
621 1039-1048

622 Singh HB, Salas LJ, Cantrell BK, Redmond RM (1985) Distribution of aromatic hydrocarbons in the
623 ambient air. *Atmos Environ* 19, 1911-1919

624 Suh I, Zhang R, Molina LT, Molina MJ (2003) Oxidation mechanism of aromatic peroxy and bicyclic
625 radicals from OH-toluene reactions. *J Am Chem Soc* 125, 12655-12665

626 Sun Y, Zhuang G, Tang A, Wang Y, An Z (2006) Chemical characteristics of PM_{2.5} and PM₁₀ in haze-fog
627 episodes in Beijing. *Environ Sci Technol* 40, 3148-3155

628 Tao J, Chai FH, Zhu LH, Gao J, Cao JJ, Wang QY, Luo L (2011) Characteristics and sources of
629 carbonaceous aerosol in the urban Chengdu during spring of 2009. *Acta Scientiae Circumstantiae* 31,
630 2756-2761

631 Volkamer R, Platt U, Wirtz K (2001) Primary and secondary glyoxal formation from aromatics:
632 Experimental evidence for the bicycloalkyl-radical pathway from benzene, toluene, and p-xylene. *J Phys*
633 *Chem A* 105, 7865-7874

634 Wang G, Kawamura K, Xie M, Hu S, Gao S, Cao J, An Z, Wang Z (2009a) Size-distributions of n-alkanes,
635 PAHs and hopanes and their sources in the urban, mountain and marine atmospheres over East Asia. *Atmos*
636 *Chem Phys* 9, 8869-8882

637 Wang G, Kawamura K, Cheng C, Li J, Cao J, Zhang RJ, Zhang T, Liu S, Zhao Z (2012) Molecular
638 distribution and stable carbon isotopic composition of dicarboxylic acids, ketocarboxylic acids, and
639 alpha-dicarbonyls in size-resolved atmospheric particles from Xi'an City, China. *Environ Sci Technol* 46,
640 4783-4791

641 Wang GH, Niu SL, Liu CE, Wang LS (2002) Identification of dicarboxylic acids and aldehydes of PM₁₀
642 and PM_{2.5} aerosols in Nanjing, China. *Atmos Environ* 36, 1941-1950

643 Wang GH, Kawamura K (2005) Molecular characteristics of urban organic aerosols from Nanjing: A case
644 study of a mega-city in China. *Environ Sci Technol* 39, 7430-7438

645 Wang GH, Kawamura K, Xie MJ, Hu SY, Cao JJ, An ZS, Weston JG, Chow JC (2009b) Organic molecular
646 compositions and size distributions of Chinese summer and autumn aerosols from Nanjing: Characteristic
647 haze event caused by wheat straw burning. *Environ Sci Technol* 43, 6493-6499

648 Wang SL, Chai FH, Zhang YH, Zhou LD, Wang QL (2004) Analysis on the Sources and Characters of
649 Particles in Chengdu. *Scientia Geographica Sinica* 24, 488-492

650 Wang Y, Zhuang G, Chen S, An Z, Zheng A (2007) Characteristics and sources of formic, acetic and oxalic
651 acids in PM_{2.5} and PM₁₀ aerosols in Beijing, China. *Atmos Res* 84, 169-181

652 Wang Y, Che H, Ma J, Wang Q, Shi G, Chen H, Goloub P, Hao X (2009c) Aerosol radiative forcing under
653 clear, hazy, foggy, and dusty weather conditions over Beijing, China. *Geophys Res Lett* 36, L06804

654 Warneck P (2003) In-cloud chemistry opens pathway to the formation of oxalic acid in the marine
655 atmosphere. *Atmos Environ* 37, 2423-2427

656 Watson JG, Chow JC, Houck JE (2001) PM_{2.5} chemical source profiles for vehicle exhaust, vegetative

657 burning, geological material, and coal burning in Northwestern Colorado during 1995. *Chemosphere* 43,
658 1141-1151

659 Yang F, Huang L, Duan F, Zhang W, He K, Ma Y, Brook J, Tan J, Zhao Q, Cheng Y (2011) Carbonaceous
660 species in PM 2.5 at a pair of rural/urban sites in Beijing, 2005–2008. *Atmos Chem Phys* 11, 7893-7903

661 Yao XH, Lau AP, Fang M, Chan CK, Hu M (2003) Size distributions and formation of ionic species in
662 atmospheric particulate pollutants in Beijing, China: 2—dicarboxylic acids. *Atmos Environ* 37, 3001-3007

663 Ye B, Ji X, Yang H, Yao X, Chan CK, Cadle SH, Chan T, Mulawa PA (2003) Concentration and chemical
664 composition of PM2.5 in Shanghai for a 1-year period. *Atmos Environ* 37, 499-510

665 Yu JZ, Huang XF, Xu JH, Hu M (2005) When aerosol sulfate goes up, so does oxalate: Implication for the
666 formation mechanisms of oxalate. *Environ Sci Technol* 39, 128-133

667 Zhao YL, Hu M, Slanina S, Zhang YH (2007) Chemical compositions of fine particulate organic matter
668 emitted from Chinese cooking. *Environ Sci Technol* 41, 99-105

669 Zimmermann J, Poppe D (1996) A supplement for the RADM2 chemical mechanism: The photooxidation
670 of isoprene. *Atmos Environ* 30, 1255-1269
671
672
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Table 1. Concentrations (ng m^{-3}) of diacids, oxoacids and α -dicarbonyls as well as concentrations ($\mu\text{g m}^{-3}$) of EC and OC in the $\text{PM}_{2.5}$ aerosols collected in Chengdu, China in winter 2012.

Components, abbreviation	Daytime (n=15)				Nighttime (n=17)			
	Min	Max	Mean	Med	Min	Max	Mean	Med
Dicarboxylic acids								
Oxalic, C_2	639	1996	1440	1590	443	2310	1320	1270
Malonic, C_3	85.5	341	227	244	3.49	394	206	222
Succinic, C_4	197	636	436	417	155	683	428	412
Glutaric, C_5	41.2	166	114	124	51.5	163	105	91.6
Adipic, C_6	37.5	156	88.3	92.4	44.4	128	79.0	80.3
Pimelic, C_7	8.27	47.3	34.4	41.9	6.06	51.34	30.5	32.7
Sebacic, C_8	BDL	63.9	8.41	BDL	BDL	BDL	BDL	BDL
Azelaic, C_9	64.0	296	154	150	52.4	191	116	108
Decanedioic, C_{10}	2.66	36.4	21.8	23.7	BDL	35.5	17.5	20.2
Undecanedioic, C_{11}	11.4	75.1	27.7	24.4	11.5	41.3	20.1	16.4
Methylmalonic, iC_4	4.57	17.2	11.0	10.5	3.10	17.6	9.87	8.68
Methylsuccinic, iC_5	30.9	72.5	51.5	50.1	29.8	107	59.3	58.4
2-Methylglutaric, iC_6	3.45	31.7	13.3	9.31	3.66	16.4	8.93	9.05
Maleic, M	22.4	60.2	41.6	42.0	33.1	155	59.5	54.0
Fumaric, F	BDL	17.2	9.00	8.93	BDL	22.8	9.94	8.45
Methylmaleic, mM	14.4	42.1	31.0	31.1	20.8	86.8	38.5	32.6
Phthalic, Ph	147	382	278	274	164	436	304	296
Isophthalic, iPh	4.99	22.8	12.8	13.2	6.37	21.4	14.6	14.1
Terephthalic, tPh	125	571	356	357	200	872	422	400
Malic, hC_4	7.51	18.5	12.1	11.6	6.45	15.0	11.3	11.4
Oxomalonic, kC_3	15.5	154	76.1	69.3	11.5	145	65.4	54.6
4-oxopimelic, kC_7	3.90	17.7	11.8	13.2	2.21	19.2	11.2	11.6
Subtotal	1490	4690	3450	3810	1410	5250	3330	3160
Oxocarboxylic acids								
Pyruvic, Pyr	46.7	176	121	119	41.3	189	114	110
2-Oxoethanoic (Glyoxylic),	178	864	445	412	149	882	416	365
3-Oxopropanoic, ωC_3	9.06	55.9	27.5	21.1	3.63	52.2	24.0	21.1
4-Oxobutanoic, ωC_4	48.6	151	103	93.2	48.4	157	97.1	85.9
5-Oxopentanoic, ωC_5	8.55	28.8	18.4	16.9	6.61	26.3	16.8	16.9
7-Oxoheptanoic, ωC_7	7.30	30.6	21.6	23.2	5.15	34.3	21.4	23.2
8-Oxooctanoic, ωC_8	8.69	42.3	25.8	26.3	7.63	47.5	25.7	25.7
9-Oxononanoic, ωC_9	3.55	22.9	12.0	11.8	5.14	29.0	14.1	14.2
Subtotal	320	1320	774	720	272	1380	729	674
α-Dicarbonyls								
Glyoxal, Gly	44.1	102	69.4	71.1	37.1	130	72.8	61.5
Methylglyoxal, MeGly	34.8	111	75.6	73.5	31.9	118	68.6	65.5
Subtotal	90.0	213	145	144	87.9	215	141	131
OC ($\mu\text{g m}^{-3}$)	28.8	83.5	57.0	57.1	35.1	82.7	57.5	57.6
EC ($\mu\text{g m}^{-3}$)	8.71	42.4	27.5	27.8	11.9	65.6	30.4	27.9

BDL: below detection limit. BDL is ca. 0.01 ng m^{-3} .

678 **Table 2.** Concentrations (ng m⁻³) and weight ratios of selected diacids and related compounds in the PM_{2.5} aerosols collected in Chengdu, China in winter 2012,
679 along with the literature data for urban aerosols.

Site/Type	Sampling Period	Particle Size	C ₂	C ₃	C ₄	C ₆	C ₉	Ph	tPh	Total Diacids	ωC ₂	C ₃ /C ₄	C ₆ /C ₉	Ph/C ₉	C ₂ /Diacids	Reference
14 Cities, China	Winter 13-14 Jan 2003	PM _{2.5}	558	40.6	79.7	15.0	28.9	78.2	—	904	37.8	0.51	0.52	2.71	0.62	Ho et al., 2007
Hong Kong, Yuen Long	Winter Jan 2004	PM _{2.5}	1224	100	34.9	16.4	91.1	94.5	—	1655	119	2.86	0.18	1.04	0.74	Li and Yu, 2010
Hong Kong, roadside	Winter Jan + Nov 2003	PM _{2.5}	478	89.1	71.9	10.7	16.8	78.0	—	858	43.2	1.24	0.64	4.64	0.56	Ho et al., 2006
Hong Kong, Urban/roadside	Winter Dec 2006~Jan 2007	PM _{2.5}	464	26.5	23.2	7.00	24.3	34.1	19.9	644	38.5	1.14	0.29	1.40	0.72	Ho et al., 2011
Guangzhou, China	Winter Dec 2006~Jan 2007	PM _{2.5}	182	13.3	18.4	5.39	16.3	91.8	31.8	384	13.0	0.72	0.33	5.63	0.47	Ho et al., 2011
Ulaanbaatar, Mongolia	Winter Nov 2007~Jan 2008	PM _{2.5}	107	13.0	63.0	11.0	42.0	54.0	130	536	55.0	0.21	0.26	1.29	0.20	Jung et al., 2010
Beijing, China Urban/PKU	Autumn Aug ~ Sep 2006	PM _{2.5}	449	28.1	49.9	17.4	34.8	77.5	32.4	760	25.1	0.56	0.50	2.23	0.59	Ho et al., 2010
Nanjing, China	Winter 15 Feb ~ 1 Mar	PM _{2.5}	880	126	195	49.3	102	—	—	1684	—	0.65	0.49	—	0.52	Wang et al., 2002
Chennai, India	Winter Jan ~ Feb 2007	PM ₁₀	435	58.7	43.3	5.6	12.6	19.4	45.6	673	41.2	1.36	0.45	1.55	0.65	Pavuluri et al., 2010
Xi'an, China	Winter Jan ~ Feb 2009	PM ₁₀	1162	58	98	6.9	15.0	196	250	1843	179	0.59	0.46	13.1	0.63	Chen et al., 2013
Tokyo, Japan	Winter 20-21 Nov 1989	TSP	186	40.5	47.4	14.2	20.6	24.0	—	438	41.5	0.85	0.69	1.17	0.42	Kawamura and Yasui, 2005
New Delhi, India	Sep 2006 ~ Apr 2007	TSP	955	181	273	29.0	70	40	5.0	1777	114	0.66	0.41	0.57	0.55	Miyazaki et al., 2009b
New Delhi, India	Daytime Sep 2006 ~ Apr 2007	TSP	1906	194	330	32.0	72	45	4.0	2875	128	0.59	0.44	0.63	0.66	Miyazaki et al., 2009b
Chengdu, China	Nighttime Winter Jan 2013 Daytime	PM _{2.5}	1439	227	436	88.3	154	278	356	3454	445	0.53	0.61	1.94	0.42	This study
Chengdu, China	Winter Jan 2013 Nighttime	PM _{2.5}	1315	206	428	79.0	116	304	422	3331	416	0.47	0.71	2.72	0.39	This study

Note: Ratios of C₃/C₄, C₆/C₉, and Ph/C₉ were calculated from average values in the references; “—” represents no data or no result.

Figure captions:

Fig. 1 Map showing the sampling site in Chengdu within the Sichuan Basin, southwest China.

Fig. 2 Molecular distributions of diacids, oxoacids, and α -dicarbonyls in fine aerosols (PM_{2.5}) collected in Chengdu, China in winter. Bars show the standard deviation of the concentrations of each species.

Fig. 3 Pie diagrams of relative abundances (%) of individual diacids in total straight-chain diacids (C₂–C₁₁) and two major aromatic diacids (Ph and tPh) in the fine aerosols.

Fig. 4 Temporal variations (ng m⁻³) of (a) total diacids, (b) total oxoacids, and (c) total α -dicarbonyls in the aerosol samples, along with API, in winter (open circle for daytime; solid circle for nighttime). API is the “Air Pollution Index”.

Fig. 5 Temporal variations (ng m⁻³) of (a) C₂, (b) C₃, (c) C₄, (d) C₆, (e) C₉, (f) Ph, and (g) tPh acids in the aerosol samples (PM_{2.5}) (open circle for daytime; solid circle for nighttime).

Fig. 6 Temporal variations (ng m⁻³) of (a) Pyr, (b) ω C₂, (c) Gly, and (d) MeGly in the aerosol samples (PM_{2.5}) (open circle for daytime; solid circle for nighttime).

Fig. 7 Correlation plots for the concentrations of oxalic acid (C₂) and sulfate (SO₄²⁻) in the aerosol samples.

Fig. 8 Scatterplots of (a) C₄, (b) C₃, (c) kC₃, (d) ω C₂, (e) ω C₄, and (f) Pyr with C₂ in the aerosol samples.

Fig. 9 Temporal variations of (a) C₃/C₄, (b) C₆/C₉, (c) Ph/C₉, and (d) Ph/C₆ in the aerosol samples (open circle for daytime; solid circle for nighttime).

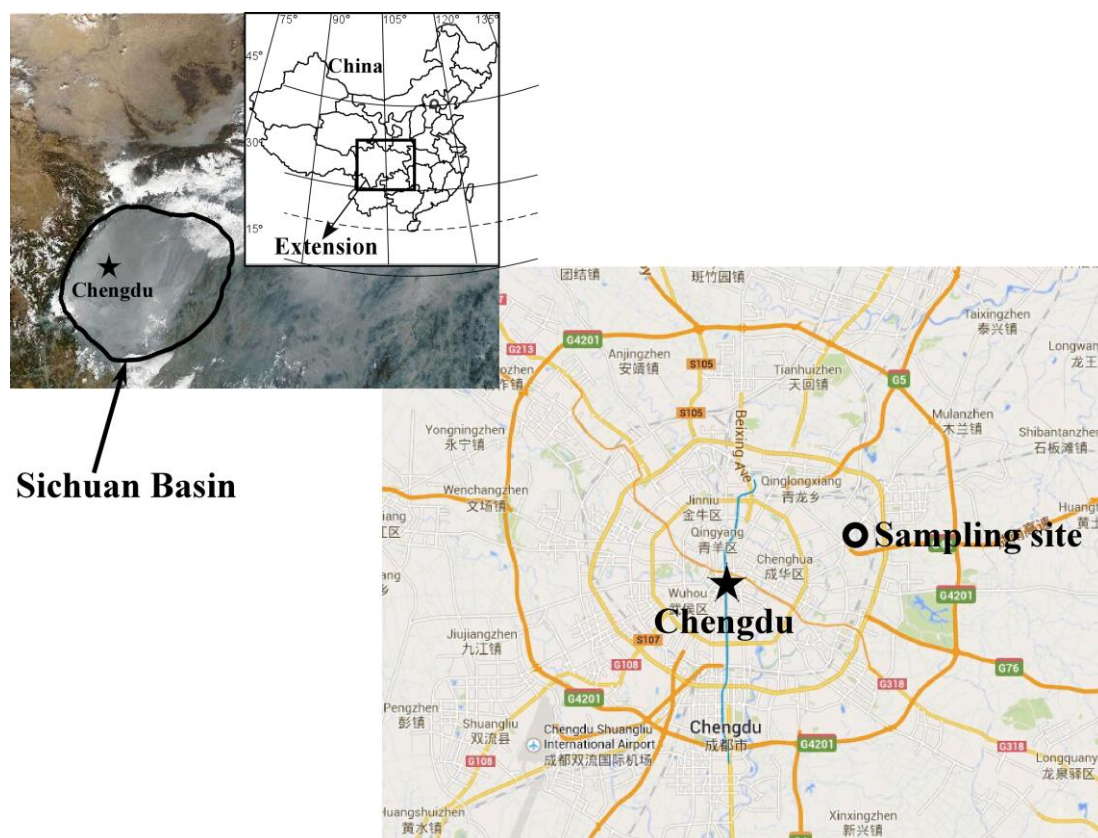


Fig. 1 A map showing the sampling site in Chengdu within the Sichuan Basin, Southwest China.

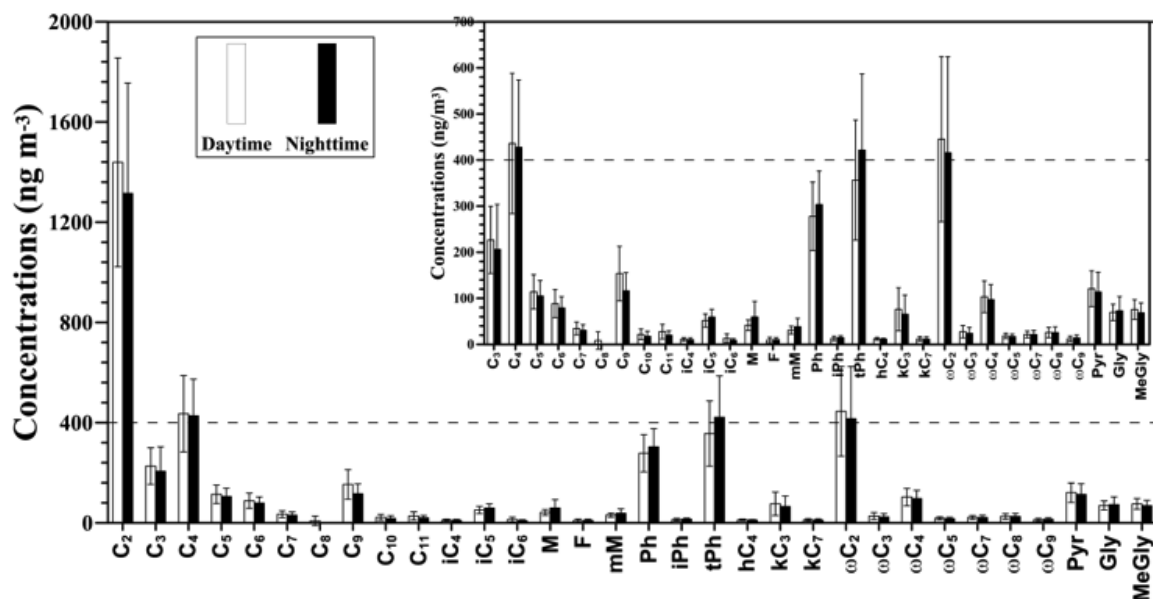


Fig. 2 Molecular distributions of diacids, oxoacids, and α -dicarbonyls in fine aerosols (PM_{2.5}) collected in Chengdu, China in winter. Bars show the standard deviation of the concentrations of each species.

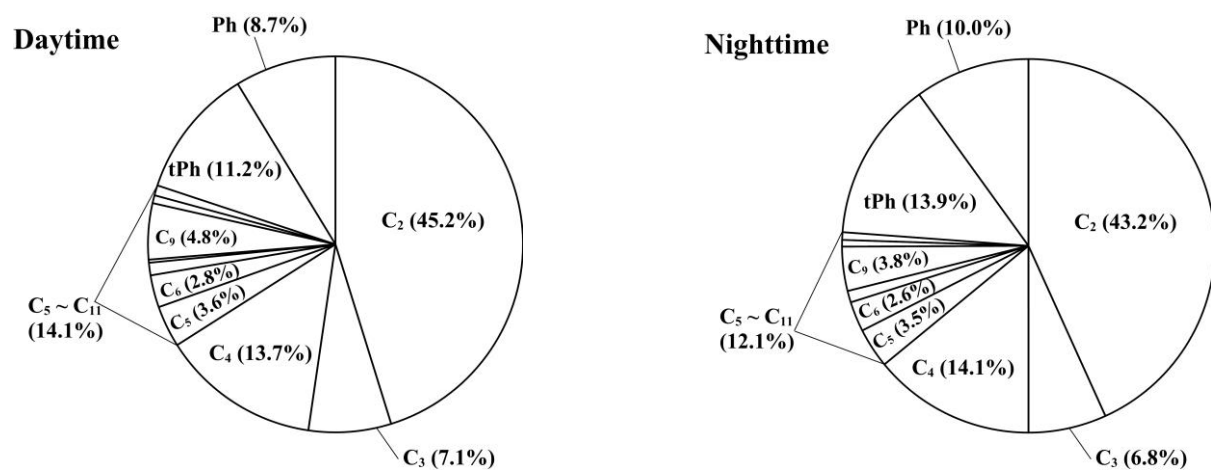


Fig. 3 Pie diagrams of relative abundances (%) of individual diacids in total straight-chain diacids (C₂–C₁₁) and two major aromatic diacids (Ph and tPh) in the fine aerosols.

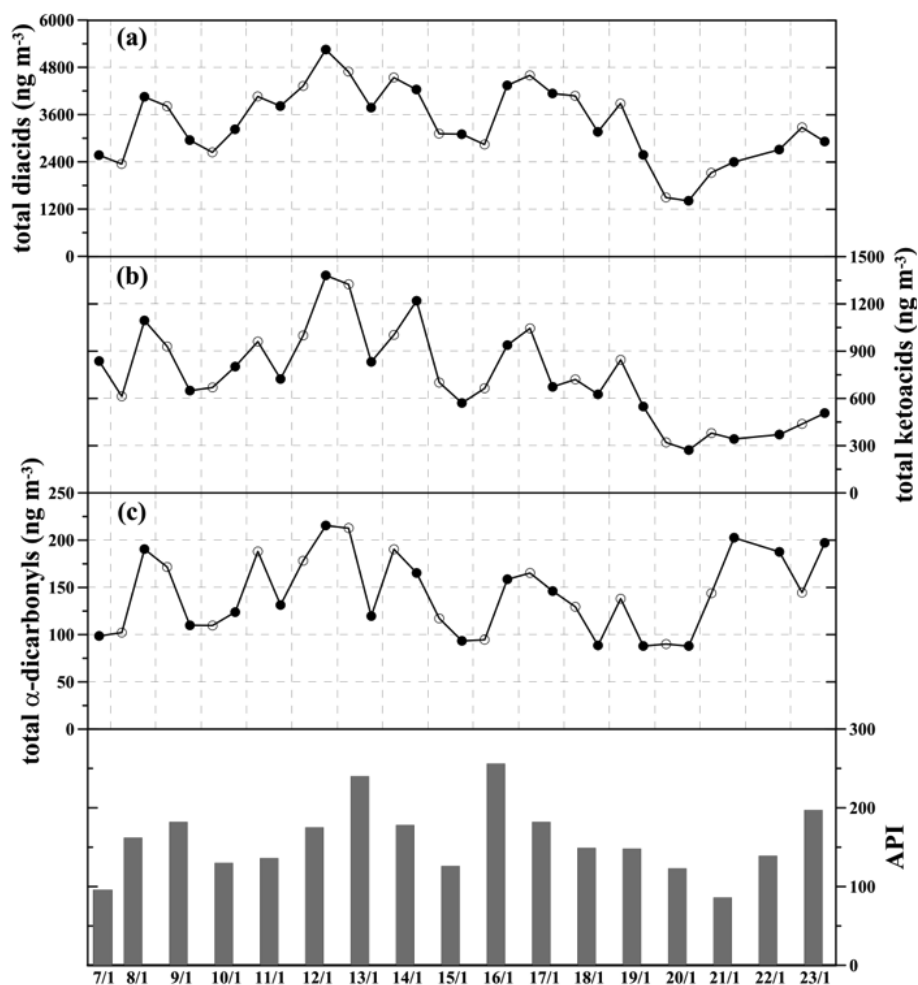


Fig. 4 Temporal variations (ng m⁻³) of (a) total diacids, (b) total oxoacids, and (c) total α-dicarbonyls in the aerosol samples, along with API, in winter (open circle for daytime; solid circle for nighttime). API is the “Air Pollution Index”.

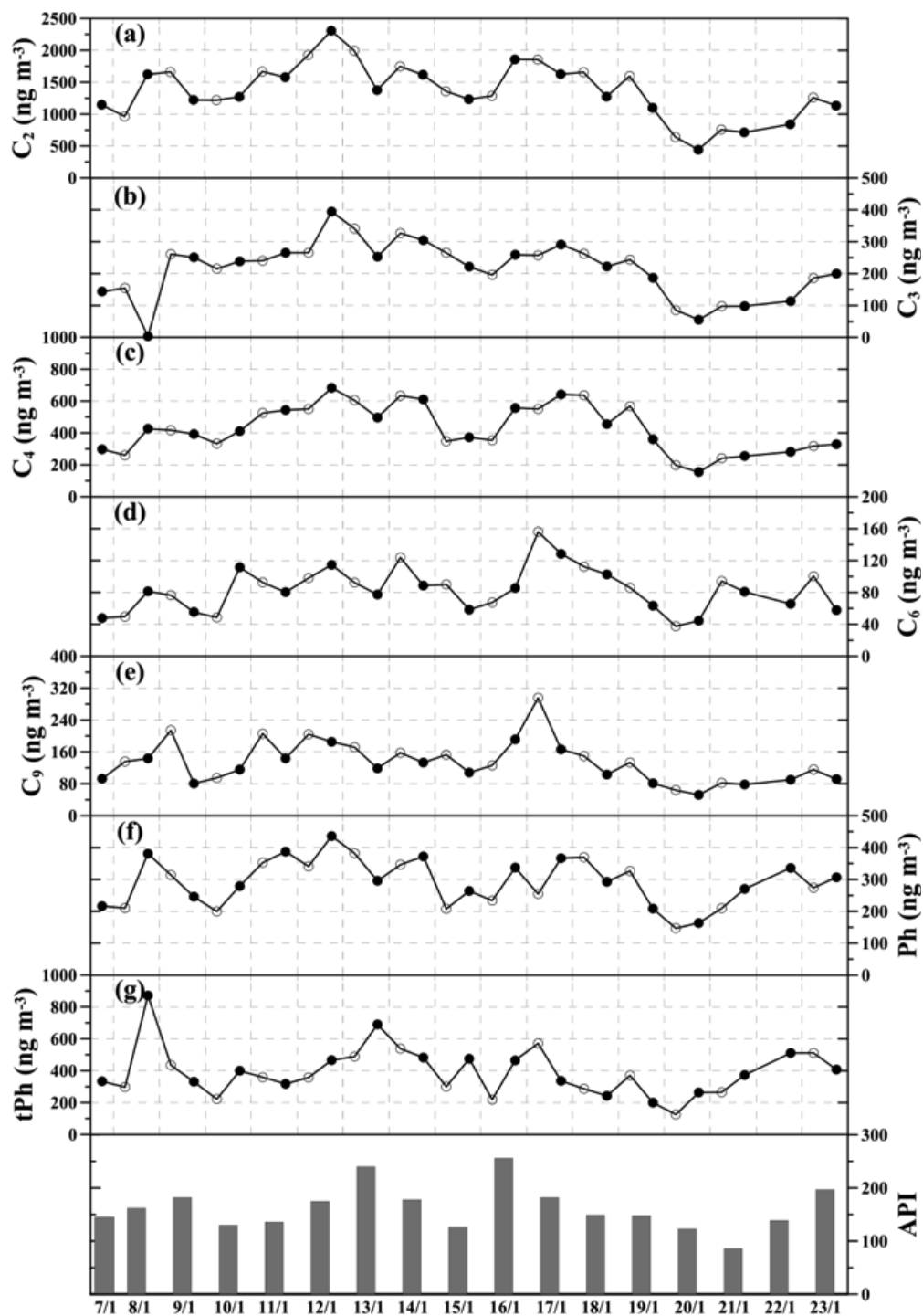


Fig. 5 Temporal variations (ng m⁻³) of (a) C₂, (b) C₃, (c) C₄, (d) C₆, (e) C₉, (f) Ph, and (g) tPh acids in the aerosol samples (PM_{2.5}) (open circle for daytime; solid circle for nighttime).

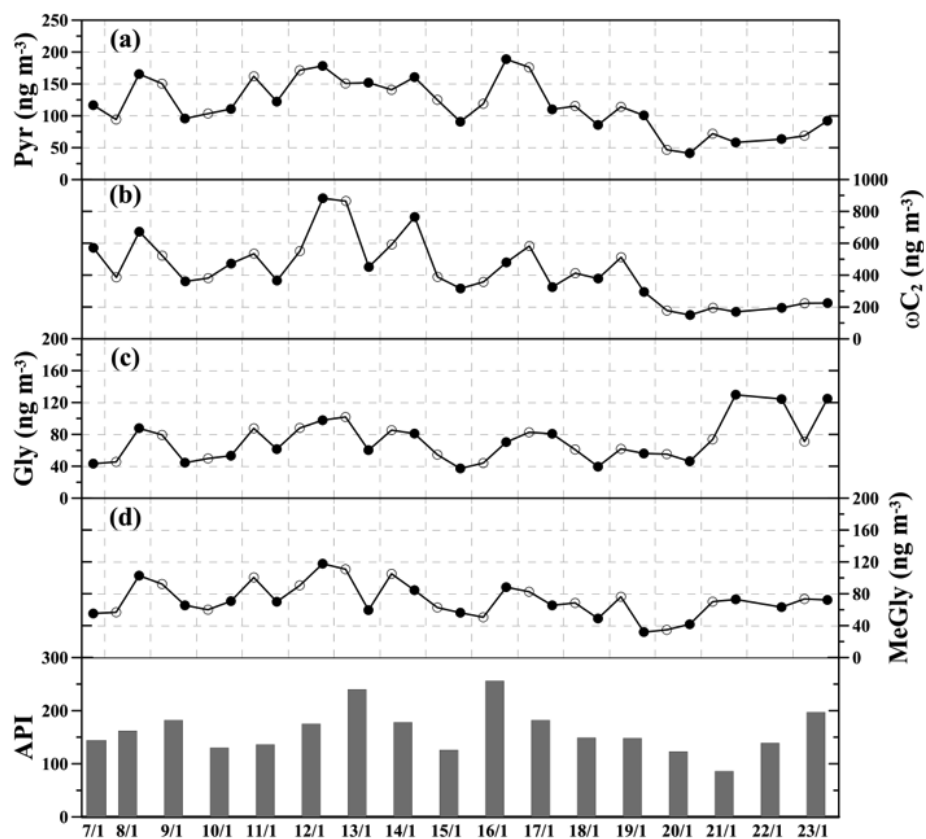


Fig. 6 Temporal variations ($ng\ m^{-3}$) of (a) Pyr, (b) ωC_2 , (c) Gly, and (d) MeGly in the aerosol samples ($PM_{2.5}$) (open circle for daytime; solid circle for nighttime).

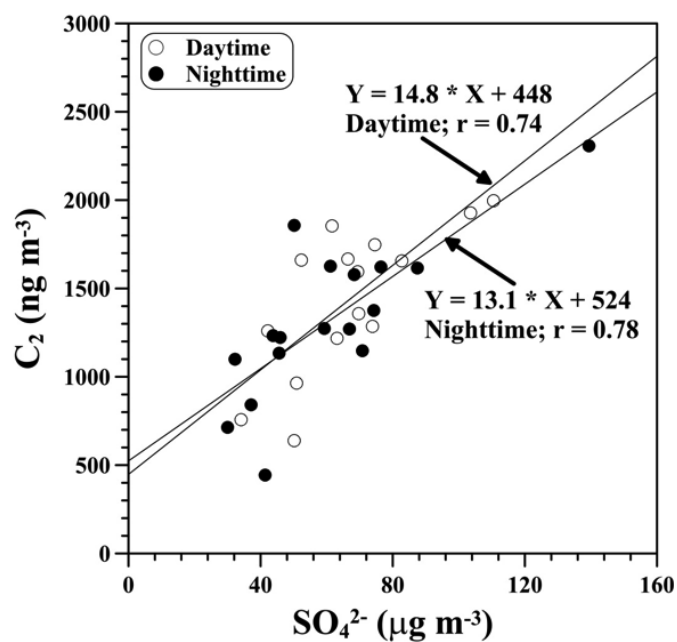


Fig. 7 Correlation plots for the concentrations of oxalic acid (C_2) and sulfate (SO_4^{2-}) in the aerosol samples.

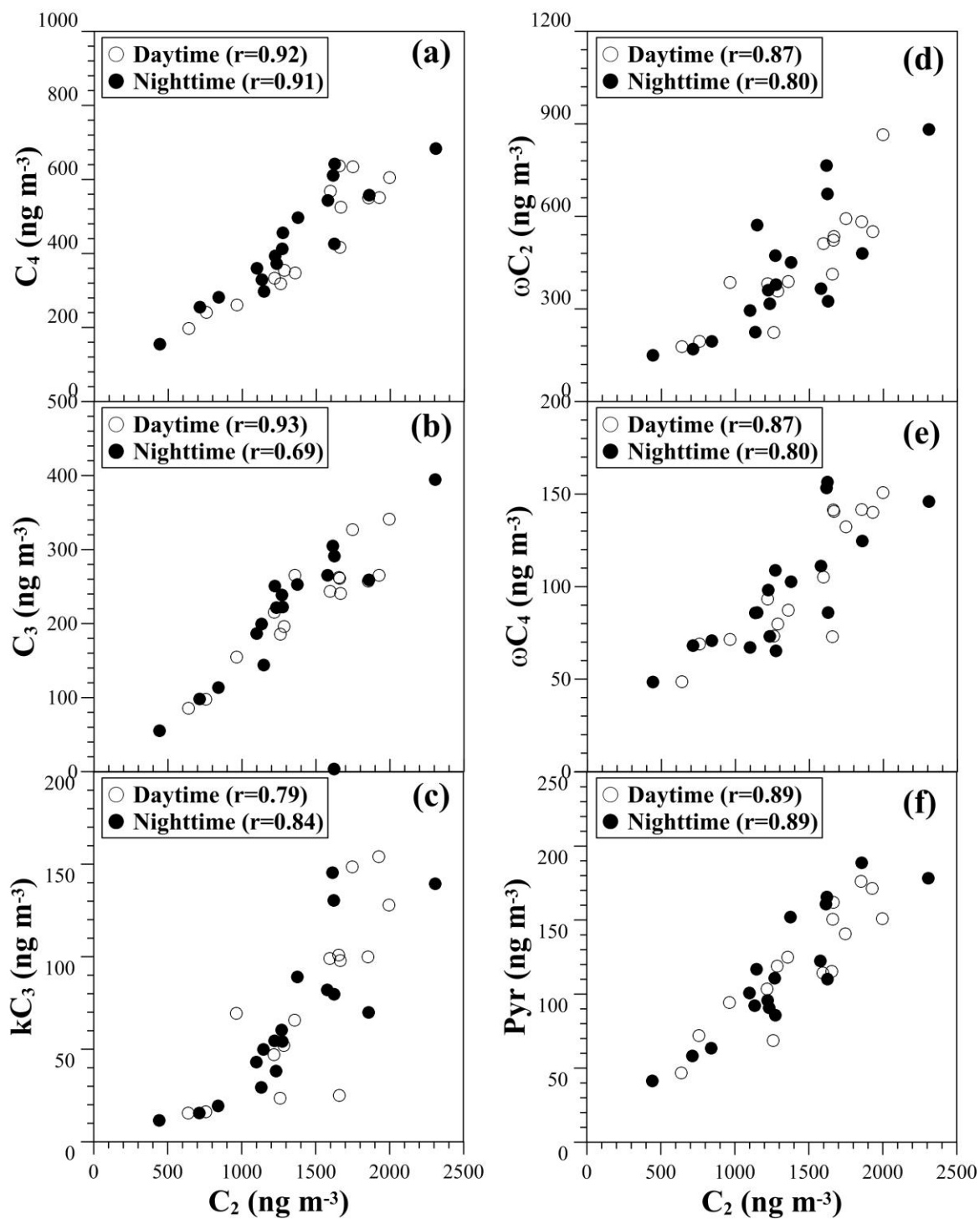


Fig. 8 Scatterplots of (a) C_4 , (b) C_3 , (c) kC_3 , (d) ωC_2 , (e) ωC_4 , and (f) Pyr with C_2 in the aerosol samples.

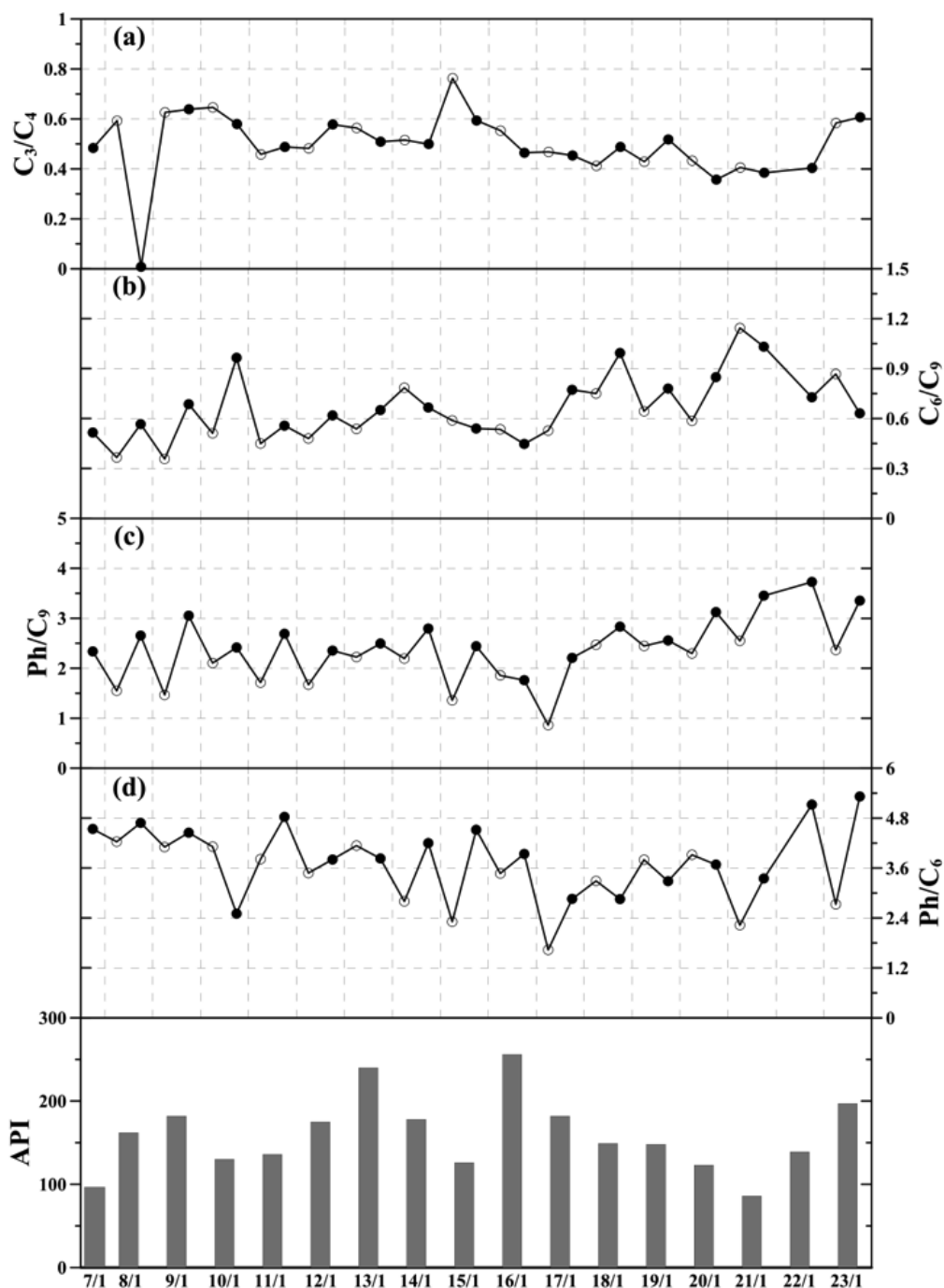


Fig. 9 Temporal variations of (a) C_3/C_4 , (b) C_6/C_9 , (c) Ph/C_9 , and (d) Ph/C_6 in the aerosol samples (open circle for daytime; solid circle for nighttime).