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Abstract

Purpose: Fine roots play an essential role in global carbon cycles, but phenological variations in root function and metabolism are poorly understood. To illustrate the dynamics of fine root function and metabolism in the field, we partitioned root respiration (R_r) into growth (R_g), maintenance (R_m), and ion uptake (R_{ion}) components using a modified traditional model.

Methods: A year-round experiment was conducted in a young larch-dominated forest regrowing on bare soil. Soil respiration was measured with a chamber method and partitioned into R_r and heterotrophic respiration by trenching. Fine root biomass and production were measured simultaneously. Using the field data, the model was parameterized, and R_r was further partitioned.

Results: Annually, R_r ($210\text{--}253\text{ g C m}^{-2}\text{ yr}^{-1}$) accounts for 45–47% of the total soil respiration. The contribution of fine root R_g , fine root R_m , coarse root R_m , and fine root R_{ion} were 26–40, 46–51, 10–16, and 12%, respectively. The R_g contribution showed a clear seasonal variation, with a peak in mid-spring and a minimum in early fall, mainly because of different seasonality between fine root production and soil temperature.

Conclusion: The model parameters were consistent with those from our previous study conducted by the same method in the same site. Thus, we believe that our approach was robust under a relatively simple condition. However, our growth respiration parameter resulting from only field data was much higher than those from laboratory experiments. To further improve our understanding of root respiration, more field data should be accumulated.

Keywords: Chamber, fine root, root biomass, root production, sap flow, soil respiration

1. Introduction

Soil respiration (R_s) is composed of autotrophic respiration by plant roots (R_r) and heterotrophic respiration (R_h) by soil microorganisms and fauna. R_r corresponds to mycorrhizosphere respiration consisting of living root, rhizomicrobial, and mycorrhizal respiration. R_h is equivalent to the decomposition of soil organic matter (SOM) and root litter by soil microbes (Kuzyakov 2006; Moyano et al. 2009). The global annual R_r was reported to be approximately 44 Pg C yr⁻¹ (Tang et al. 2019), which is approximately fourfold greater than anthropogenic carbon emissions (Friedlingstein et al. 2020). The contribution of R_r to R_s varies from 10 to 90% in forest ecosystems (Hanson et al. 2000) but is typically 45–50% annually (Subke et al. 2006). Because R_r and R_h respond differently to environmental factors, such as temperature and water content (Boone et al. 1998; Lavigne et al. 2004; Scott-Denton et al. 2006), R_s should be partitioned into two components to understand ecosystem-scale carbon cycling.

Plant root systems are composed of fine and coarse roots that serve contrasting functions. Fine roots, commonly defined as roots thinner than 2 mm in diameter (Brunner et al. 2013; Finér et al. 2011b), absorb water and nutrients from the soil. Despite their small biomass, fine roots have a large net primary production (NPP), accounting for 22% of terrestrial NPP (McCormack et al. 2015a) because of their fast turnover rates (Brunner et al. 2013; Finér et al. 2011b). Thus, fine roots play a dominant role in belowground carbon cycling (Finér et al. 2011b; Richter et al. 1999). Although fine root phenology strongly influences belowground carbon dynamics, seasonality and variability in fine root function are poorly understood (Abramoff and Finzi 2015; McCormack et al. 2014; Radville et al. 2016).

Autotrophic respiration is further separated into growth (R_g) and maintenance (R_m) components (Amthor 2000; McCree 1974; Penning de Vries 1974; Thornley 1970) and sometimes into an ion uptake component (R_{ion}) for fine root respiration (Chapin et al. 2011; Johnson 1990; Lambers et al. 2008) using a conceptual model. In the model, R_g , R_m , and R_{ion} were linearly correlated with growth (production), biomass, and ion uptake, respectively. However, the model has been criticized for its weak scientific basis in quantitative partitioning into R_g and R_m (Cannell and Thornley 2000; Sweetlove et al. 2013; Thornley 2011), although experimental results indicate that the model is useful for understanding ecological control of autotrophic respiration (Chapin et al. 2011; Lambers et al. 2008). Most terrestrial biosphere models do not explicitly incorporate root respiration because of a lack of mechanistic

respiration models (Collalti et al. 2020; Hopkins et al. 2013; Sweetlove et al. 2013; Warren et al. 2015). Respiration is frequently estimated from other processes using correlation (Sweetlove et al. 2013), and important processes contributing to respiration have been oversimplified (Ballantyne et al. 2017). The partitioning model is effective for understanding and quantifying intrinsic processes, such as acclimation to warming.

The partitioning model has been applied to R_r measured through laboratory experiments in controlled environments (Lambers et al. 2008; Thongo M'Bou et al. 2010). In the field, (George et al. 2003) partitioned R_r into R_g , R_m , and R_{ion} (nitrogen uptake), based on laboratory data and a short-term field experiment. (Sun et al. 2020) conducted a year-round experiment in two mature forests and estimated annual R_g and R_m . Because experimental conditions in the field are complex, especially in mature natural forests, the uncertainty in partitioning should be large. To decrease the uncertainty, (Cui et al. 2021) conducted an experiment in a young forest regenerating on bare soil in relatively simple conditions resulting from homogeneous tree age, little litter accumulation, limited coarse roots, little ground vegetation, and poor SOM. However, their study's spatial replication was small ($n = 10$), and R_{ion} was ignored. More field data are indispensable for the robustly parameterization and verification of the partitioning model. Thus, a field experiment was conducted at the same site as (Cui et al. 2021). We increased the number of replications and added R_{ion} as a respiratory component. The objectives of this study were 1) to modify the partitioning model including R_{ion} , 2) to robustly parameterize the modified model using a large size of field data, 3) to quantitatively partition R_r into R_g , R_m , and R_{ion} , and 4) to show the seasonal variations of the R_r components.

2. Material and Methods

2.1. Study site

An experiment was conducted in a young forest in Hokkaido, Japan ($42^{\circ}44.27'N$, $141^{\circ}31.42'E$, 116 m above sea level), which was the same site as (Cui et al. 2021). The forest was dominated by Japanese larch (*Larix kaempferi*) and dotted with Japanese white birch (*Betula platyphylla*). The site was used as a mature larch plantation growing on volcanogenous regosol but was severely damaged by a windstorm in 2004 (Sano et al. 2010). All tree stems were removed in 2005. In 2006, a thin layer of organic topsoil

(A horizon), coarse woody debris, stumps, accumulated litter, regenerating ground vegetation, and buried seeds were removed. As a result, volcanic pumice stones (C horizon) were exposed because B horizon was originally lacking. Wind-blown larch seeds germinated in 2007. The ground was sparsely covered with understory species. Aboveground biomass of trees taller than 2 m was 27.3 ± 12.9 and $32.8 \pm 9.8 \text{ t ha}^{-1}$ (mean $\pm$ standard deviation (SD) of three plots of $20 \text{ m} \times 20 \text{ m}$) in September 2019 and September 2020, respectively, of which larch accounted for 82% and 87%, respectively. Belowground biomass of larch coarse roots was 4.50 ± 2.10 and $5.57 \pm 1.75 \text{ t ha}^{-1}$ in the respective years. Tree biomass was estimated from diameter at breast height (DBH) using allometric equations (Yazaki et al. 2016). Soil bulk density, total carbon (C) and nitrogen (N) concentrations were $0.446 \pm 0.042 \text{ g cm}^{-3}$, $15.2 \pm 13.8 \text{ g kg}^{-1}$, and $0.742 \pm 0.794 \text{ g kg}^{-1}$ (mean $\pm$ SD, $n = 32$) for the top 15 cm fine soil layer ($< 2 \text{ mm}$) in 2016. Soil C concentration decreased with distance from tree stems, mainly because of decreasing litter fall with distance.

Mean annual air temperature and precipitation were $8.1 \pm 0.3^{\circ}\text{C}$ and $1305 \pm 202 \text{ mm yr}^{-1}$, respectively, from 2011 through 2020 at an observatory (Tomakomai) 14 km from the study site. The mean monthly air temperature was highest in August (21.0°C) and lowest in January (-3.9°C). Snow usually covers the ground from early December to early April.

2.2. Experimental design

Nine pairs of aluminum collars ($0.5 \text{ m} \times 0.5 \text{ m}$) were concentrically installed 0.5 m and 1.0 m, respectively, from nine isolated larch trees in $50 \text{ m} \times 50 \text{ m}$ in July 2019 (Fig. 1). Because root density and root respiration were expected to depend on distances from tree stems, we installed collars at the two positions to ensure a wide range of data. For each pair, collars were positioned at a 0.3–0.4 m interspace and inserted 3 cm deep into the soil. To exclude root respiration, four PVC boards were inserted 30 cm into the soil around the collar (TC) for each pair in July 2019. The soil profiles showed that almost all roots were distributed in the top 15 cm. Fine roots were sampled from the other collar (SC).

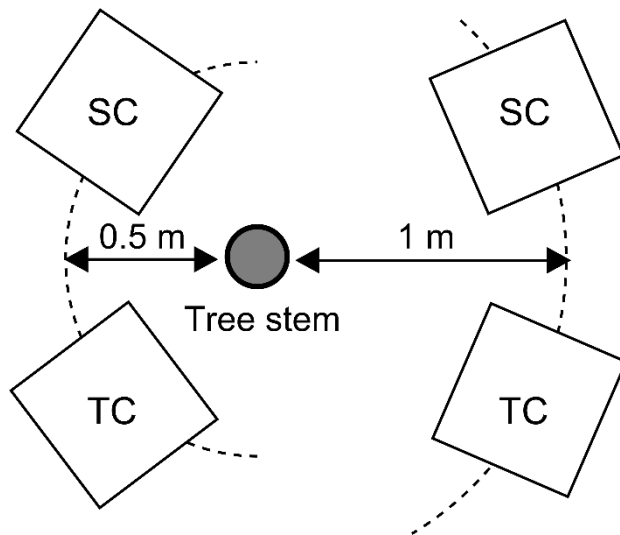


Fig.1. Layout of collars for root sampling (SC) and trenching (TC) 0.5 m and 1.0 m from the tree stem.

2.3. Soil CO₂ flux

Soil CO₂ flux was measured on each collar using a previously described method (Cui et al. 2021; Sun et al. 2017) between 10:00 and 16:00 at intervals of approximately three weeks from August 2019 through November 2020 with a five-month suspension from mid-November 2019 through mid-April 2020. Seedlings were carefully pulled from the collars before flux measurements, although they were rarely found. Flux was measured using a portable system composed of two 0.5-m-tall cubic chambers and a CO₂ analyzer (LI820; Li-Cor Inc., Lincoln, NB, USA). The two chambers were automatically closed for 3 min and opened sequentially. Thus, the flux measurement on the two collars took 6 min. During closing, the CO₂ concentration was measured every 5 s, and its rate of increase was determined using the least-squares method to calculate CO₂ flux. In each collar, soil temperature at a depth of 5 cm and volumetric soil moisture of the top 5 cm were measured immediately after the flux measurement. In addition, soil temperature and soil moisture were recorded half-hourly at depths of 6 cm and 3 cm, respectively, at a station (Hirano et al. 2017) approximately 150 m from the study site.

After trenching, CO₂ flux from the TC (R_{TC}) was equivalent to the sum of the original R_h and additional CO₂ flux resulting from the decomposition of dead roots (R_{DR}) caused by trenching. R_{DR} estimation method is described later. Although the CO₂ flux from the SC (R_{SC}), which was equivalent to the total R_s , was influenced by core sampling, the sampling effect on the CO₂ flux was expected to

negligible owing to the limited sampling area (Sun et al. 2020). R_r was estimated for each collar pair as $R_r = R_{SC} - R_{TC} + R_{DR}$ under the assumption that R_h was identical in each pair. Because they were concentrically installed (Fig. 1), SOM and litter fall were assumed to be the same between the two collars in each pair.

The following equation was applied to relate the CO₂ flux (R_c , $\mu\text{mol m}^{-2} \text{s}^{-1}$) to soil temperature (T_s , °C) for each collar.

$$R_c = a \cdot \exp(b \cdot T_s) \quad (1)$$

where a and b are the fitting parameters. Using this equation, R_{SC} and R_{TC} were calculated half-hourly from half-hourly soil temperature, and daily R_r was calculated from daily R_{SC} , R_{TC} and R_{DR} .

2.4. Decomposition of dead roots

The daily R_{DR} ($\text{g C m}^{-2} \text{d}^{-1}$) in TC at elapsed time t (days) was calculated using the following equation for fine and coarse roots, respectively:

$$R_{DR} = C_c \cdot (X_{t-1} - X_t) = C_c \cdot X_0 \cdot \exp(-k \cdot t) \cdot \{\exp(k) - 1\} \quad (2)$$

where C_c is the C concentration of roots (g g^{-1}), X_t is the dry weight of the remaining dead roots (g m^{-2}) at t , X_0 is the initial dry weight of dead roots (g m^{-2}), and k is the decay constant (d^{-1}). X_0 was set for each TC based on the root biomass measured in its paired SC by soil core sampling in September 2019 for fine roots and soil bulk sampling in April 2021 for coarse roots. The root biomass (X_0) was 134 ± 54 at 0.5 m and 111 ± 75 g m^{-2} at 1.0 m for fine roots, and 41.6 ± 10.0 at 0.5 m and 18.3 ± 6.3 g m^{-2} at 1.0 m for coarse roots (mean $\pm$ SD, $n = 9$). The diameter of the coarse roots was mostly less than 10 mm with a rough average of 5 mm. The C_c was set separately for fine and coarse roots based on the CN analysis of root samples collected in April 2021; C and N concentrations were 0.455 ± 0.026 g g^{-1} and 9.22 ± 0.83 mg g^{-1} , respectively, for fine roots, and 0.489 ± 0.014 g g^{-1} and 7.10 ± 1.25 mg g^{-1} , respectively, for coarse roots (mean $\pm$ SD, $n = 10$). The C and N concentrations differed significantly between fine and coarse roots ($P < 0.01$, two-sided t -test). We used k values determined from litter bag experiments in our previous studies. For fine roots, the k was set at $2.1 \times 10^{-3} \pm 7.4 \times 10^{-4} \text{d}^{-1}$ ($\pm$ standard error (SE)) before mid-November 2019, 0.0d^{-1} in winter, and $1.7 \times 10^{-3} \pm 4.5 \times 10^{-6} \text{d}^{-1}$ after mid-April 2020 (Cui et al. 2021), whereas it was set at $6.8 \times 10^{-4} \pm 4.5 \times 10^{-5} \text{d}^{-1}$ for coarse roots throughout the period (Sun et al.

2020).

2.5. Biomass and production of fine roots

The same method as in (Cui et al. 2021) was applied to measure fine root biomass and production. Biomass density (B_f , g m⁻²) was determined by soil coring eight times from September 2019 through November 2020 with a suspension for five months in winter. Soil cores were collected down to 15 cm using a stainless-steel edged tube with an inner diameter of 2.4 cm. Three cores were collected from randomly selected single-use grid positions with an 8 cm spacing in each SC. The area of pits caused by core sampling came to 109 cm² in total for each SC ($= 4.52 \text{ cm}^2 \times 3 \text{ positions} \times 8 \text{ times}$), accounting for 4.3% of the collar area (0.5 m $\times$ 0.5 m). Core samples were stored in PVC tubes in a freezer. Living fine roots were visually extracted from the soil samples dispersed in tap water, dried at 70°C for 48 h, and weighed to determine biomass.

The ingrowth core method was applied to measure production (P_f , g m⁻² period⁻¹). Plastic hair curlers with an outer diameter of 2.3 cm were wrapped in a 2 mm mesh fabric and filled with air-dried root-free soil. The soil was collected from the study site and sieved through 2 mm meshes. In each SC, three ingrowth cores were inserted down to 15 cm into the pit dug through soil core sampling, and the three ingrowth cores installed the previous time were collected. Fine root biomass in the cores, corresponding to root production during the interval, was analyzed using the same method as described above. In addition, annual mortality (M_f , g m⁻² yr⁻¹) were estimated by subtracting the annual difference in root biomass (ΔB_f) from annual production ($M_f = P_f - \Delta B_f$). The dry weight was converted to C by using C_c (0.455 g g⁻¹). In addition, turnover rates were determined by dividing the annual production by the mean biomass (Brunner et al. 2013).

2.6. Sap flow

The sap flow was measured to quantify ion uptake respiration (R_{ion}). In laboratory experiments, R_{ion} was usually correlated with N uptake as a major nutrient ion, as determined by destructive sampling (Lambers et al. 2008). However, applying this method to trees is difficult in the field because the N analysis of whole trees is expensive and labor-intensive. It was reported that N uptake depends on root

water uptake, causing water mass flow in the soil, especially when root density is low (Henriksson et al. 2021; McMurtrie and Nasholm 2018; Oyewole et al. 2014). We adopted the sap flow rate as a proxy for water uptake and incorporate it into the partitioning model described below.

Sap flow velocity was measured using the thermal dispersion method (Granier 1987) in three trees from July 2019 through November 2020 with a suspension during the leafless season. Sap flow sensors (CUP-SPF-M; Climatec Inc., Tokyo, Japan) were installed 25 cm below the lowest branch to measure the entire sap flow of each tree. From the destructive sampling, we presumed that the stem area at the sensor height (25–72 cm²) was occupied by sapwood, excluding bark. Thus, sap flow rates were calculated as the product of the sap flow velocity and the stem area. Although sap flow, transpiration, and water uptake were not the same, their daily rates were almost identical. Because DBH of larch trees averaged 54.4 ± 28.3 mm in 2019 ($n = 215$) and 58.2 ± 30.4 mm in 2020 ($n = 225$) in the study site, the three trees with DBH of 37–84 mm in 2019 and 43–91 mm in 2020 were almost within the range of mean $\pm$ SD in size. The total coarse root biomass of each sample tree was estimated from the DBH using an allometric equation, and the total fine root biomass of each tree was estimated from the coarse root biomass using a factor of 0.185, which was determined from boreal trees (Yuan and Chen 2010). Specific sap flow rates normalized by fine root biomass (S_s ; g H₂O g dry matter (DM)⁻¹ d⁻¹) were calculated for each tree and averaged. The water uptake rate in each SC was estimated as a product of the average S_s and fine root biomass.

2.7. Partitioning of root respiration

R_r can be partitioned into R_g , R_m , and R_{ion} using the conceptual model below (Amthor 2000; Lambers et al. 2008; Thornley 1970).

$$R_r = R_g + R_m + R_{ion} = g \cdot P_f + m \cdot B_f + u \cdot U_i \quad (3)$$

where g , m , and u are the coefficients of growth, maintenance, and ion uptake respiration, and U_i is the ion uptake rate. Considering the field conditions, we modified the model, including one without R_{ion} to be consistent with our previous study (Cui et al. 2021), as follows.

$$R_r = R_g + R_m + R_{ion} = g \cdot P_f + d \cdot \exp(f \cdot T_s) \cdot (B_f + B_c) + u \cdot S_s \cdot B_f: \text{Model 1 (4)}$$

$$R_r = R_g + R_m = g \cdot P_f + d \cdot \exp(f \cdot T_s) \cdot (B_f + B_c): \text{Model 2 (5)}$$

where g is the growth coefficient (g C g DM^{-1}), d is the R_m of the unit biomass at 0°C ($\text{g C g DM}^{-1} \text{ d}^{-1}$), f is the temperature coefficient ($^\circ\text{C}^{-1}$), B_c is the coarse root biomass (g m^{-2}), and u is the ion uptake coefficient ($\text{g C g H}_2\text{O}^{-1}$). Although respiration components and root production were originally expressed per unit root biomass, they are per unit ground area here, because respiration per area is easier to measure in the field and more useful to quantify ecosystem carbon cycles. In these models, temperature affects only R_m exponentially (Moyano et al. 2009; Thornley 2011). Although the tree survey indicated coarse root growth during the study period, the growth was ignored in R_g because of the lack of data. However, coarse root biomass was incorporated into R_m based on the assumption that the temperature responses of fine and coarse roots were the same. A nonlinear mixed-effects model was applied to the time-series (eight times) datasets for 18 pairs ($n = 8 \times 18$) to parameterize the models. The data preparation for the parametrization is summarized in Fig. 2. Fine root production was measured eight times and converted into daily values using the day length of the measurement intervals. Fine root biomass was the average of two consecutive measurements at the beginning and end of each interval. Soil temperature and specific sap flow are the mean values of the interval. Coarse root biomass was set to a fixed value measured in April 2021 for each pair.

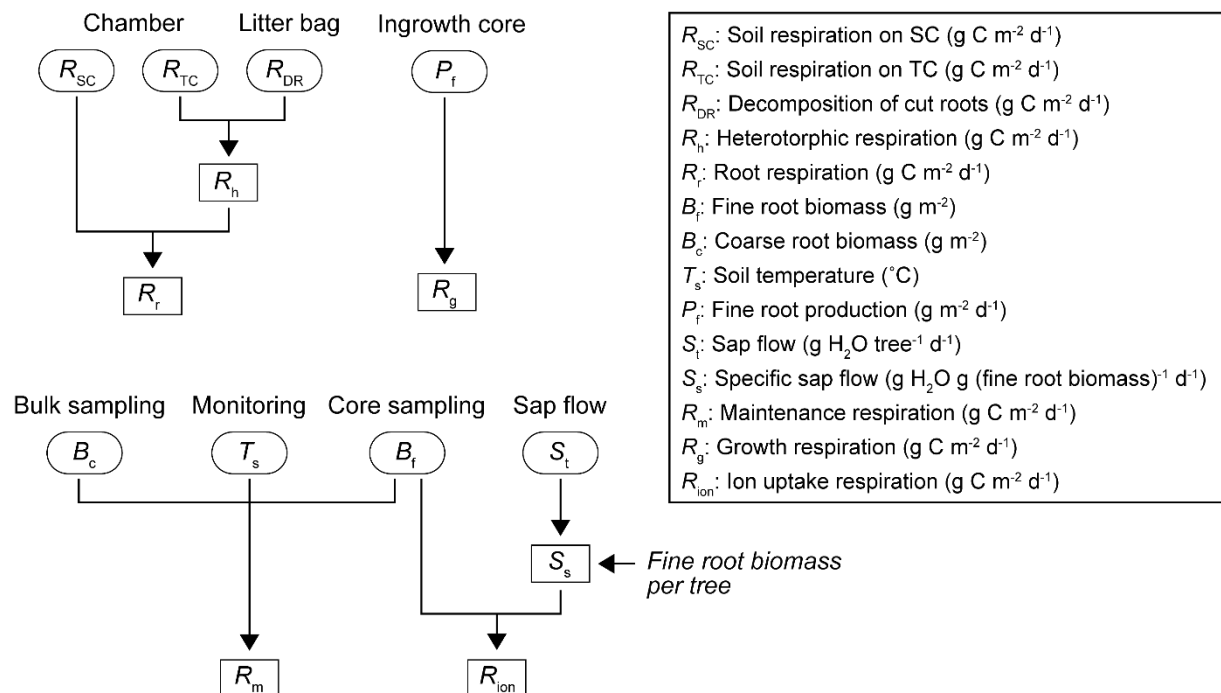


Fig. 2. Workflow of data preparation for model parameterization.

2.8. Data analysis

Student's *t*-test was applied to compare two means, assuming homoscedasticity. Two-way repeated measures analysis of variance (ANOVA) was used to test the effects of factors. For the mixed-effects model application, we used the package 'nlme' in R (Pinheiro et al. 2022). Uncertainties (SD) in the annual summation of respiration components due to model parameterization (Eqs. 1, 2, 4 and 5) were determined by a bootstrap approach ($n = 1000$), in which model parameters were randomly generated according to a normal distribution with the mean $\pm$ SE of each parameter. The uncertainties were propagated by the law of error propagation. Finally, the uncertainties due to parameterization and spatial variation were combined (combined SD) at 0.5 m and 1.0 m, respectively.

3. Results

3.1. Environmental conditions

The field experiment was conducted from September 2019 through November 2020. In the annual period from October 2019 to September 2020, the mean air temperature was 8.6°C, which was higher than its decadal mean + 1 SD, whereas the total precipitation of 1168 mm was within the range. The five-day moving average of soil temperature broadly peaked at 22–23°C from early August to mid-September and was below -1°C from late December to mid-February, with a negative peak at -3°C in mid-January (Fig. 3a). Soil moisture fluctuated between 0.1 and 0.2 m³ m⁻³ in the snowless season mainly according to precipitation but rapidly decreased in December by soil freezing and then gradually increased under snow accumulation because of the thawing of frozen soil (Fig. 3b). Sap flow showed a seasonal pattern like temperature seasonality but with a large fluctuation mainly due to variable solar radiation (Fig. 3c).

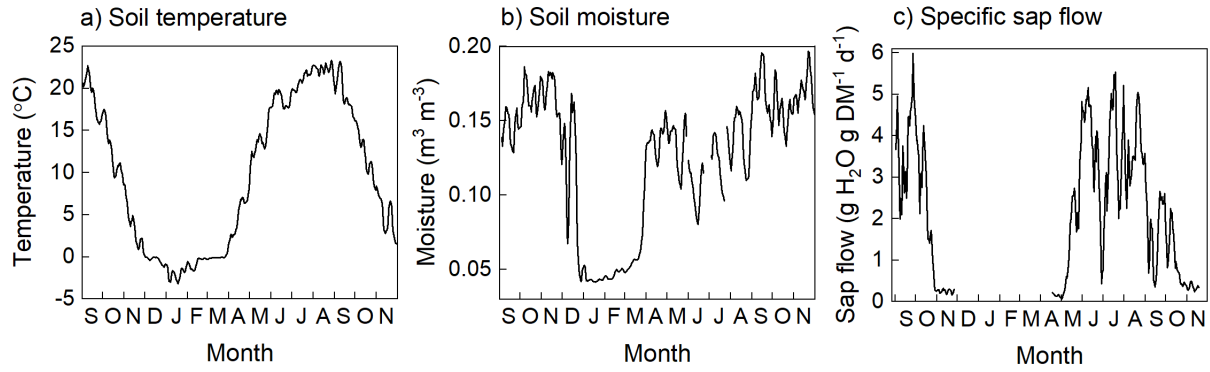


Fig. 3. Seasonal variations in daily means of soil temperature at a depth of 6 cm (a), volumetric soil moisture at a depth of 3 cm (b), and mean specific sap flow ($n = 3$) (c) from September 2019 to November 2020. Five-day moving averages are shown.

3.2. Soil CO₂ flux

Soil CO₂ flux varied seasonally, following temperature variation (Fig. 4a). Fluxes were 2.54 ± 1.53 (R_{SC}) and 1.44 ± 0.89 (R_{TC}) $\mu\text{mol m}^{-2} \text{s}^{-1}$ at 0.5 m and 2.03 ± 1.53 (R_{SC}) and 1.15 ± 0.80 (R_{TC}) $\mu\text{mol m}^{-2} \text{s}^{-1}$ at 1.0 m (mean $\pm$ SD, $n = 135$). According to two-way repeated measures ANOVA by setting the date as a block factor, R_{TC} was significantly smaller than R_{SC} at both positions ($P < 0.0001$, $n = 9$). In addition, R_{TC} was smaller than R_{SC} in each pair, with a few exceptions when R_{SC} was smaller than $1.5 \mu\text{mol m}^{-2} \text{s}^{-1}$, and a significant correlation was found between the fluxes ($P < 0.0001$) (Fig. 5). Meanwhile, a significant exponential relationship was found between soil CO₂ flux and soil temperature (Eq. 1) on each collar ($R^2 = 0.48\text{--}0.77$, $P < 0.01$). However, no significant linear or curvilinear relationship was found between temperature-normalized fluxes and soil moisture, as in our previous studies (Cui et al. 2021; Sun et al. 2020). Thus, half-hourly R_{SC} and R_{TC} were calculated for each collar from monitoring soil temperature (Fig. 3a), using each exponential equation. Both soil temperature and soil moisture measured in collars did not differ significantly between positions (0.5 m vs. 1.0 m) or treatments (SC vs. TC).

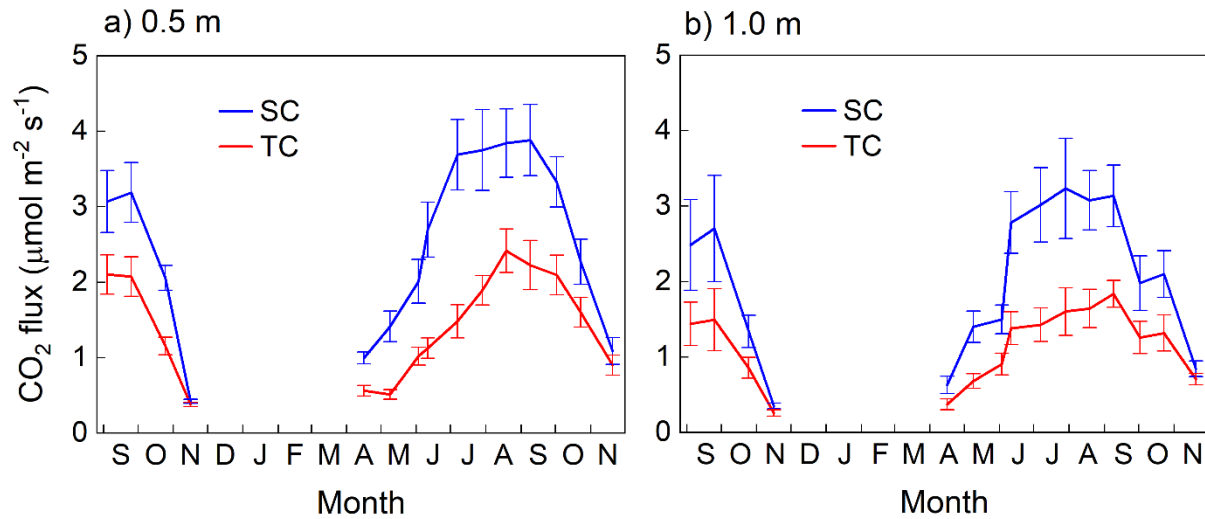


Fig. 4. Seasonal variations in soil CO₂ flux on control collars (SC) and trenched collars (TC) at 0.5 m (a) and 1.0 m (b) from September 2019 to November 2020. Vertical bars denote standard errors ($n = 9$).

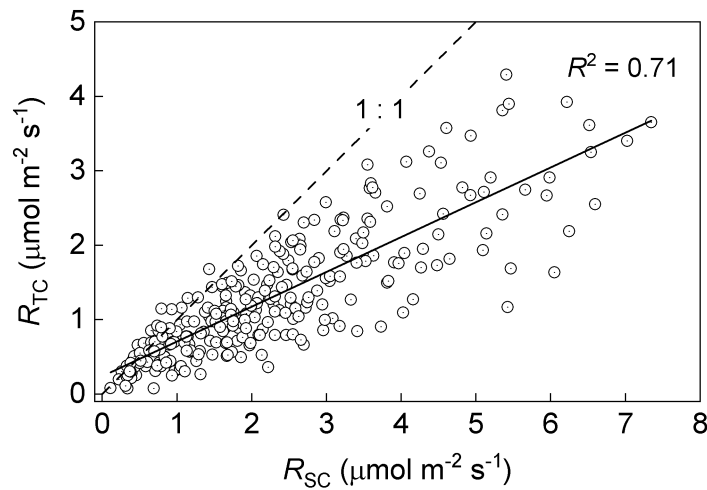


Fig. 5. Relationship between soil CO₂ fluxes measured in trenched collars (R_{TC}) and control collars (R_{SC}) of each pair. The dashed line denotes a 1:1 relationship. The solid line denotes liner regression.

In each pair of collars, daily soil respiration ($R_s = R_{SC}$) was partitioned into R_h and R_r . R_s , R_r , and R_h varied similarly to the seasonal pattern of soil temperature, because both R_{SC} and R_{TC} were calculated from soil temperature (Figs. 6a and 6b). In addition, the contribution of R_r to R_s (R_r / R_s) showed a clear seasonal variation (Fig. 6c). The R_r / R_s decreased sharply during the fall until late November and continued decreasing to 0.36–0.37 until late March under snow accumulation. After snow disappearance

in April, R_r / R_s rapidly increased and reached a broad peak of approximately 0.5 in summer from late June to early September 2020.

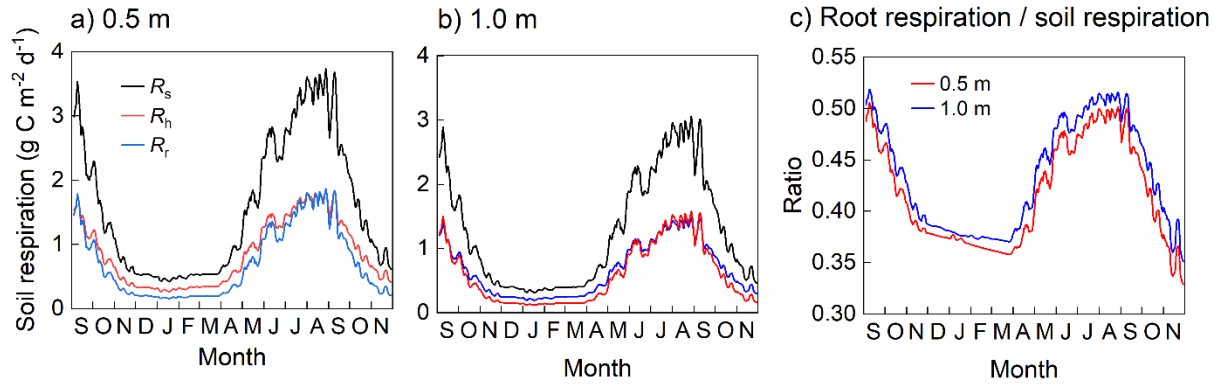


Fig. 6. Seasonal variations in mean daily soil respiration (R_s), heterotrophic respiration (R_h) and root respiration (R_r) at 0.5 m (a) and 1.0 m (b), and the ratio of root respiration and soil respiration (c) ($n=9$) from September 2019 to November 2020. Five-day moving averages are shown.

Annually, R_s was partitioned into R_h and R_r by 55% vs. 45% at 0.5 m and 53% vs. 47% at 1.0 m (Table 1). The CO₂ emissions through dead root decomposition (R_{DR_all}) accounted for 5.9% and 4.5% of R_s at 0.5 m and 1.0 m, respectively.

Table 1. Annual soil CO₂ fluxes (g C m⁻² yr⁻¹) (mean ± combined standard deviation).

Position	R_{SC} (R_s)	R_{TC}	R_{DR_f}	R_{DR_c}	R_{DR_all}	R_h	R_r
0.5 m	562 ± 123 (100)	341 ± 100 (61)	27 ± 16 (4.8)	6 ± 3 (1.1)	33 ± 16 (5.9)	308 ± 101 (55)	253 ± 159 (45)
1.0 m	447 ± 107 (100)	256 ± 74 (57)	17 ± 21 (3.8)	3 ± 2 (0.7)	20 ± 21 (4.5)	237 ± 77 (53)	210 ± 132 (47)

1) R_{SC} (R_s): soil CO₂ flux in the sampling collar (SC) or total soil respiration, R_{TC} : soil CO₂ flux in trenched collars (TC), R_{DR_f} : CO₂ emissions through the decomposition of dead fine roots, R_{DR_c} : CO₂ emissions through the decomposition of dead coarse roots, R_{DR_all} : the sum of R_{DR_f} and R_{DR_c} ; R_h : heterotrophic respiration, R_r : root respiration

2) The numbers in parentheses denote the percentages of R_{SC} at each position. The annual value was calculated as the average of the sums for the two annual periods, i.e., October 2019 through September 2020 and November 2019 through October 2020.

3.3. Biomass and production of fine roots

Seasonal variation in fine root biomass was ambiguous, despite a small peak in early summer (Fig. 7). In contrast, fine root production showed a clear seasonal pattern (Fig. 8), with a peak in June to July. In the cold season from mid-November thorough mid-April, root production was small at 0.11 ± 0.005 and $0.12 \pm 0.069 \text{ g m}^{-2} \text{ d}^{-1}$ (mean $\pm$ SE) at 0.5 m and 1.0 m, respectively. Annual values were shown in Table 2. Only the turnover rates were significantly different between the two positions ($P < 0.05$).

To compare the spatial variation of biomass and production between two scales of the collar (0.5 m $\times$ 0.5 m) and the study area (50 m $\times$ 50 m), the standard deviations (SDs) of biomass and production measured at three points within each SC were calculated and averaged for nine locations at 0.5 m and 1.0 m, respectively, for each sampling date. For all sampling dates ($n = 8$), the mean $\pm$ SD values of the averaged SD (within collar) were 47.2 ± 5.7 (0.5 m) and 37.9 ± 6.2 (1.0 m) g m^{-2} for biomass, and 0.249 ± 0.155 (0.5 m) and 0.229 ± 0.130 (1.0 m) $\text{g m}^{-2} \text{ d}^{-1}$ for production. Meanwhile, the SD of the mean biomass and production from nine individual locations (within study area) was calculated and averaged for all sampling dates to be 76.6 ± 14.4 (0.5 m) and 73.6 ± 23.8 (1.0 m) g m^{-2} for biomass, and 0.363 ± 0.210 (0.5 m) and 0.309 ± 0.190 (1.0 m) $\text{g m}^{-2} \text{ d}^{-1}$ for production. Spatial variation expressed by the mean SD was smaller within the collar than within the study area by 40–50% for biomass and 25–50% for production.

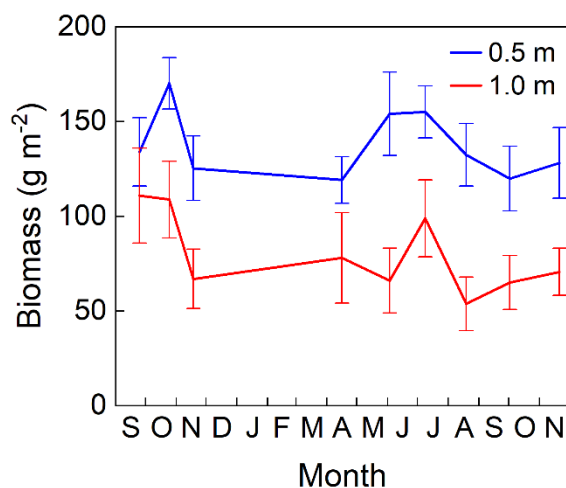


Fig. 7. Seasonal variation in fine root biomass at 0.5 m and 1.0 m ($n = 9$) from September 2019 to November 2020. Means ($\pm$ standard errors) were shown.

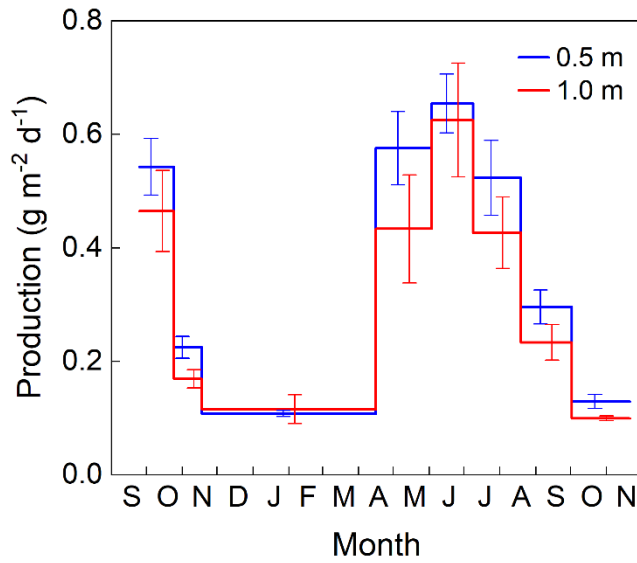


Fig. 8. Seasonal variation in fine root production at 0.5 m and 1.0 m ($n = 9$) from September 2019 to November 2020. Means ($\pm$ standard errors) were shown.

Table 2. Annual fine root dynamics (mean $\pm$ standard deviation, $n = 9$).

Position	P_f (g m ⁻² yr ⁻¹)	ΔB_f (g m ⁻² yr ⁻¹)	M_f (g m ⁻² yr ⁻¹)	Mean B_f (g m ⁻²)	Turnover rate (yr ⁻¹)*	Mean B_c (g m ⁻²)
0.5 m	116 $\pm$ 19	3 $\pm$ 48	113 $\pm$ 99	133 $\pm$ 40	0.90 $\pm$ 0.30	44 $\pm$ 42
1.0 m	103 $\pm$ 29	4 $\pm$ 26	98 $\pm$ 39	78 $\pm$ 43	1.68 $\pm$ 0.84	20 $\pm$ 26

* $P < 0.05$, between positions by two-sided t test

1) P_f : fine root production, ΔB_f : annual difference in fine root biomass, M_f : fine root mortality, B_f : fine root biomass, B_c : coarse root biomass

2) The annual value was calculated as the average of the sums for the two annual periods, i.e., October 2019 through September 2020 and November 2019 through October 2020.

3.4. Partitioning of root respiration

Both models (Eqs. 4 and 5) were significantly parameterized using the dataset ($P < 0.0001$, $R^2 = 0.57$ – 0.59) (Table 3). All parameters were determined to be statistically significant ($P < 0.012$). Parameter d , the maintenance respiration of the unit biomass at 0°C, was almost identical in the two models. The Q_{10} values calculated from the parameter f were 2.34 and 2.61 in Models 1 and 2, respectively. $d \cdot \exp(f \cdot T_s)$ in the models, corresponding to the maintenance respiration coefficient (m) in Eq. 3, was 0.0013 at 5°C, 0.0030 at 15°C, and 0.0070 g C g DM⁻¹ d⁻¹ at 25°C in Model 1 and 0.0013 at

5°C, 0.0034 at 15°C, and 0.0089 g C g DM⁻¹ d⁻¹ at 25°C in Model 2. The coefficients were the same at 5°C across the models, but 12% and 21% lower in Model 1 at 15°C and 25°C, respectively, because of a lower temperature coefficient. In addition, the coefficient of growth respiration (*g*) was 22% lower in Model 1. The decreases in *f* and *g* in Model 1 were compensated for by the coefficient of ion uptake respiration (*u*).

Table 3. Model parameters ($\pm$ standard error).

	Model 1		Model 2	
	Parameter	<i>P</i>	Parameter	<i>P</i>
Growth respiration coefficient (<i>g</i>) (g C g DM ⁻¹)	0.57 $\pm$ 0.14	0.0001	0.73 $\pm$ 0.13	< 0.0001
Maintenance respiration coefficient at 0°C (<i>d</i>) (g C g DM ⁻¹ d ⁻¹)	0.00084 $\pm$ 0.00027	0.0019	0.00081 $\pm$ 0.00025	0.0015
Temperature coefficient (<i>f</i>) (°C ⁻¹)	0.085 $\pm$ 0.015	< 0.0001	0.096 $\pm$ 0.014	< 0.0001
Ion uptake respiration coefficient (<i>u</i>) (g C g H ₂ O ⁻¹)	0.00045 $\pm$ 0.00018	0.012		
Adjusted <i>R</i> ²	0.59		0.57	
<i>P</i>	< 0.0001		< 0.0001	

Annual *R_r* was partitioned into fine root *R_g*, fine root *R_m* (*R_{m_f}*) and coarse root *R_m* (*R_{m_c}*), and *R_{ion}* using the fitting parameters (Table 4). The results from the models were overestimated by 7% at 0.5 m and underestimated by 10% at 1.0 m in total (Sum / *R_r*). *R_{m_f}* contributed the most to *R_r* (46–51%), followed by *R_g* (26–40%). *R_g*, *R_{m_f}*, and *R_{m_c}* were smaller in Model 1 by 19, 8, and 5 g C m⁻² yr⁻¹, respectively, at 0.5 m and by 16, 5, and 2 g C m⁻² yr⁻¹, respectively, at 1.0 m. *R_g* showed the largest difference between the two models, suggesting that the majority of *R_{ion}* was lumped together with *R_g* in Model 2. Fine root *R_r* (Sum – *R_{m_c}*) accounted for 86–90% and 84–89% of the total *R_r* (Sum) in Models 1 and Model 2, respectively.

Table 4. Annual sums of root respiration components (g C m⁻² yr⁻¹) (mean $\pm$ combined standard deviation).

Model 1						
Position	<i>R_r</i>	<i>R_g</i>	<i>R_{m_f}</i>	<i>R_{m_c}</i>	<i>R_{ion}</i>	Sum
0.5 m	253 $\pm$ 227 (93)	70 $\pm$ 23 (26)	130 $\pm$ 66 (48)	39 $\pm$ 45 (14)	32 $\pm$ 17 (12)	271 $\pm$ 93 (100)

1.0 m	210 ± 151 (112)	60 ± 32 (32)	87 ± 74 (46)	18 ± 44 (10)	23 ± 17 (12)	188 ± 122 (100)
Model 2						
Position	R_r	R_g	R_{m_f}	R_{m_c}	Sum	
0.5 m	253 ± 227 (93)	89 ± 25 (33)	138 ± 76 (51)	44 ± 51 (16)	271 ± 99 (100)	
1.0 m	210 ± 151 (112)	76 ± 39 (40)	92 ± 83 (49)	20 ± 49 (11)	188 ± 124 (100)	

1) R_r : root respiration (from Table 2), R_g : growth respiration, R_{m_f} : fine root maintenance respiration, R_{m_c} : coarse root maintenance respiration, R_{ion} : ion uptake respiration, Sum: $R_g + R_{m_f} + R_{m_c} + R_{ion}$ (Model 1) or $R_g + R_{m_f} + R_{m_c}$ (Model 2)

2) The numbers in parentheses denote percentages of the sum. The annual value was calculated as the average of the sums for the two annual periods, e.g., October 2019 to September 2020 and November 2019 to October 2020.

The respiration components varied seasonally with different peak periods. Even though Fig. 9 shows the result at 0.5 m, the seasonal variation was similar at 1.0 m. R_g peaked from June to July in proportion to root production (Fig. 8), whereas R_{m_f} peaked from July to August owing to its positive relationship with soil temperature (Fig. 3a) and fine root biomass (Fig. 7). In addition, R_{m_c} peaked following the temperature variation under the assumption of no seasonality in coarse root biomass. R_{ion} peaked from July to August according to the seasonality of fine root biomass and sap flow (Fig. 3c) and was zero in winter because of defoliation. The relative contribution of each respiration component to R_r also showed seasonality (Figs. 9c and 9d). Seasonal variation in R_g contribution was relatively large, with a peak from April to May at 0.40 (Model 1) and 0.49 (Model 2) and reached a minimum of 0.16 (Model 1) and 0.20 (Model 2) from August to September.

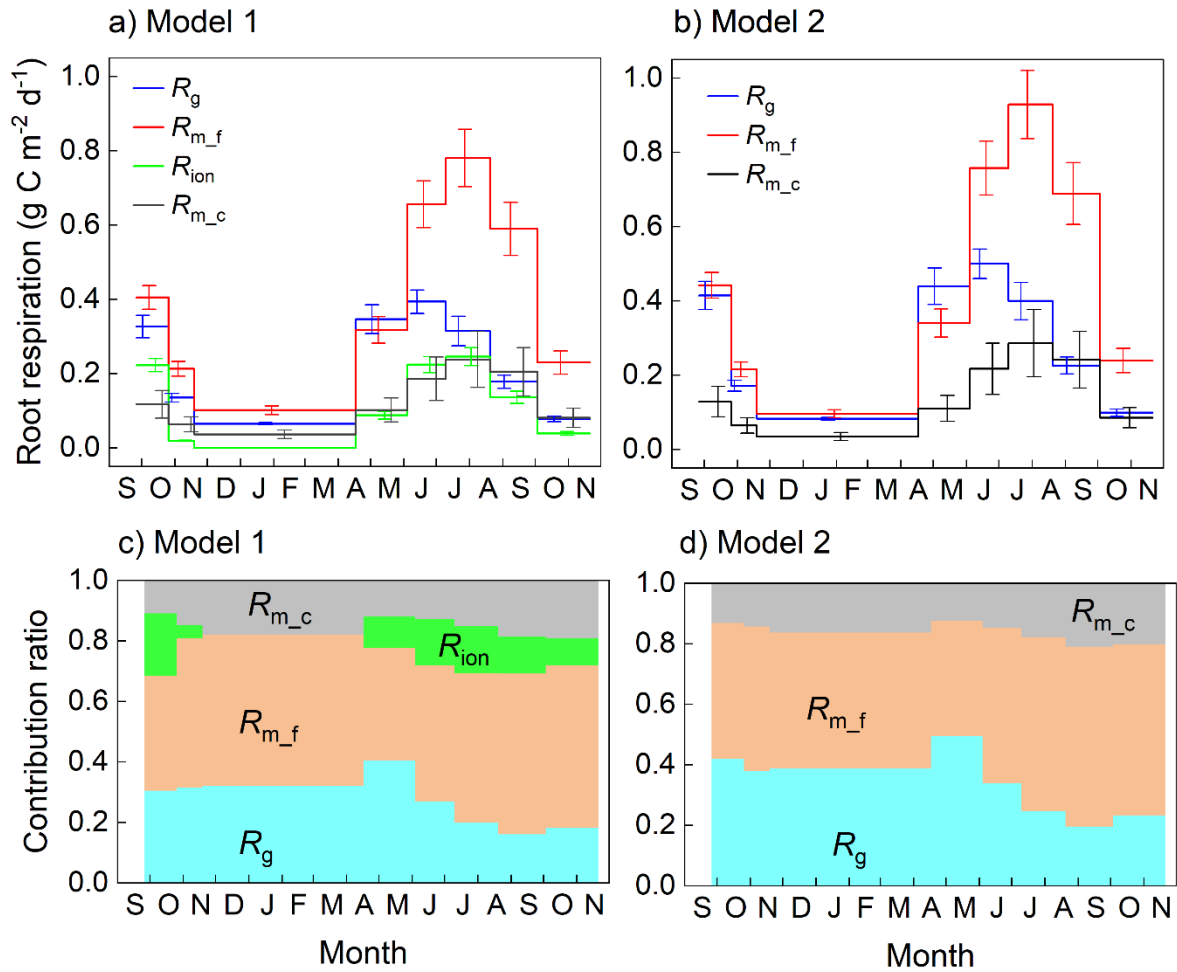


Fig. 9. Seasonal variations in fine root growth respiration (R_g), fine root maintenance respiration (R_{m_f}), coarse root maintenance respiration (R_{m_c}), and ion uptake respiration (R_{ion}) at 0.5 m estimated by Model 1 (a) and Model 2 (b) from September 2019 to November 2020. Means ($\pm$ standard errors, $n = 9$) were shown. Contribution ratios of the respiration components from Model 1 (c) and Model 2 (d) are also shown.

4. Discussion

4.1. Partitioning of soil respiration

The CO₂ emission (R_{DR}) from the decomposition of dead roots caused by trenching was estimated using the decay constant from root litterbag experiments, which were reported to underestimate fine root decomposition by 10–20% because of rhizosphere disturbance (Dornbush et al. 2002). However, the effect of the underestimation would be limited, because annual fine root decomposition was small, accounting for 4–5% of total soil respiration (Table 1). Meanwhile, trenching would have decreased R_h

because of the lack of litter supply from the living roots. In addition, no priming effect of labile exudate from roots on microbial decomposition was expected (Kuzyakov 2010). Thus, R_h was potentially underestimated, although the underestimation could have been limited because of low root density, which was much lower at 78–133 g m⁻² (Table 2) than the average density (505 g m⁻²) of temperate deciduous forests (Finér et al. 2011a). Although soil moisture was expected to increase in TC because of no water uptake by roots (Subke et al. 2006), there was no difference between SC and TC, probably because of the low root density.

Periodically measured CO₂ fluxes were extrapolated to sequential data for each collar using an exponential equation (Eq. 1) from the continuously measured soil temperature. It is common to determine representative R_h and R_r values from the spatial averages of soil CO₂ fluxes in a study area because of their large spatial variations (Sun et al. 2020). However, we determined R_h and R_r in each pair of collars to ensure the spatial variation of R_r based on the assumption that R_h values on the two collars (SC and TC) in each pair were similar. The assumption was because the two collars were closely installed (0.3–0.4 m apart) and equally distant from each isolated larch stem, suggesting that SOM, leaf litter accumulation, and root density were at a similar level in the two neighboring collars. We confirmed that soil CO₂ fluxes in the two collars were similar before trenching, and soil C concentration decreased with increasing distance from tree stems (Cui et al. 2021). The significant linear relationship between R_{TC} and R_{SC} (Fig. 4) supports this assumption. In addition, the result that spatial variation in fine root biomass was considerably smaller on a collar scale than on the scale of the study area suggests a similar soil condition in each collar pair.

Compared to our previous study conducted two years ago (Cui et al. 2021), annual R_h slightly decreased by 2% at 0.5 m and 9% at 1.0 m, but annual R_r increased by 40% at 0.5 m and 412% at 1.0 m. The R_r increase was due to tree growth; fine root biomass increased twofold at 0.5 m and sixfold at 1.0 m in two years (Table 2). The annual contribution of R_r to R_s was 45–47% (Table 1), which is comparable to the results of meta-analyses (Hanson et al. 2000; Subke et al. 2006).

4.2 Partitioning of root respiration

Soil CO₂ flux and fine root dynamics were simultaneously measured within the same collar (SC) to

minimize spatial mismatches. Although soil core sampling disturbed soil conditions, the disturbance would be insignificant to CO₂ flux (Sun et al. 2020) because the total pit area caused by eight-time core samplings was only 4.3% in each SC, and sampling intervals were more than a month.

We applied the ingrowth core method to measure root production. The method tends to underestimate root production compared to the minirhizotron method which potentially yields more reliable estimates (Addo-Danso et al. 2016; Finér et al. 2011b; Hendricks et al. 2006). The annual fine root production was 103–116 g m⁻² yr⁻¹ in this study (Table 2), which was much smaller than the average (337 g m⁻² yr⁻¹) of temperate forests (Finér et al. 2011b). However, turnover rates (0.90–1.68 yr⁻¹, Table 2) were comparable to those of temperate forests (Brunner et al. 2013; Finér et al. 2011b). Thus, our results might not have been underestimated, because ingrowth cores were sampled at relatively short intervals of 5–7 weeks using thinner cores with a diameter of 2.3 cm. Short intervals probably enhanced root production in conditions of low root competition and less decomposition, and thin cores could minimize the delayed root recolonization (Hertel and Leuschner 2002; Li et al. 2013).

All parameters were significantly determined in the models, whereas the models explained R_r variations by only 57–59% (Table 3). We excluded coarse root growth from R_g . (Wieser and Bahn 2004) reported that R_m of coarse roots (R_{m_c}) accounted for 73–83% of total coarse root respiration. Accordingly, coarse root growth respiration was roughly estimated from R_{m_c} (Table 4) to be 8–12 g C m⁻² yr⁻¹ at 0.5 m and 4–5 g C m⁻² yr⁻¹ at 1.0 m, which accounted for 3–4% and 2–3% of total R_r , respectively. In addition, we set coarse root biomass at the measurements in April 2021 after the experiment, ignoring its phenological variation. Coarse root biomass would have increased during the growing season because tree surveys indicated an annual growth in the study site from 2019 to 2020. Thus, the root biomass in April 2021 would be close to the maximum for the experimental period, and consequently, R_{m_c} was probably overestimated to some extent. The parameters related to R_m (d and f) were common for fine and coarse roots in the models to make the model simpler and nonlinear fitting converge. However, the parameters might be smaller for coarse roots, because N concentration was significantly lower in coarse roots than in fine roots (7.10 ± 1.25 vs. 9.22 ± 0.83 mg g⁻¹); R_m is positively related to protein content (Lambers et al. 2008). Because trenching was conducted using PVC boards, calculated R_r consists of respiration by roots alone and rhizomicrobial and mycorrhizal respirations

(Moyano et al. 2009). Rhizomicrobial and mycorrhizal respirations were not incorporated in the models. However, these respirations were probably stimulated by the translocation of photosynthate to roots, depending on root-derived compounds. Therefore, rhizomicrobial and mycorrhizal respirations should be included in R_g estimated from root production. In a mature Japanese larch forest, the contribution of respiration by roots plus rhizomicrobial respiration to R_s and mycorrhizal respiration to R_s were 42% and 6%, respectively, during the growing season (Makita et al. 2021).

The annual partitioning of R_r to four (Model 1) or three components (Model 2) at 0.5 m and 1.0 m were similar (Table 4), although the absolute values were smaller at 1.0 m. R_{m_f} accounted for the majority (46–51%), followed by R_g (26–40%) and R_{m_c} (10–16%) or R_{ion} (12%). In comparison between the two models, the largest difference was found in R_g , suggesting that the majority of R_{ion} was lumped together with R_g when R_{ion} was not considered, which is supported by the strong correlation between plant growth and ion uptake (Lambers et al. 2008).

Fine root growth even in the snow season (Radville et al. 2016) was suggested by ingrowth core sampling in mid-April (Fig. 8). Thus, relatively high contribution of R_g to R_r was possible during winter (Fig. 9). However, the root production measured in mid-April may have resulted from rapid growth after snow disappearance in early April (Fig. 3a) (McCormack et al. 2015b). The contribution of R_g showed a clear seasonal variation, with a small peak in mid-spring and a minimum in early fall, because of the time lag of seasonal variation between fine root production (Fig. 8) and soil temperature (Fig. 3a). R_g exceeded R_{m_f} only in mid-spring. Although R_g and R_{m_f} were larger in this study compared to our previous study (Cui et al. 2021) because of large increase in fine root production and biomass in two years, the annual contribution ratios of R_g , R_{m_f} , and R_{m_c} to R_r was unchanged at 0.5 m. However, the results of this study (Table 4) were much different from those of a field study in evergreen and deciduous plantations (George et al. 2003), in which fine root R_m accounted for 86–92% of R_r and the contribution of R_{ion} was only 4%. The difference is mainly attributable to different methods as described below.

The coefficients of Models 1 and 2 are compared with those from other studies (Table 5). To convert the unit of O_2 to C of CO_2 , the respiration quotient was assumed to be one. Compared with (Cui et al. 2021), the g for growth respiration of the same model (Model 2) rose by 26% in this study, whereas the m for maintenance respiration was almost the same. Although R^2 was almost the same, the standard

errors of the parameters (g , d , and f) were smaller in this study (Table 3) mainly because of more replications. This and our previous studies (Cui et al. 2021; Sun et al. 2020) showed a larger g than those of the other tree species (*Quercus suber*, *Eucalyptus* sp., *Pinus Taeda*, and *Liquidambar styraciflua*) The difference is mainly due to different methods and different experimental conditions. For *Quercus suber* and *Eucalyptus* sp., the g for all roots, including coarse roots, was determined from relative growth rates of hydroponically cultured seedlings or cuttings in a controlled environment (Lambers et al. 2008; Thongo M'Bou et al. 2010). In addition, the fine root g for *Pinus taeda* and *Liquidambar styraciflua* was determined from a laboratory calorimetric experiment (George et al. 2003). The carbon cost of producing new root biomass was reported to be 0.536 g C g DM⁻¹ on average (Chapin et al. 2011; Poorter 1994). From the carbon cost, carbon consumed through growth respiration was roughly estimated to be 0.11 g C g DM⁻¹ under the assumption that respiration cost accounts for 20% of the total cost (Chapin et al. 2011). Our g derived from field experiments were consistently higher than the respiration cost from laboratory experiments. The large difference in g suggests that more energy is necessary for fine root production in the field. In contrast, the m was relatively stable among tree species, except for smaller m for mature spruce.

Table 5. Comparison of model parameters.

Plant species	Experiment	Growth respiration coefficient (g) (g C g DM ⁻¹)	Maintenance respiration coefficient (m) (mg C g DM ⁻¹ d ⁻¹)	Temperature (°C)	Ion Uptake respiration (R_{ion})	Reference
<i>Dactylis glomerta</i>	Laboratory	0.13	27	No data	No	Lambers <i>et al.</i> , 2008
<i>Festuca ovina</i>		0.23	22			
<i>Quercus suber</i>		0.14	6.2			
<i>Triticum aestivum</i>		0.22	23			
<i>Eucalyptus</i> sp.	Greenhouse	0.06	10	22	No	Thongo M'Bou <i>et al.</i> , 2010
<i>Pinus taeda</i>	Forest	0.061	9.2	25	Yes	George <i>et al.</i> , 2003
<i>Liquidambar styraciflua</i>		0.070	11			
Mature larch	Forest	0.32	5.7	22	No	Sun <i>et al.</i> , 2020
			8.2	25		
Mature spruce			2.8	22		
		0.24	3.7	25		

Young larch	Forest	0.58	6.8 8.1	22 25	No	Cui <i>et al.</i> , 2021
Young larch	Forest	0.57	5.5 7.0	22 25	Yes	This study (Model 1)
Young larch	Forest	0.73	6.7 8.9	22 25	No	This study (Model 2)

We conducted a year-round experiment by increasing spatial replication and incorporating R_{ion} in a young larch-dominated forest regrowing on bare ground. The experimental conditions were considerably simplified because of an almost pure stand, the same tree age, low tree density, little ground vegetation, little litter accumulation, and a small amount of SOM. From the experiments in this study and (Cui et al. 2021), we derived similar parameters (g , d , and f) for partitioning. The Q_{10} of maintenance respiration (2.33–2.61) was reasonable. Therefore, we believe that root respiration was partitioned with a certain level of reliability and the approach was robust under relatively simple conditions. However, we should note that our results were yielded in an unnaturally simple field condition, because the soil without the organic layer was not typical among forest soils. Moreover, our growth respiration parameter (g) resulting from only field data was much higher than those from laboratory experiments. To further improve our understanding of root respiration in the field, more experimental data should be accumulated.

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Statements and Declarations

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